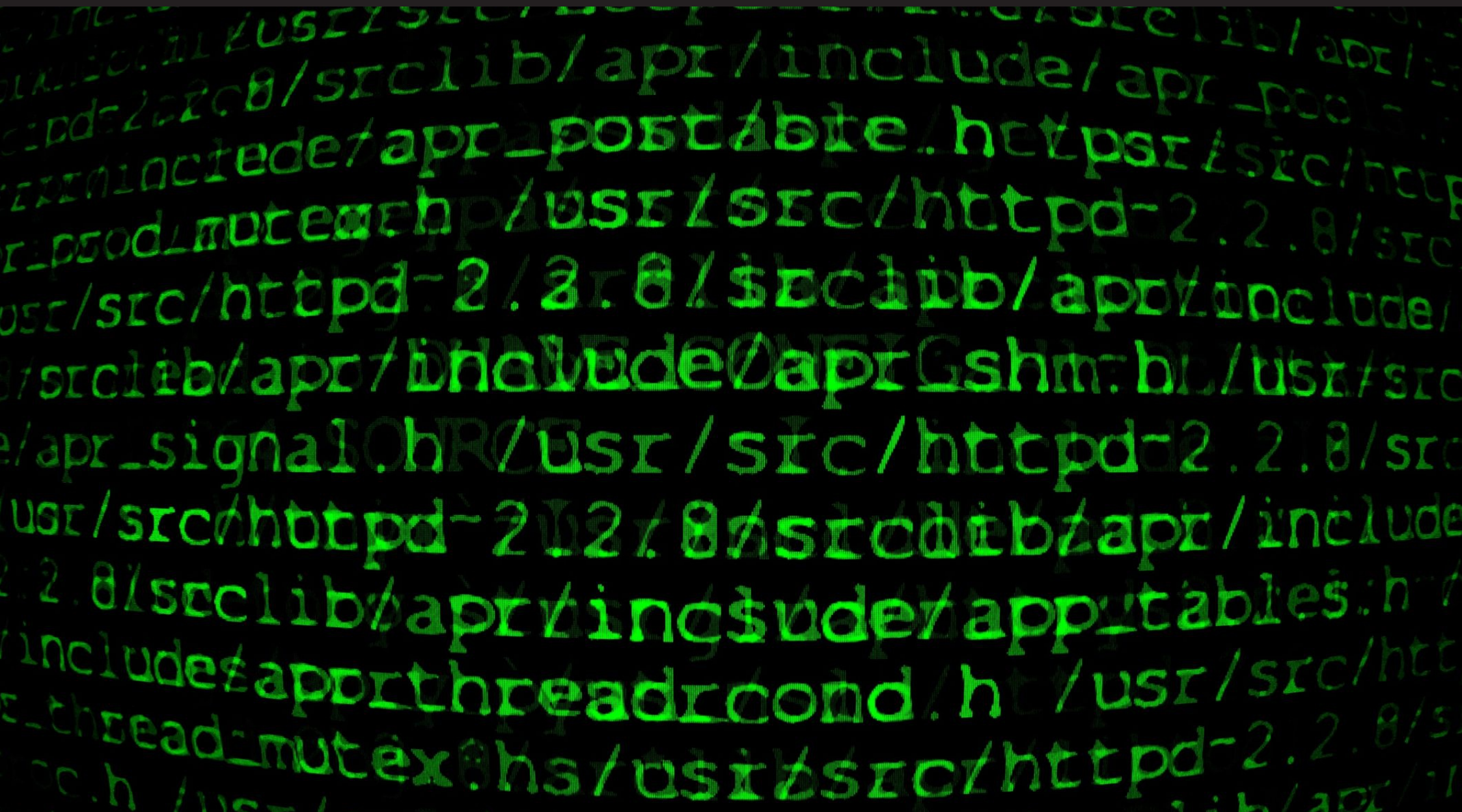


ECODINGT HEM

ARINEG EOLOGICALR ECORDD

Ndyr Idgwella



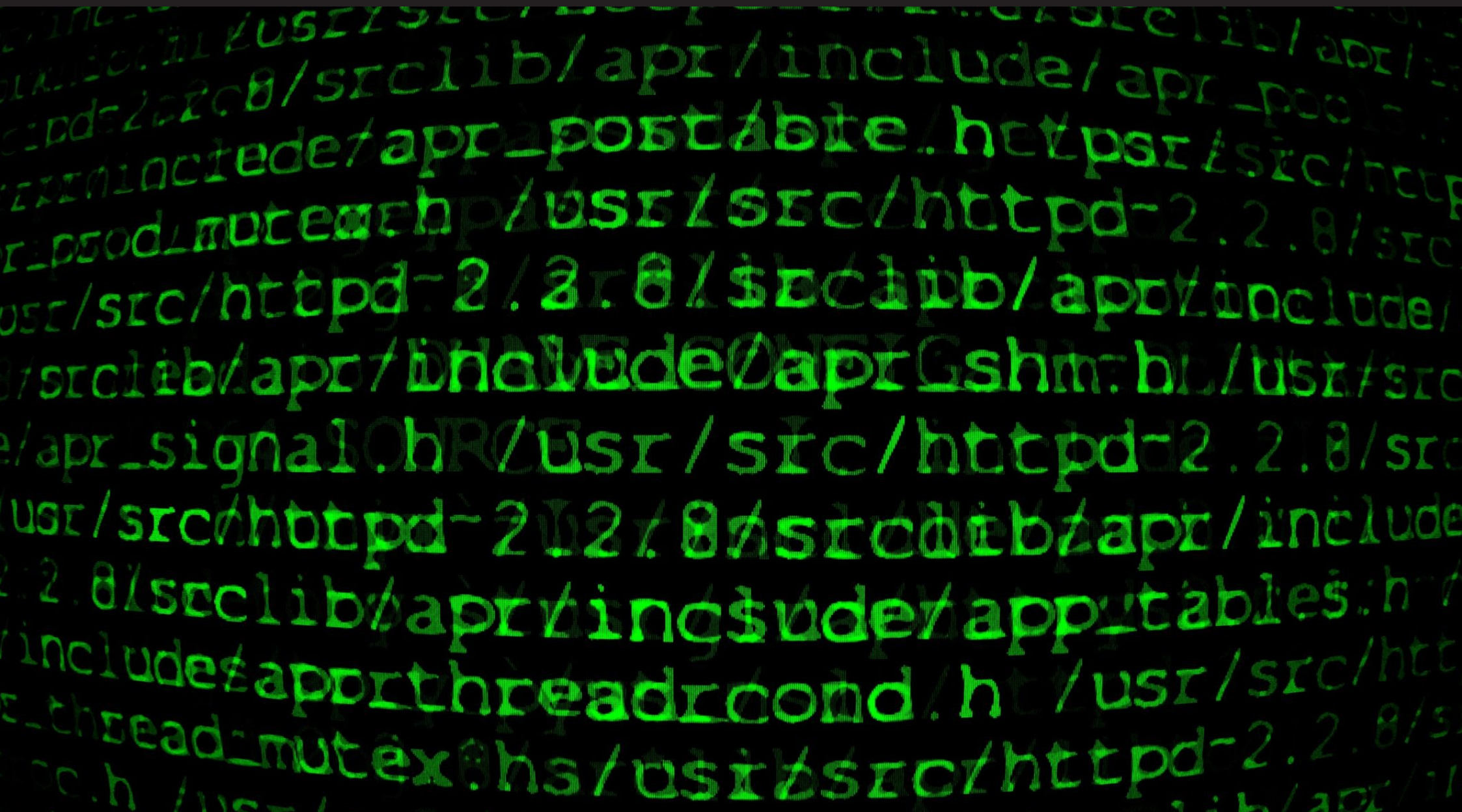


ECODINGT HEM

ARINEG E

ALR ECORDD

Ndyr Idgwella



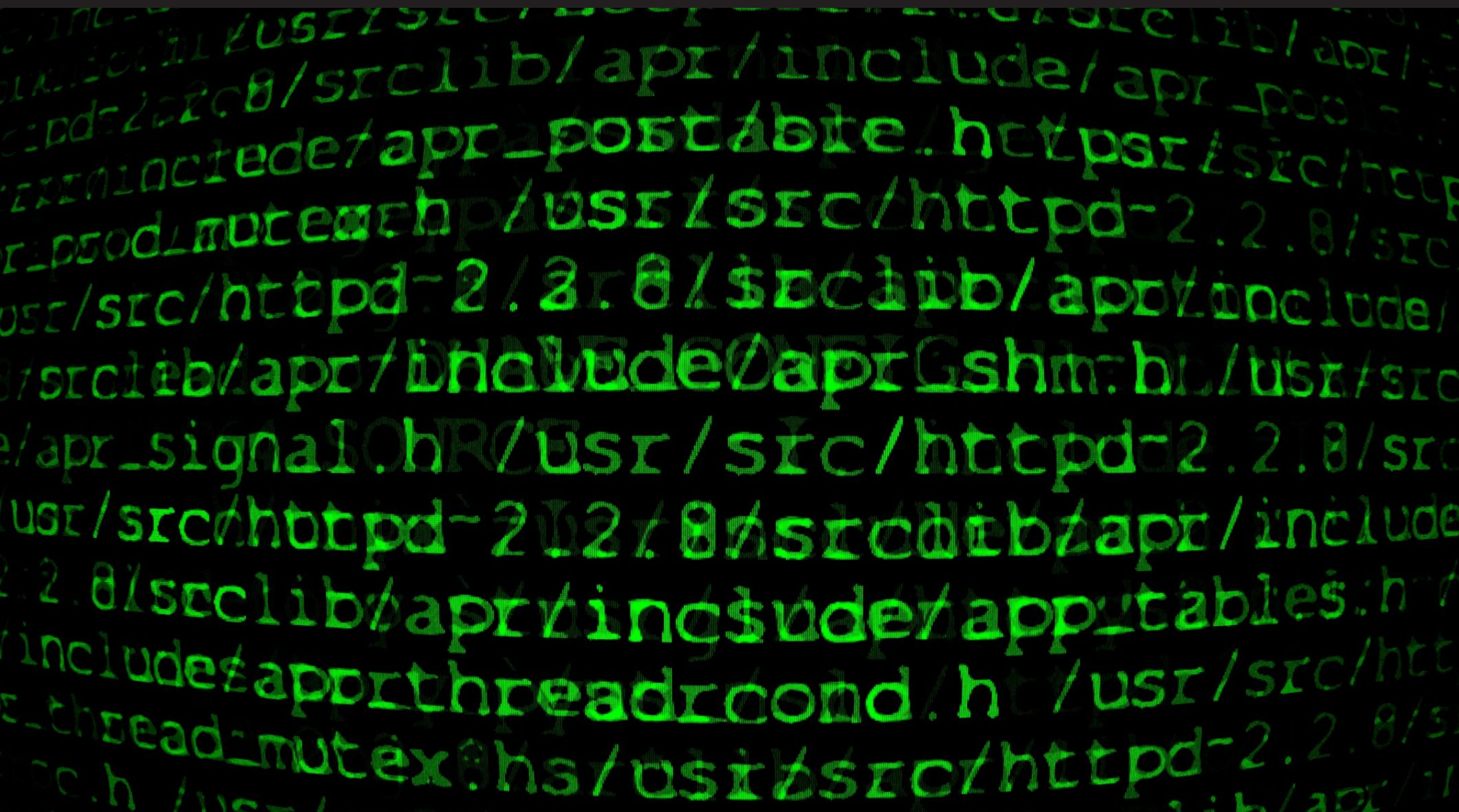


ΕΧΟΔΙΝΓΤ ΗΕΜ

ΑΡΙΝΕΓ Ε

ΑΛΡ ΕΧΟΡΑΔ

Νδψρ Ιδγωελλα

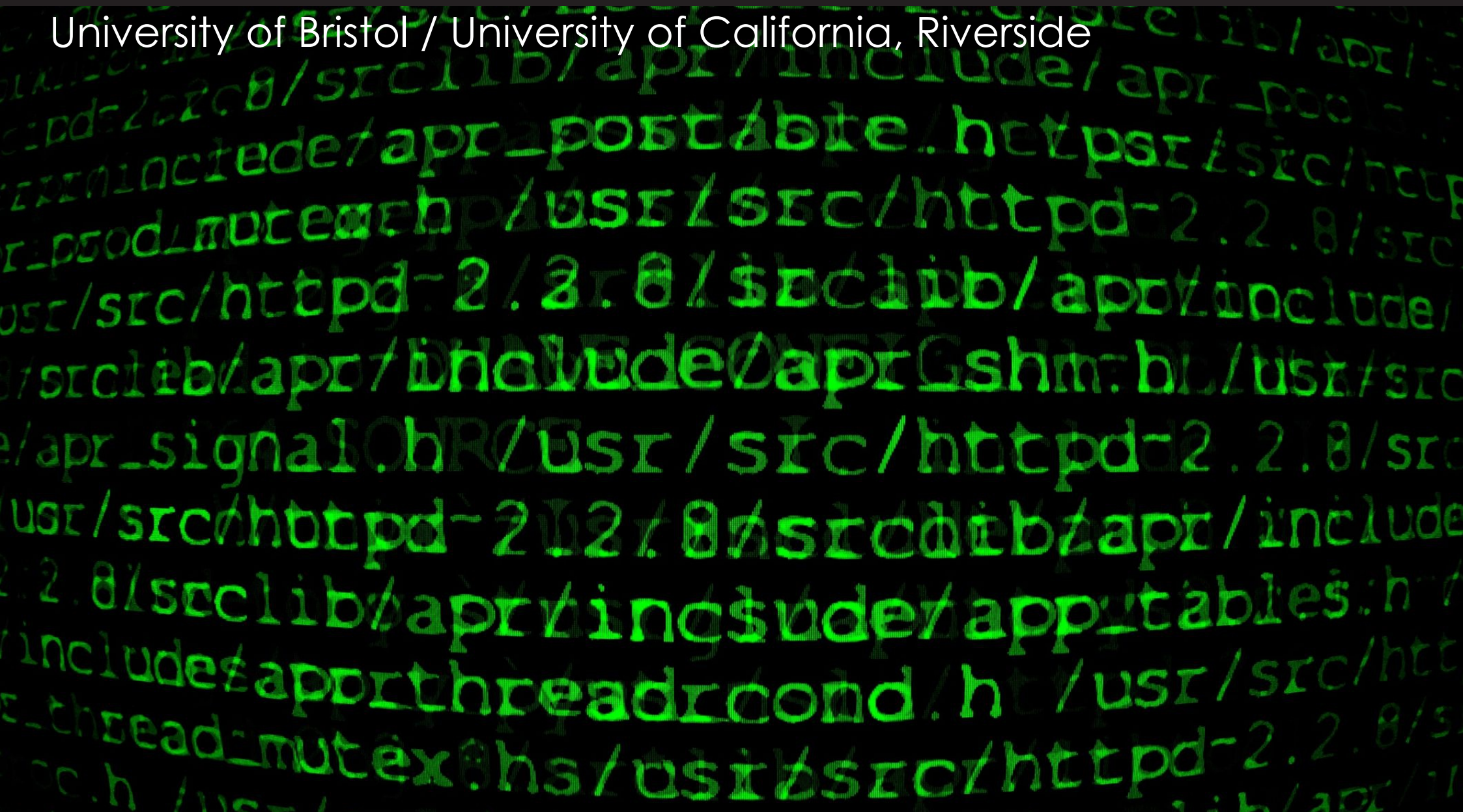




# DECODING THE MARINE GEOLOGICAL RECORD

Andy Ridgwell

University of Bristol / University of California, Riverside















How much carbon?  
(=> infer climate,  
ecosystem sensitivity etc.)







By simple mass balance, the size of an isotopic excursion is:

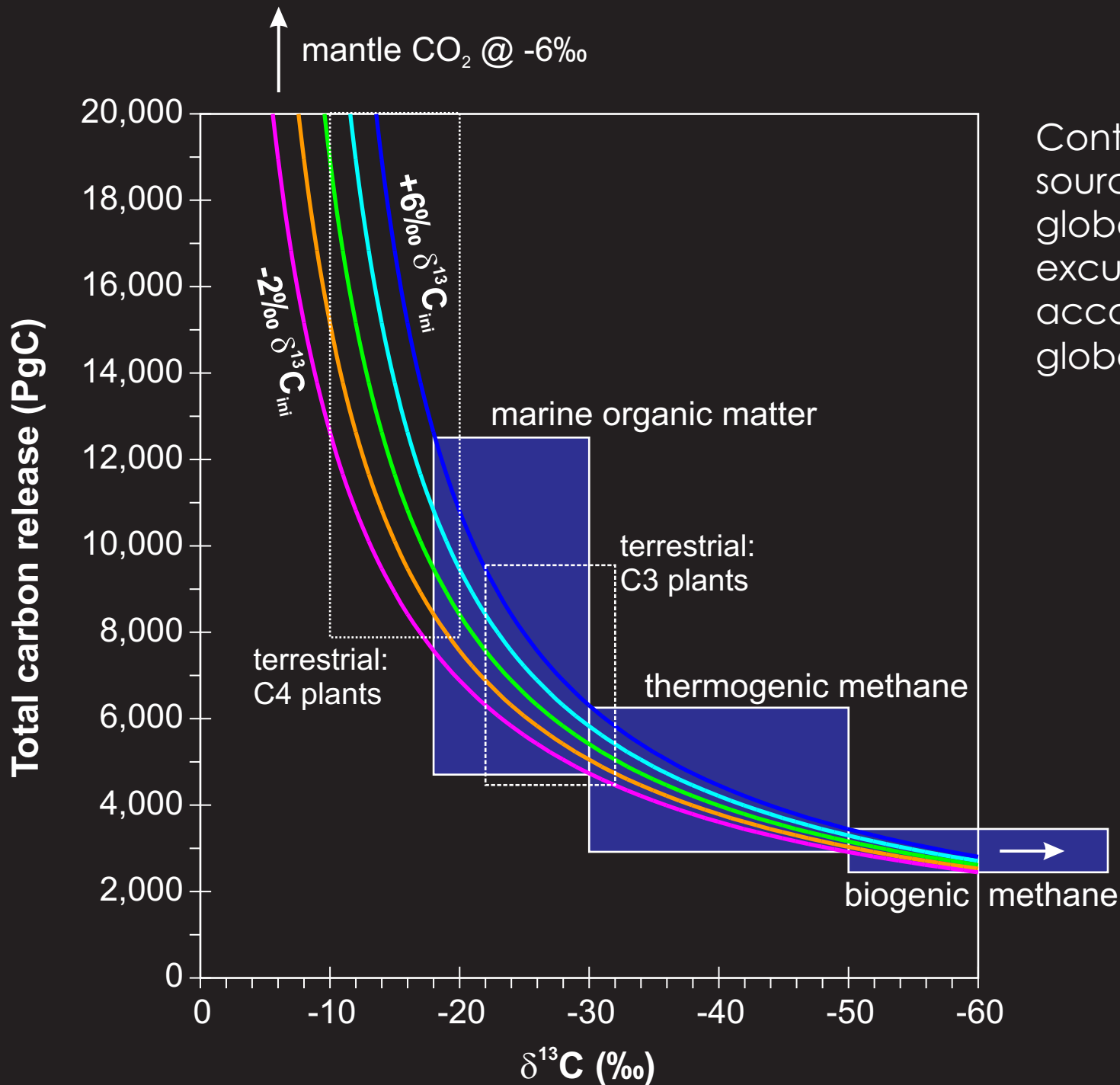
$$\Delta\delta^{13}\text{C} \sim \delta^{13}\text{C}_{\text{new}} \times \Delta M_{\text{new}} / ( M_{\text{old}} + \Delta M_{\text{new}} )$$

where  $\Delta M_{\text{new}}$  is the mass added,  $\delta^{13}\text{C}_{\text{new}}$  is its isotopic signature, and  $M_{\text{old}}$  is the original total mass of 'exchangeable' carbon. Or:

$$\Delta M_{\text{new}} \sim \Delta\delta^{13}\text{C} \times M_{\text{old}} / ( \delta^{13}\text{C}_{\text{old}} + \Delta\delta^{13}\text{C} - \delta^{13}\text{C}_{\text{new}} )$$



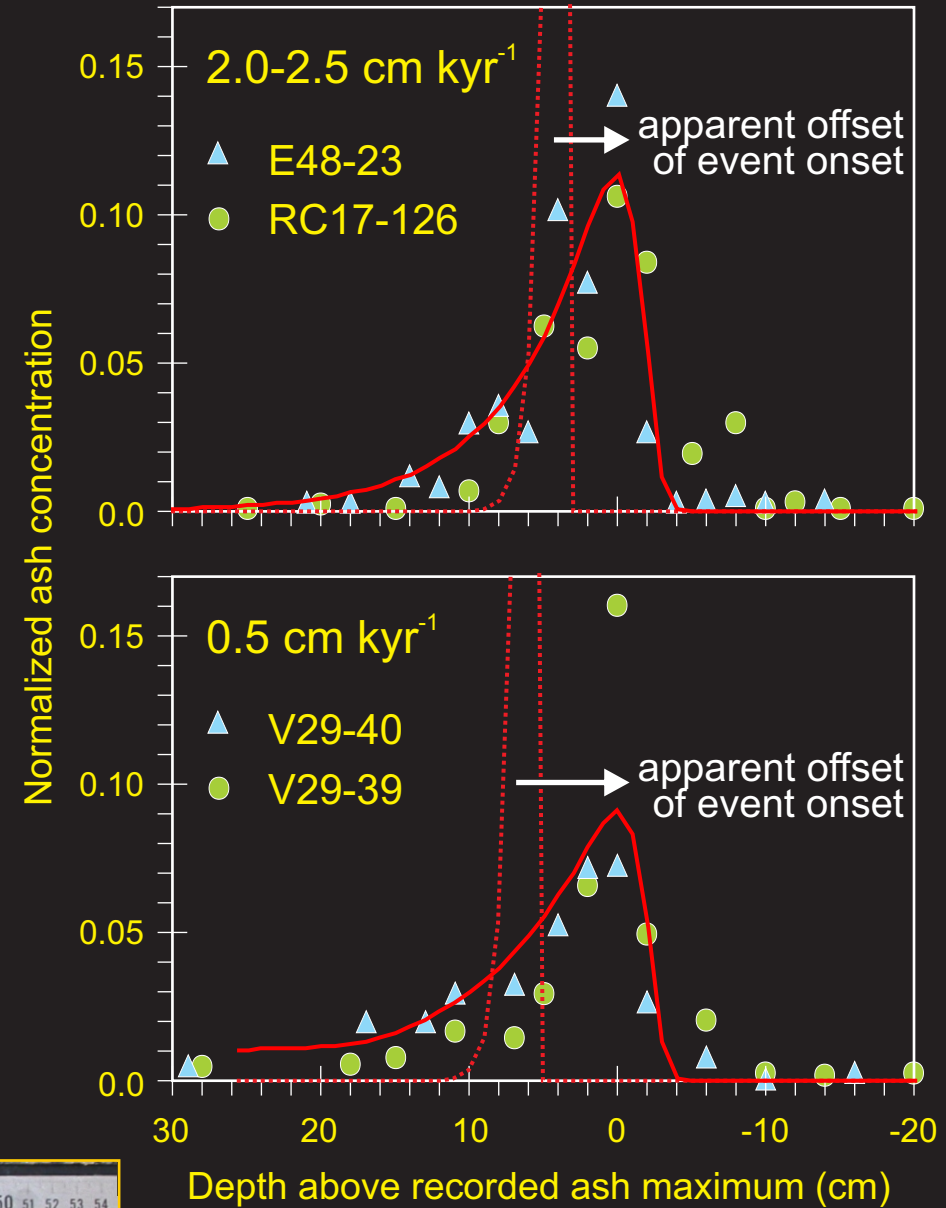
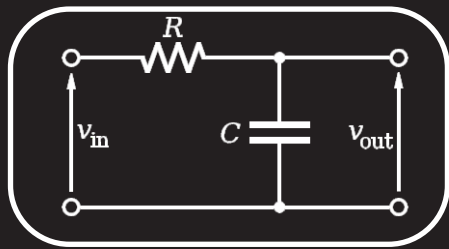




Contours of carbon release vs. source isotopic signature for a global  $-4\text{‰}$  carbon isotopic excursion. Contours differ according to the initial mean global  $\delta^{13}\text{C}$ .



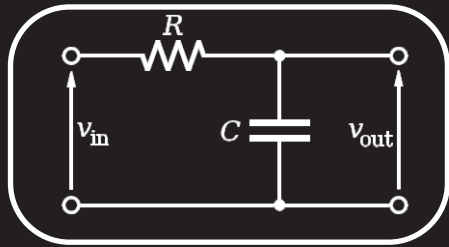
mixing  
(bioturbation)



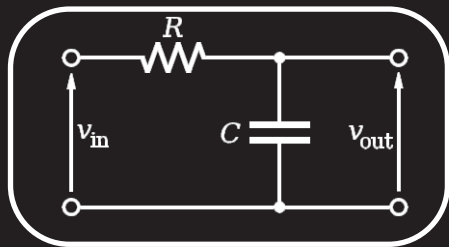




dissolution  
(preservation)

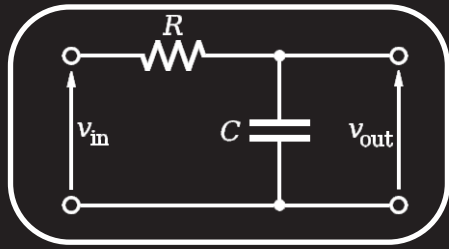


mixing  
(bioturbation)

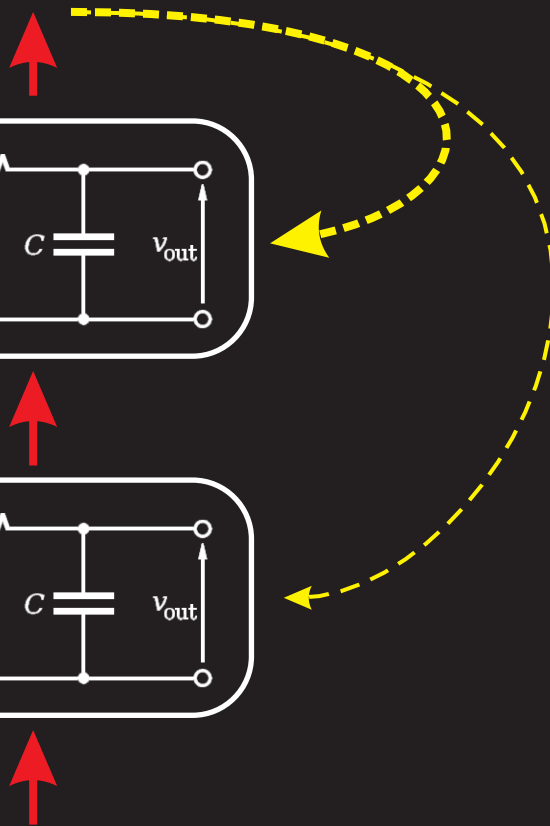
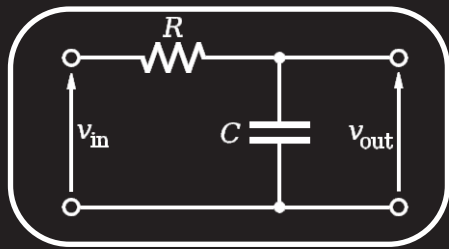




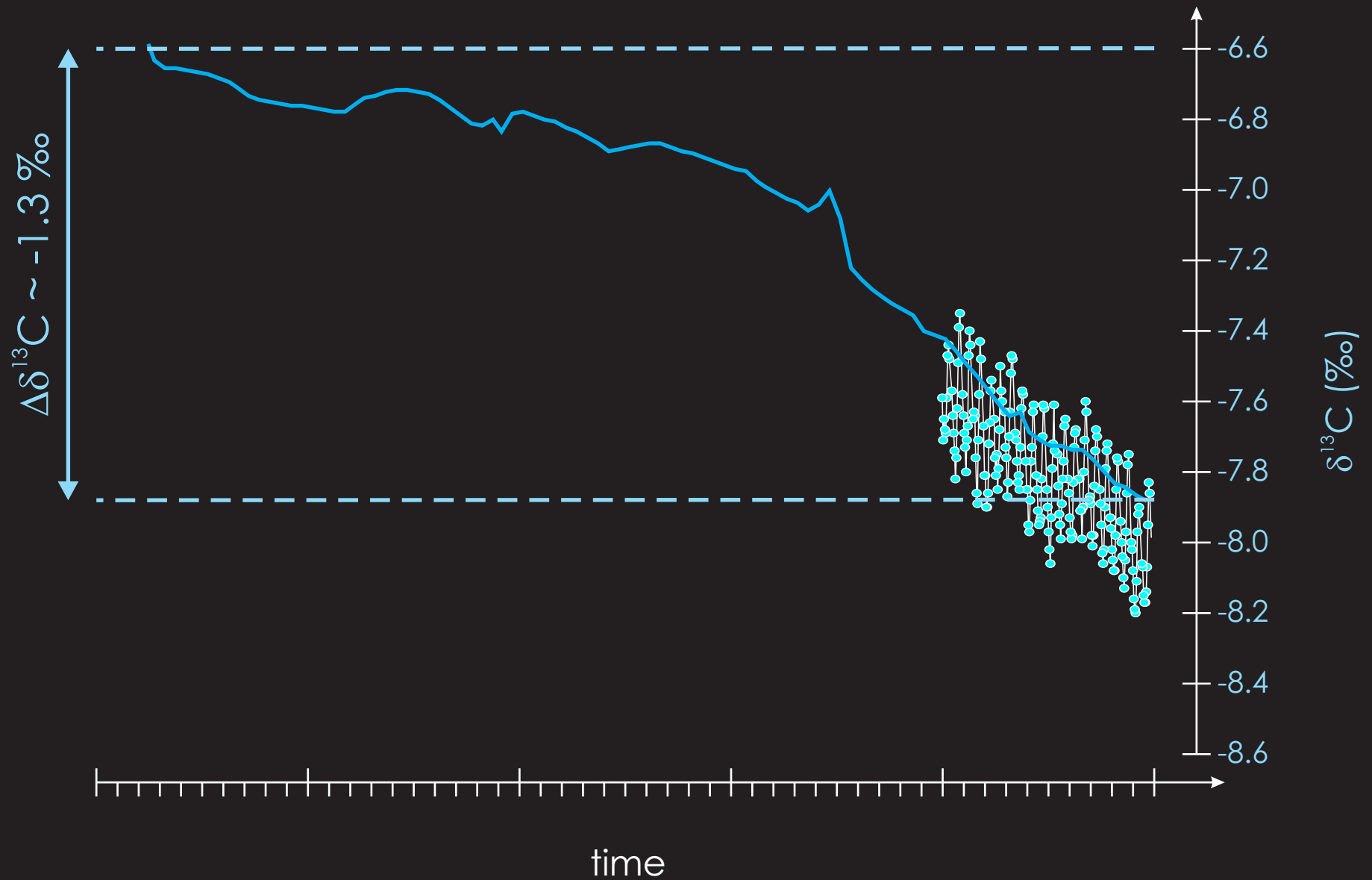
dissolution  
(preservation)

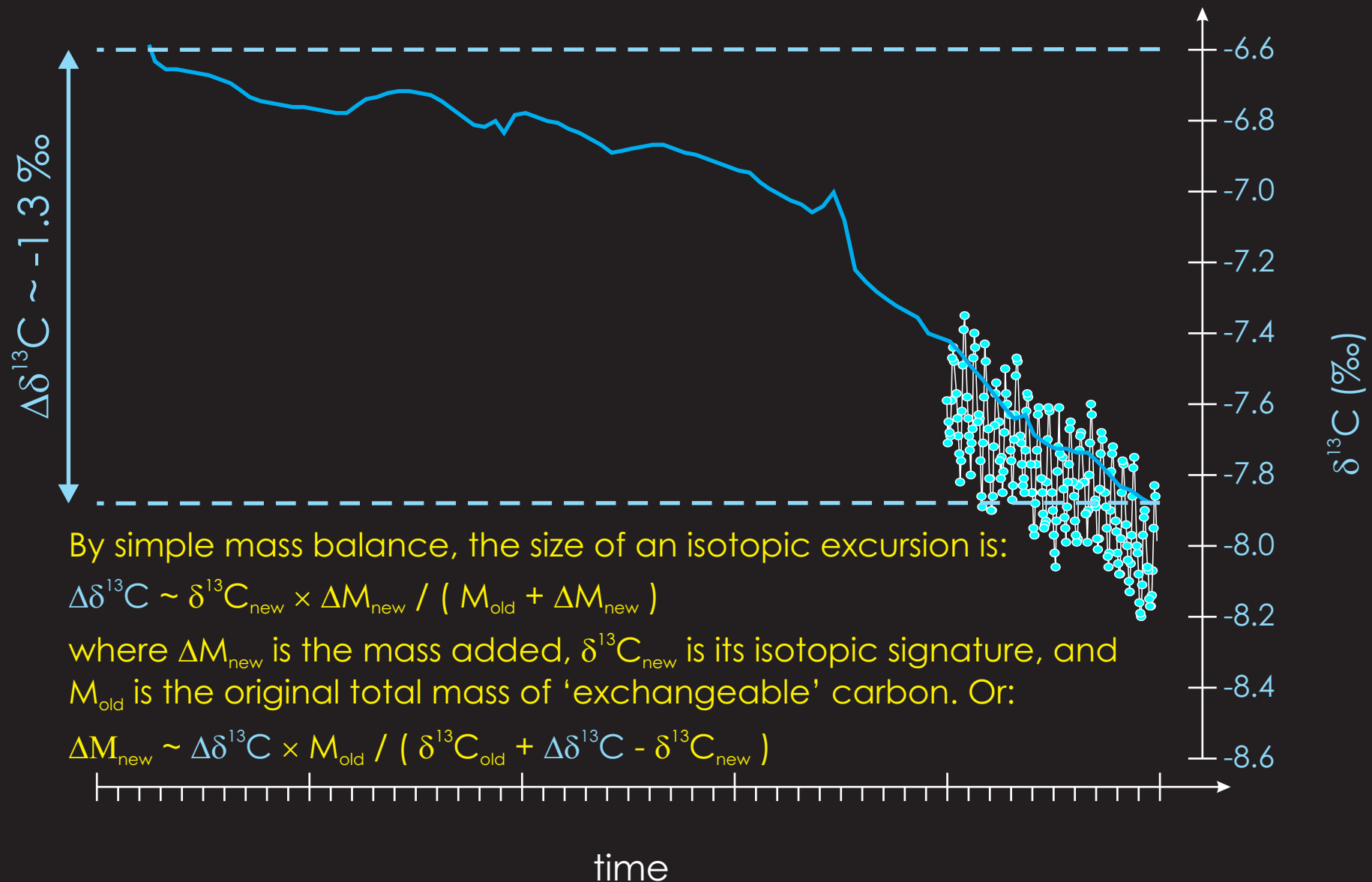


mixing  
(bioturbation)

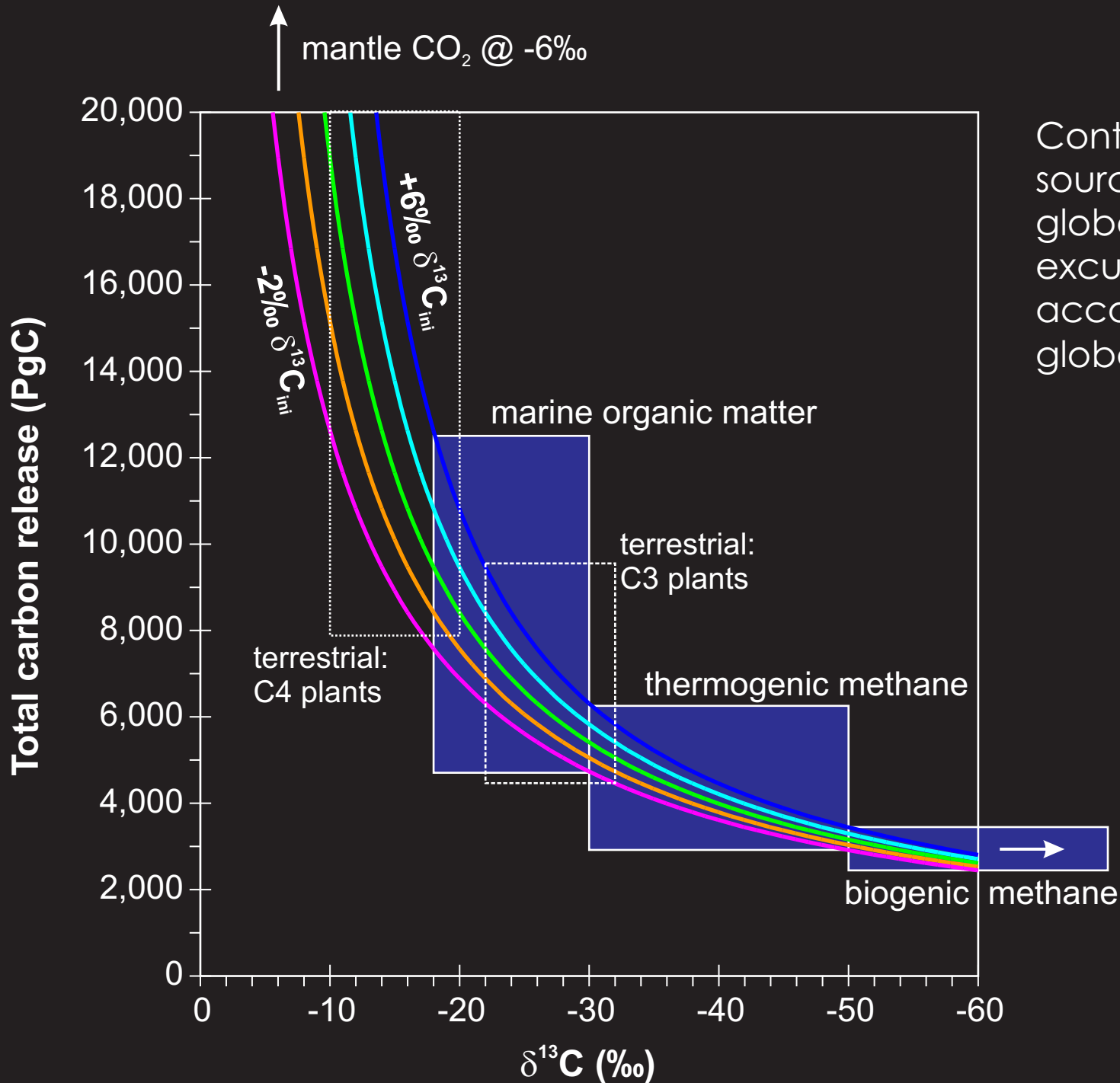




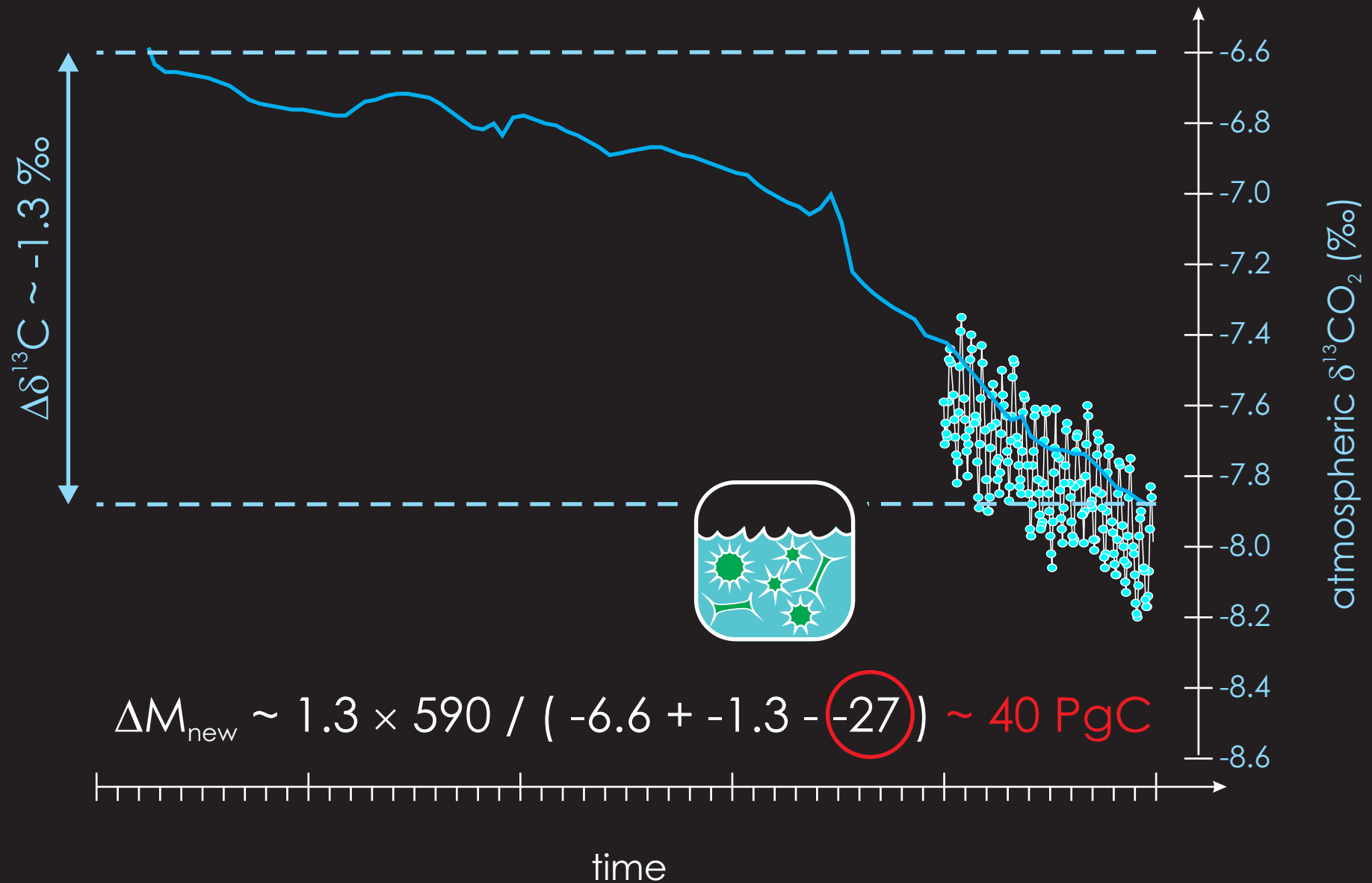




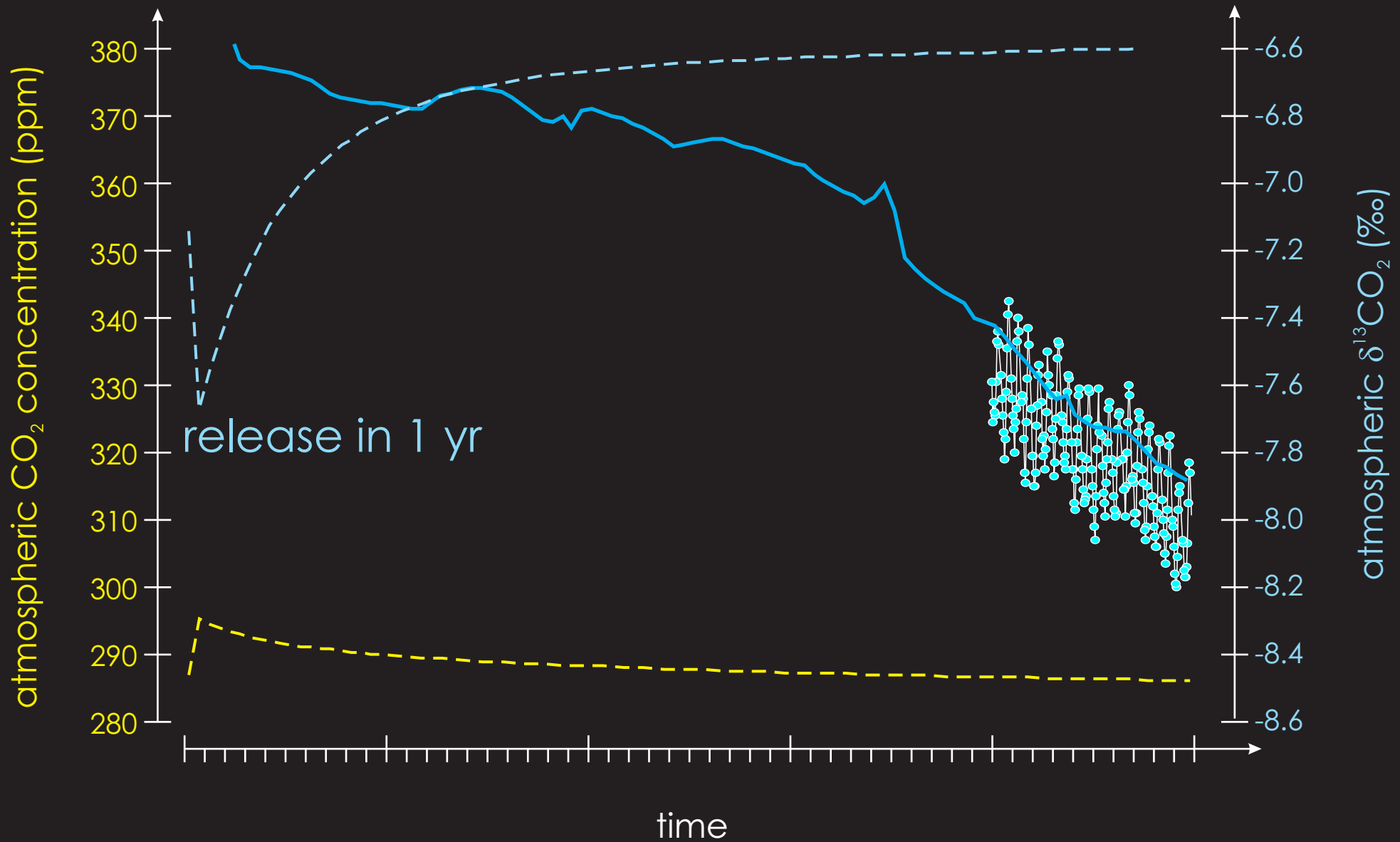


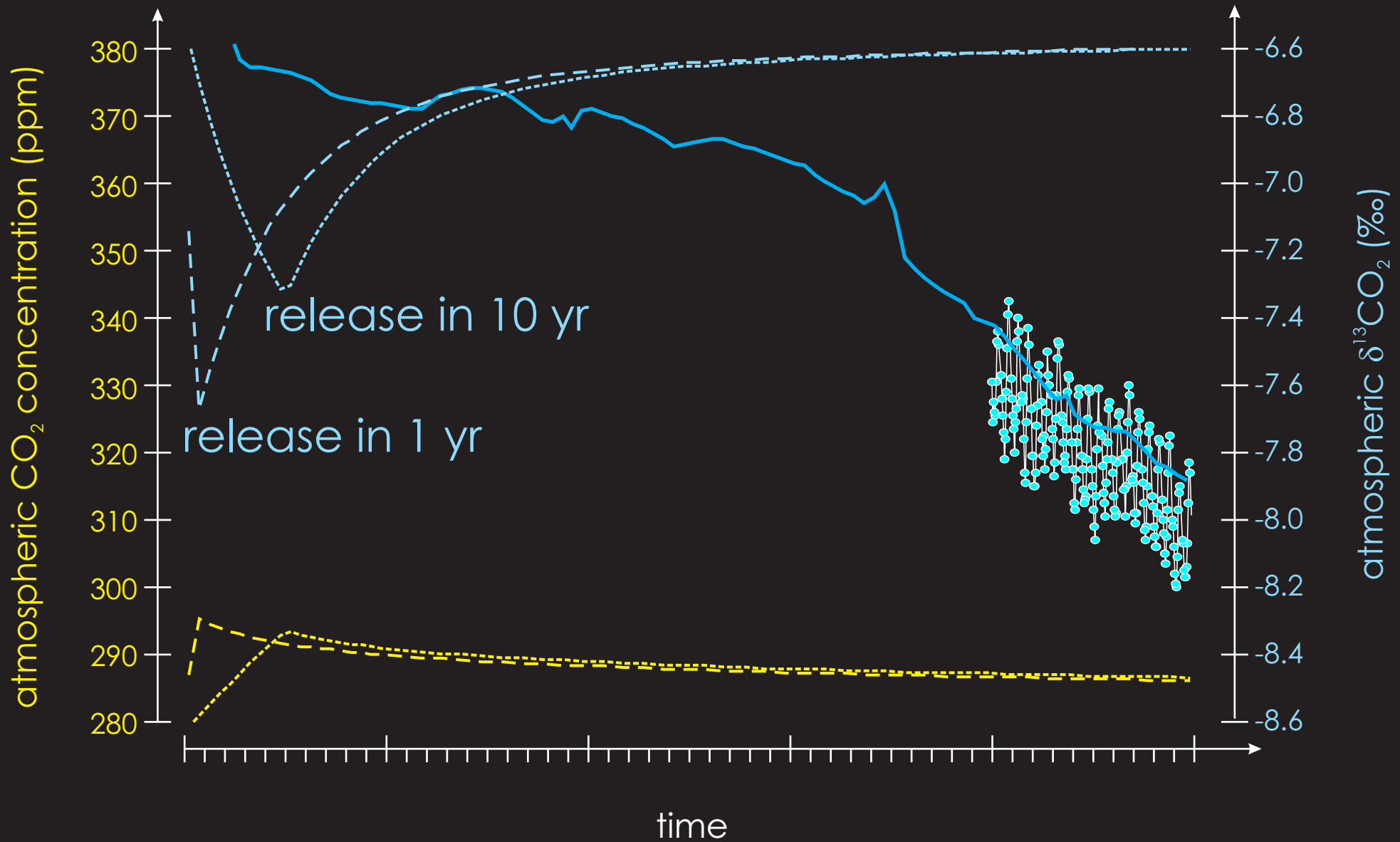


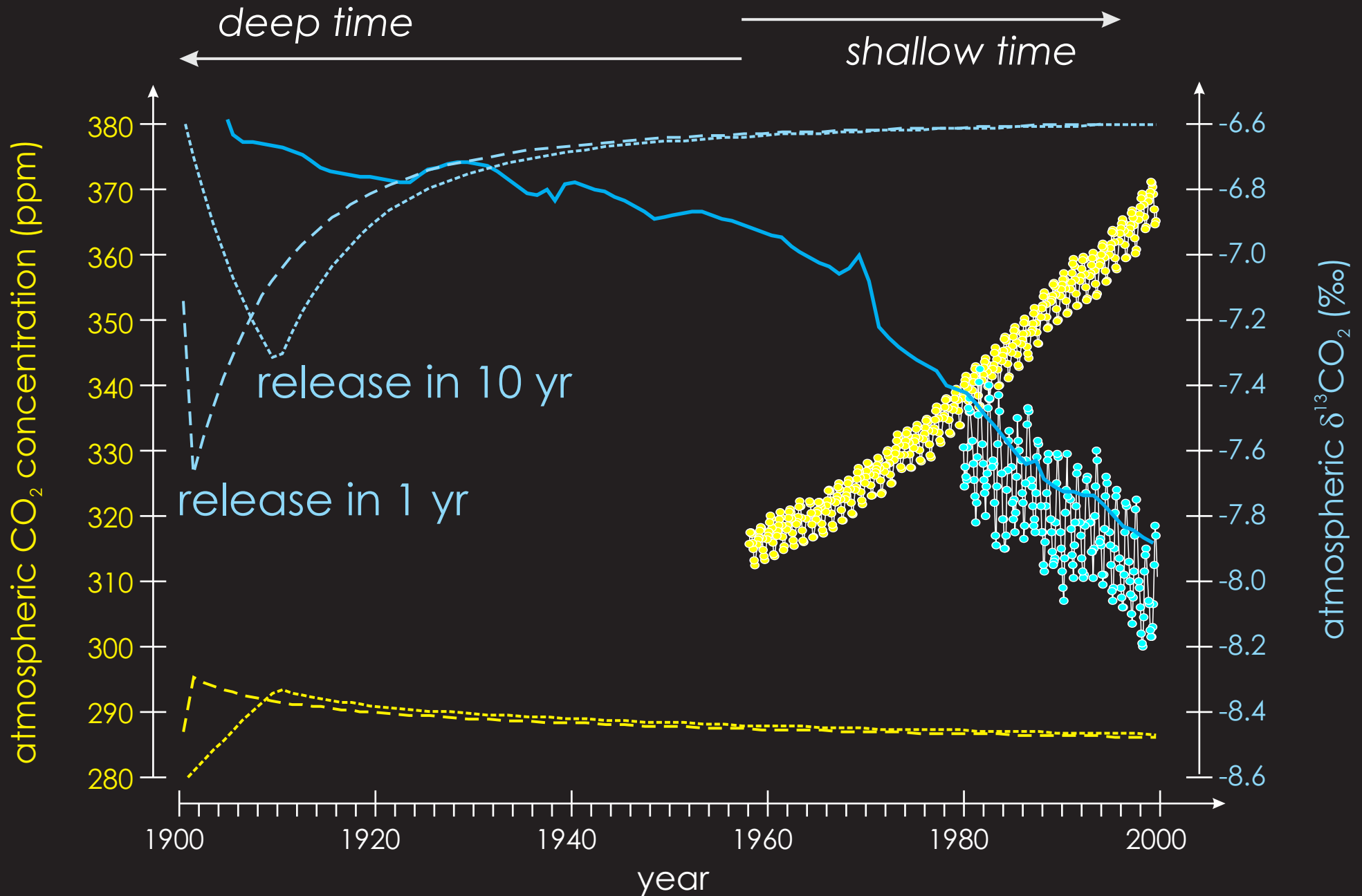
Contours of carbon release vs. source isotopic signature for a global -4‰ carbon isotopic excursion. Contours differ according to the initial mean global  $\delta^{13}\text{C}$ .





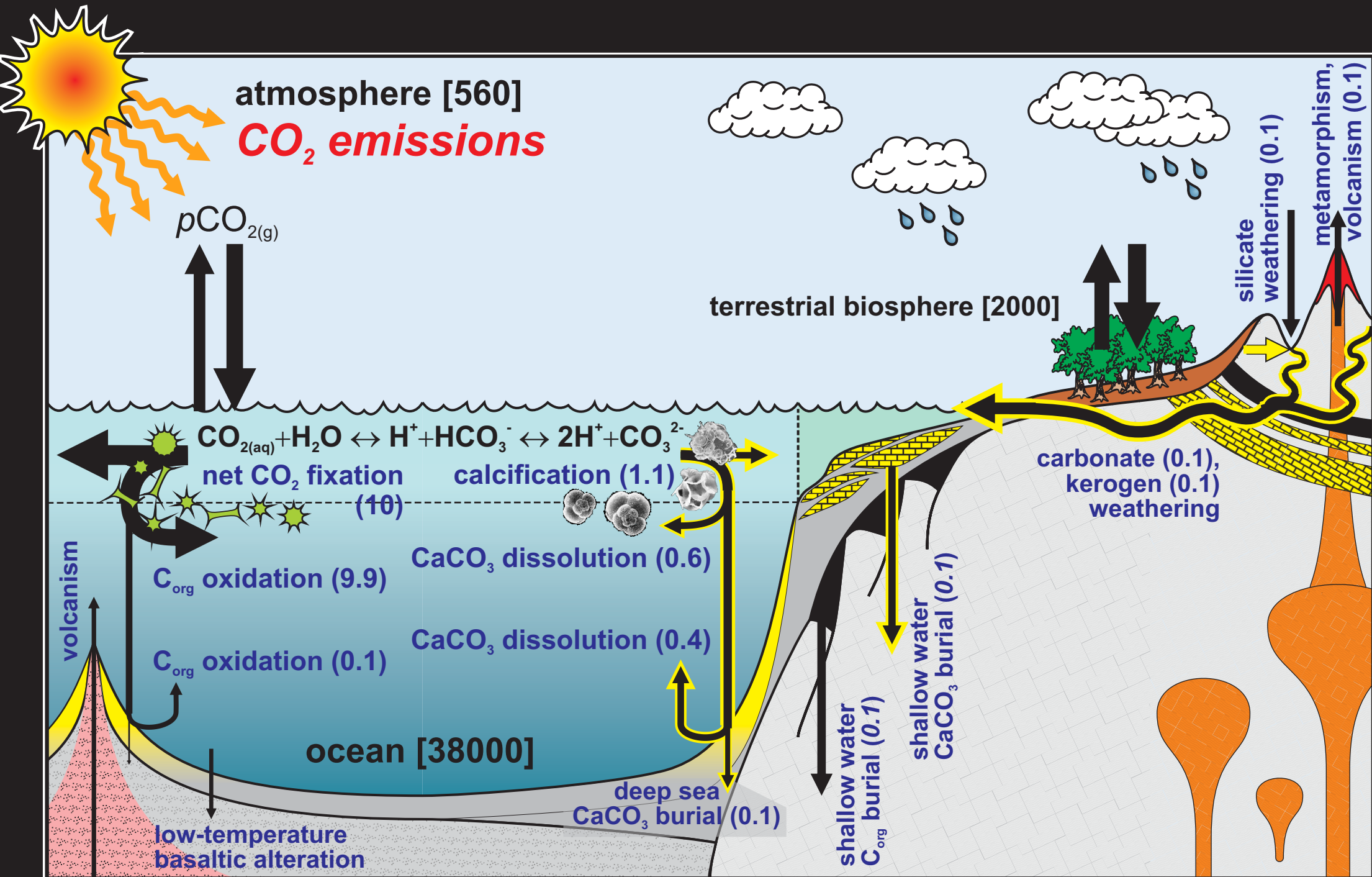




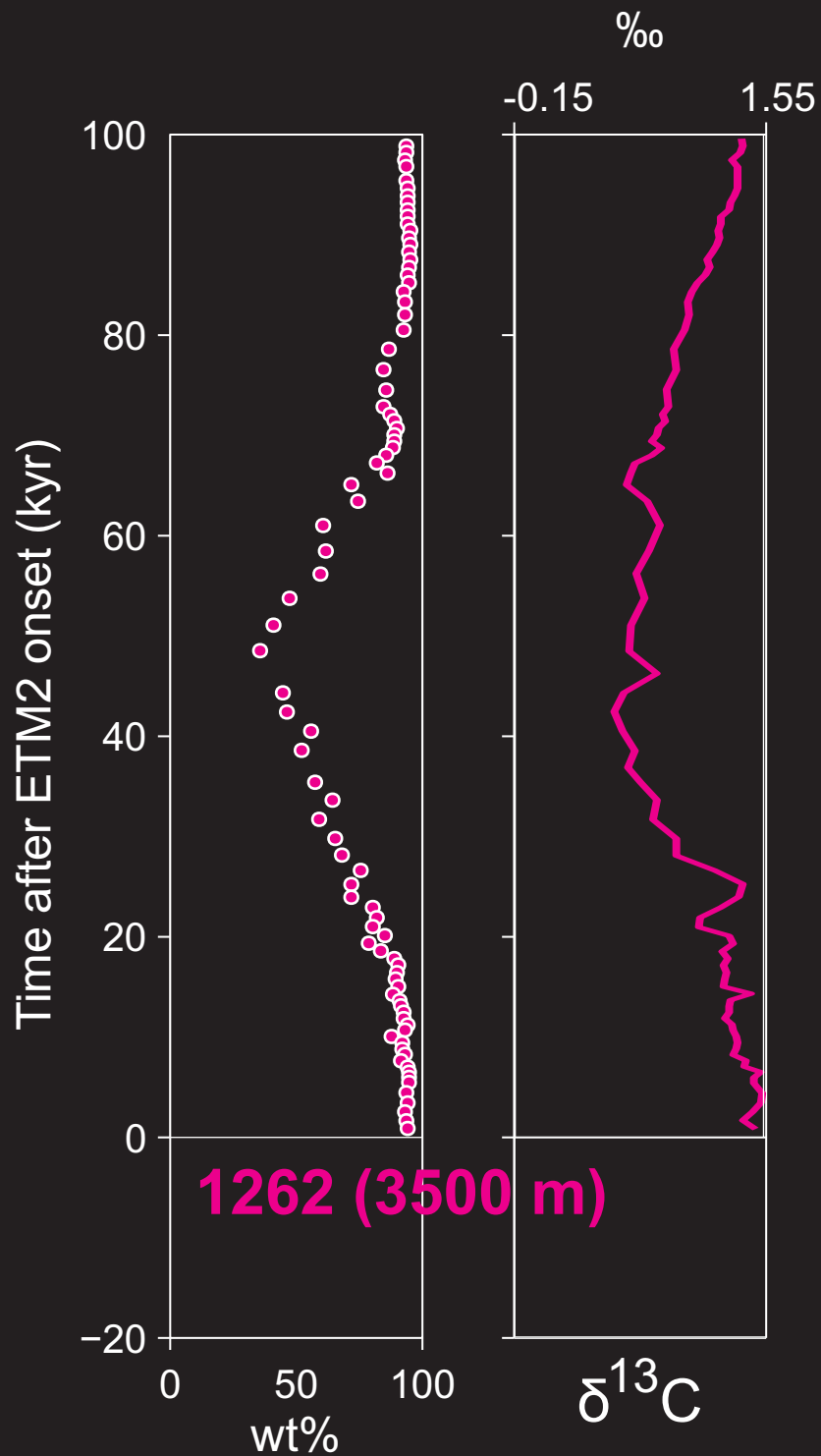




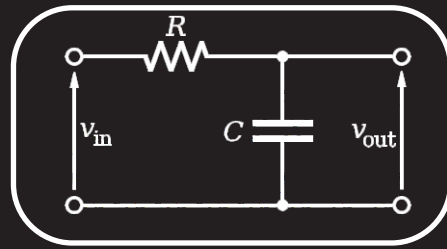
# 'Inverting' isotopic records



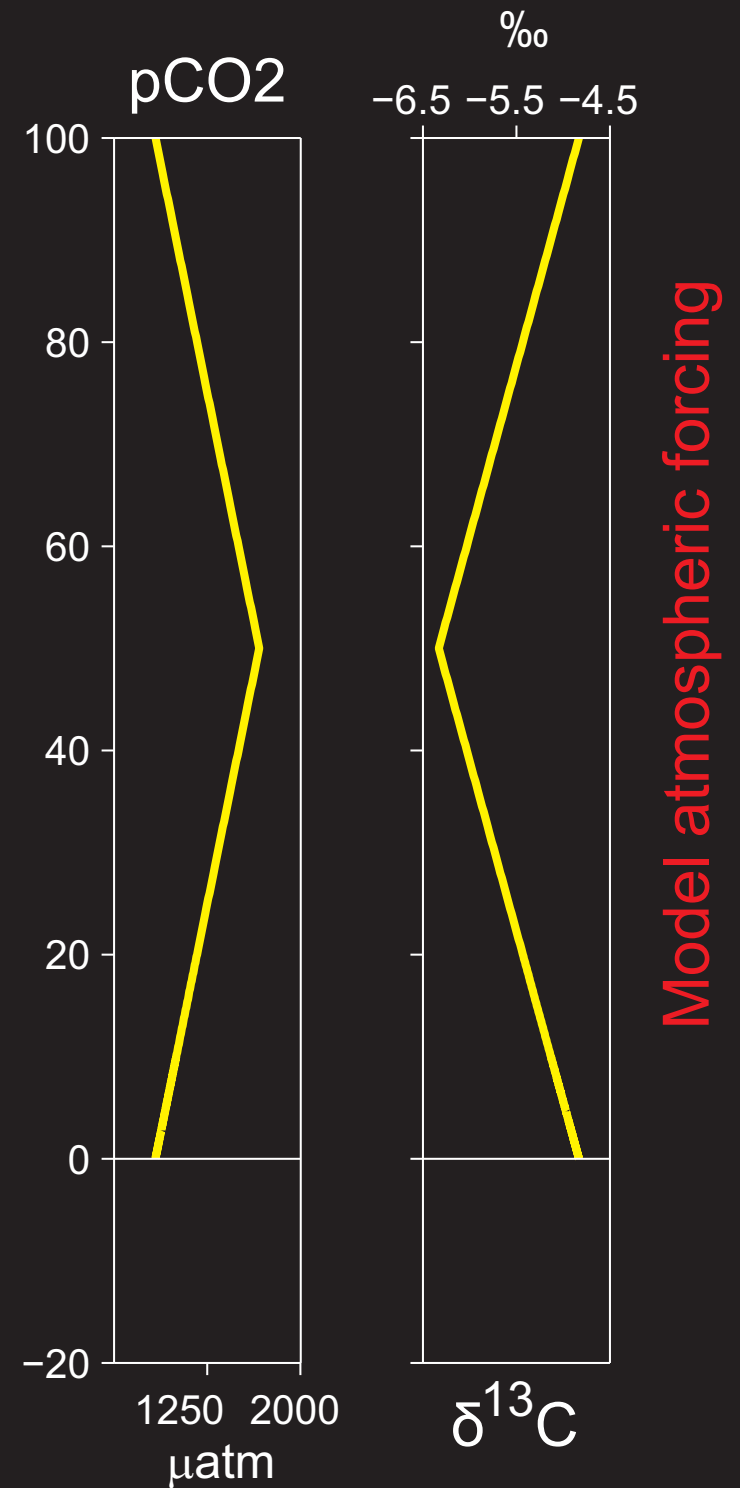


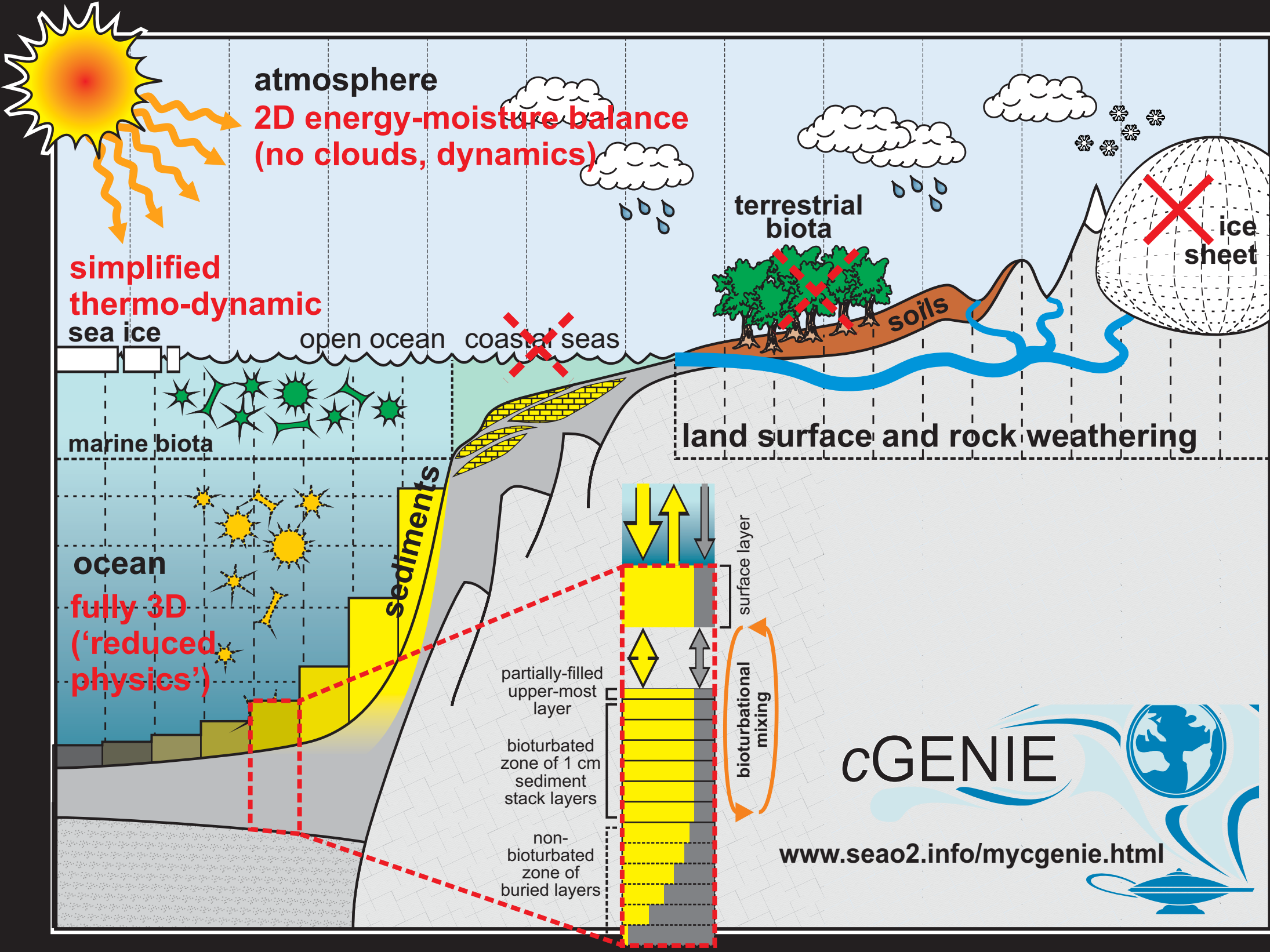




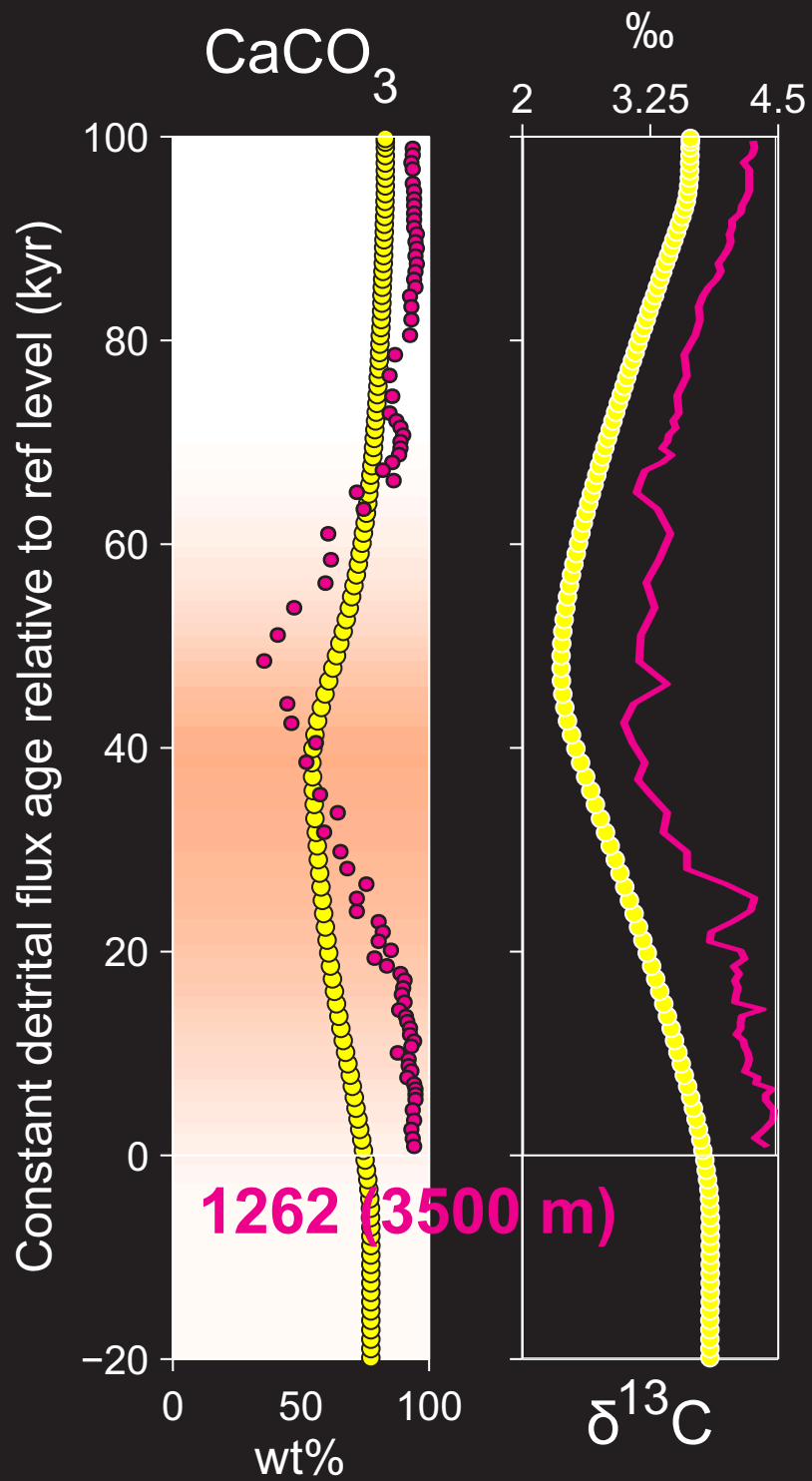


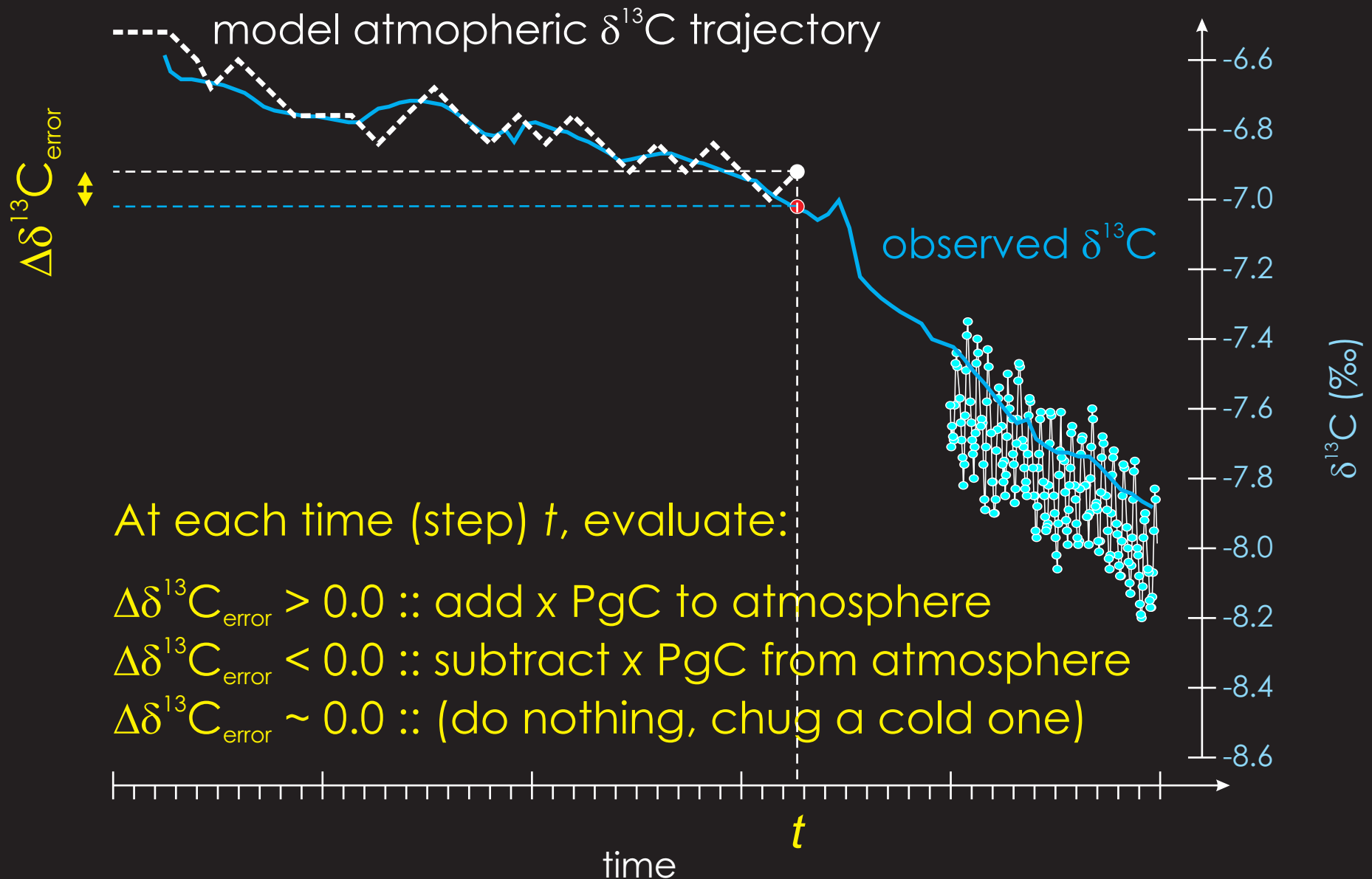
('traditional', forward-modelling approach)





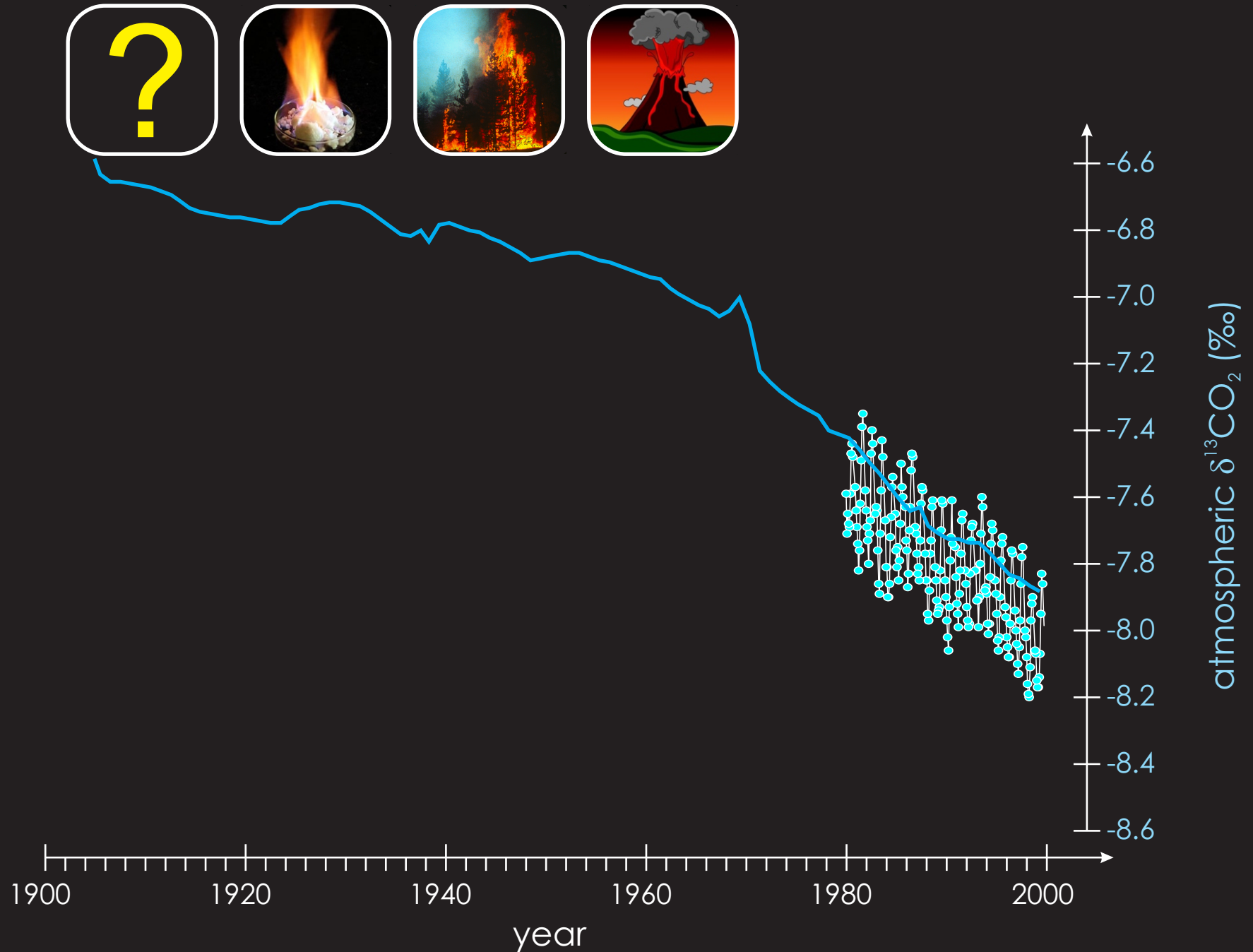
# 'Inverting' isotopic records



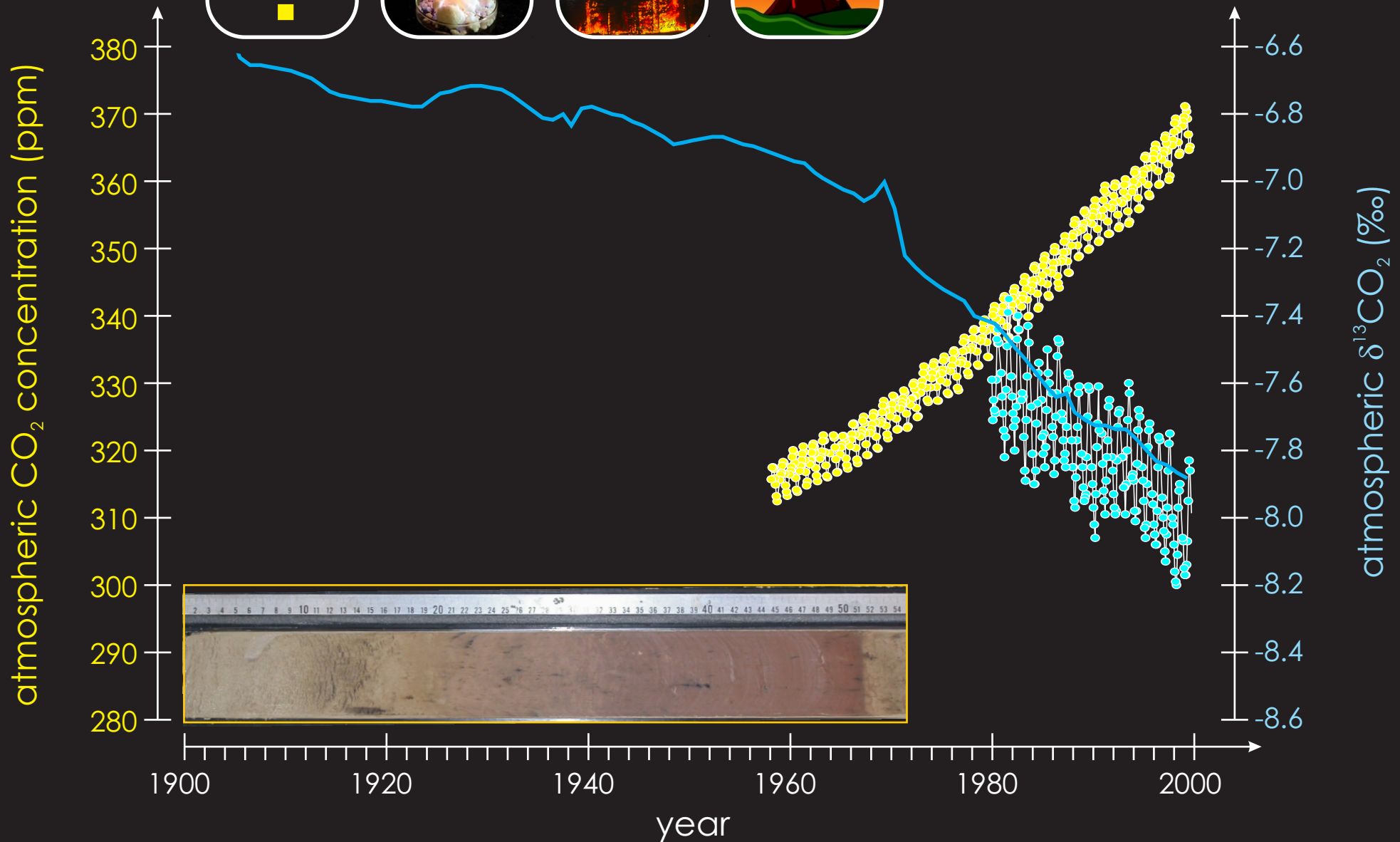


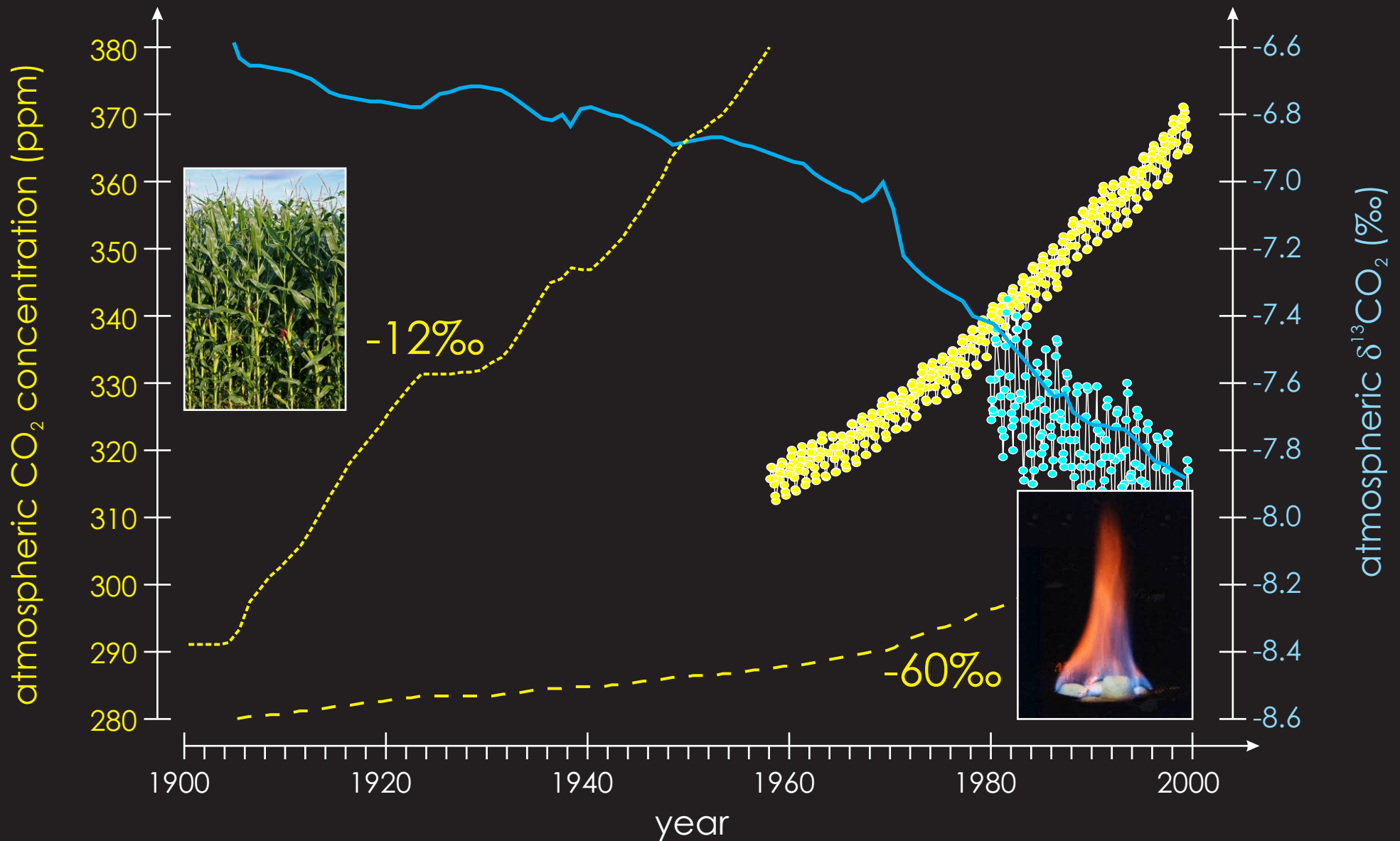


# 'Inverting' isotopic records

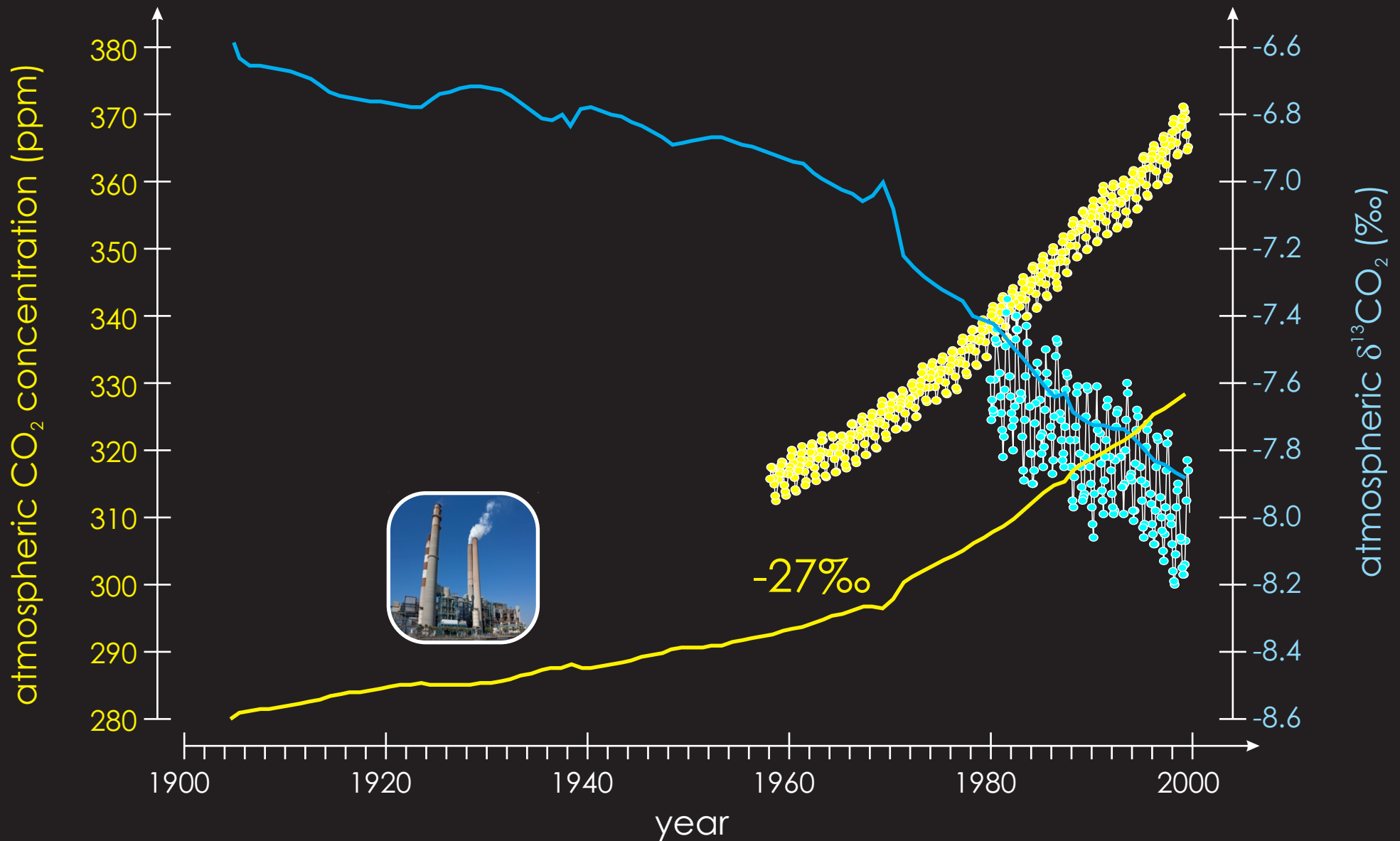


# 'Inverting' isotopic records



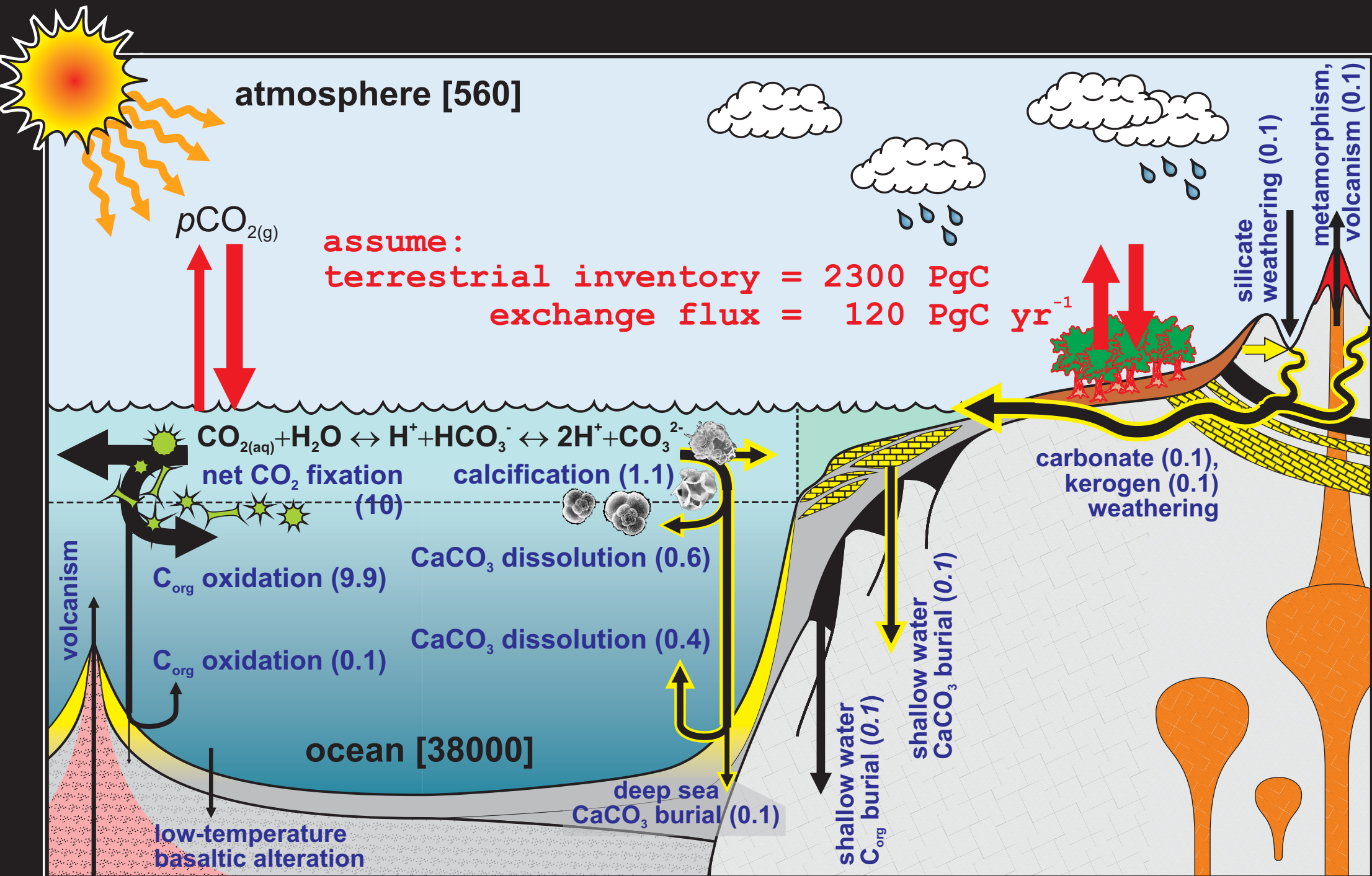


# 'Inverting' isotopic records

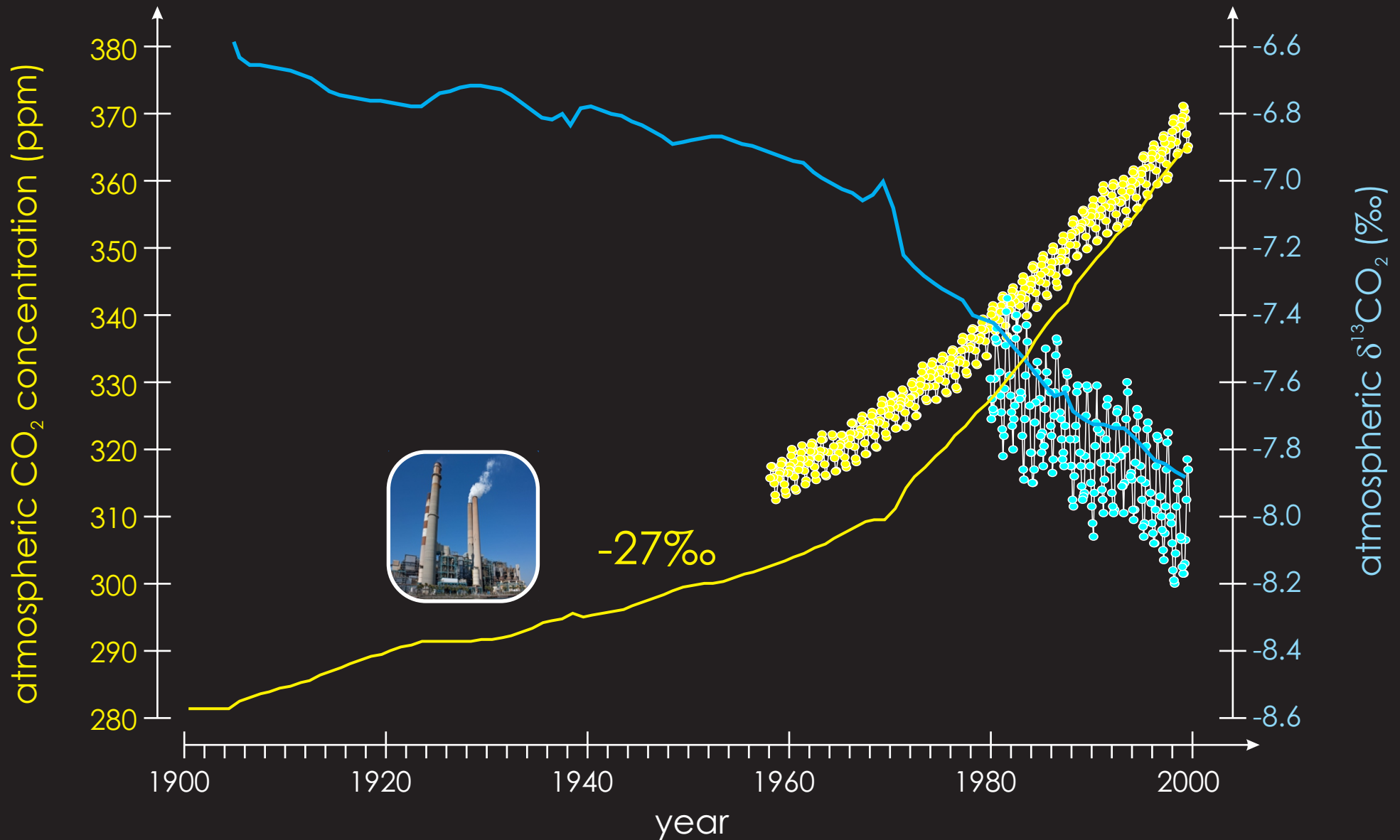




# 'Inverting' isotopic records



# 'Inverting' isotopic records

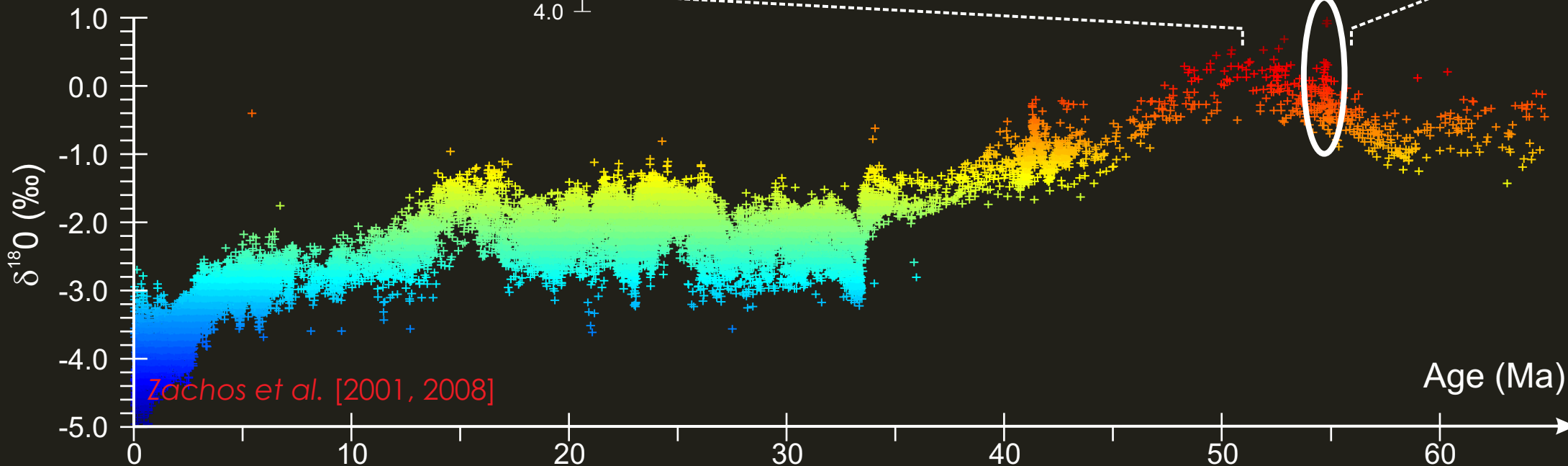
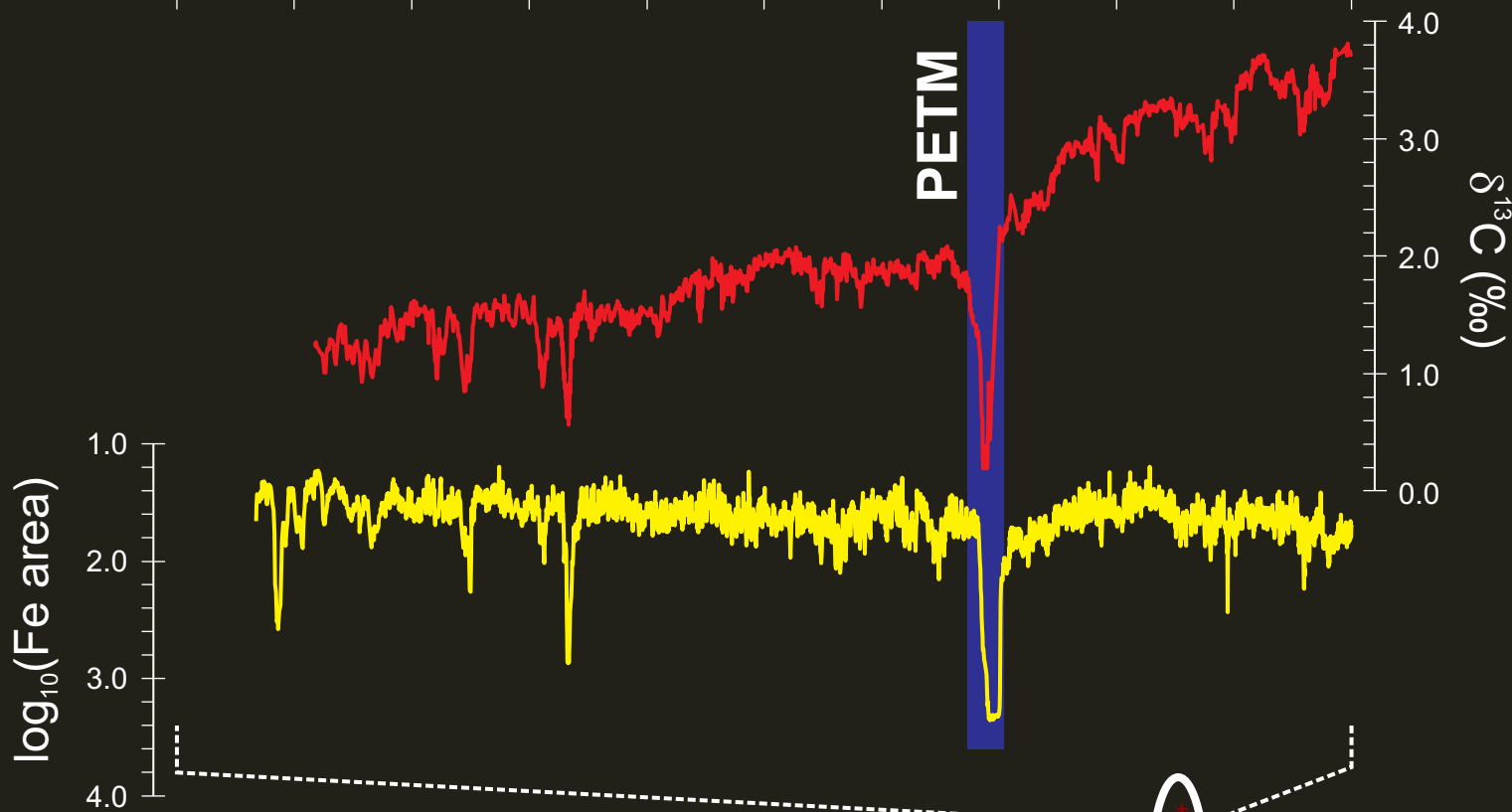


# *Simple 'inversions' of isotopic records*

*Zachos et al. [2010]*  
*Lunt et al. [2011]*

Age relative to the PETM (Ma)

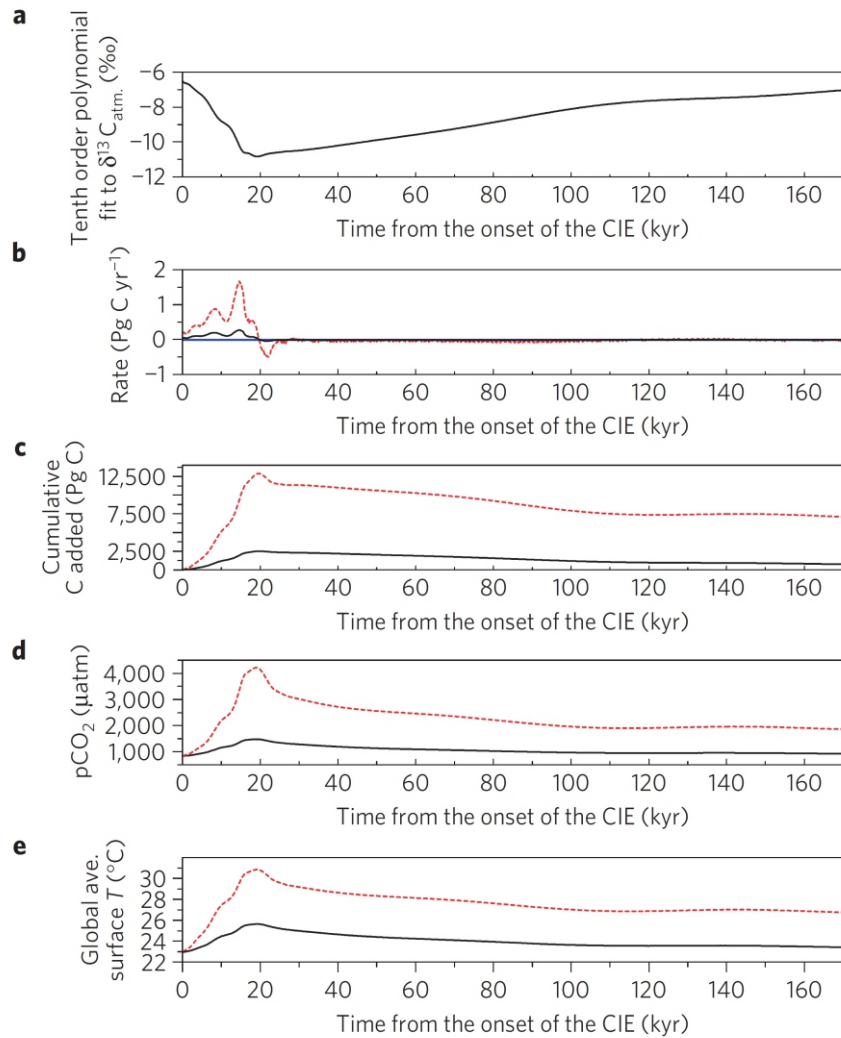
-3.5 -3.0 -2.5 -2.0 -1.5 -1.0 -0.5 0.0 0.5 1.0 1.5



*Zachos et al. [2001, 2008]*

Age (Ma)





**Figure 4 | Model results of the PETM carbon release rate and cumulative amount of carbon added versus time from the onset of the CIE (535 mbs) (age model is from ref. 2).** **a**,  $\delta^{13}\text{C}_{\text{atm}}$  that we used to force GENIE. **b**, Model results of the PETM carbon release rate. **c**, Model results of the cumulative amount of carbon added. **d**, Model results of the PETM atmospheric  $p\text{CO}_2$ . **e**, Model results of the PETM global average temperature ( $^{\circ}\text{C}$ ). The two best-fit simulations are shown in **b-e**: (1)  $\text{CH}_4$  simulation (black solid line); (2)  $\text{C}_{\text{org}}$  simulation (red dotted line). Both simulations are with bioturbation on.

nature  
geoscience

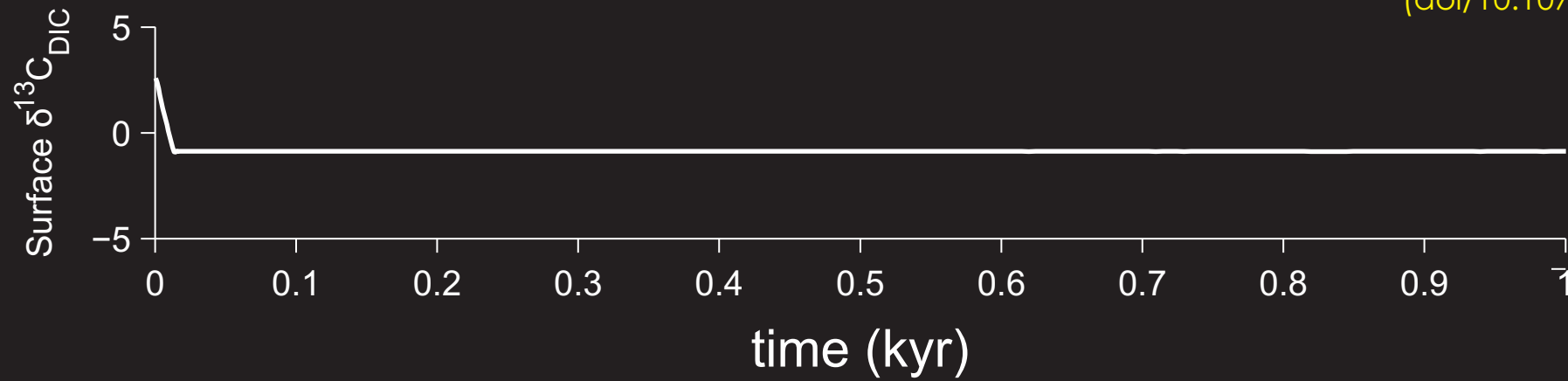
ARTICLES

PUBLISHED ONLINE: 5 JUNE 2011 | DOI: 10.1038/NNGEO1179

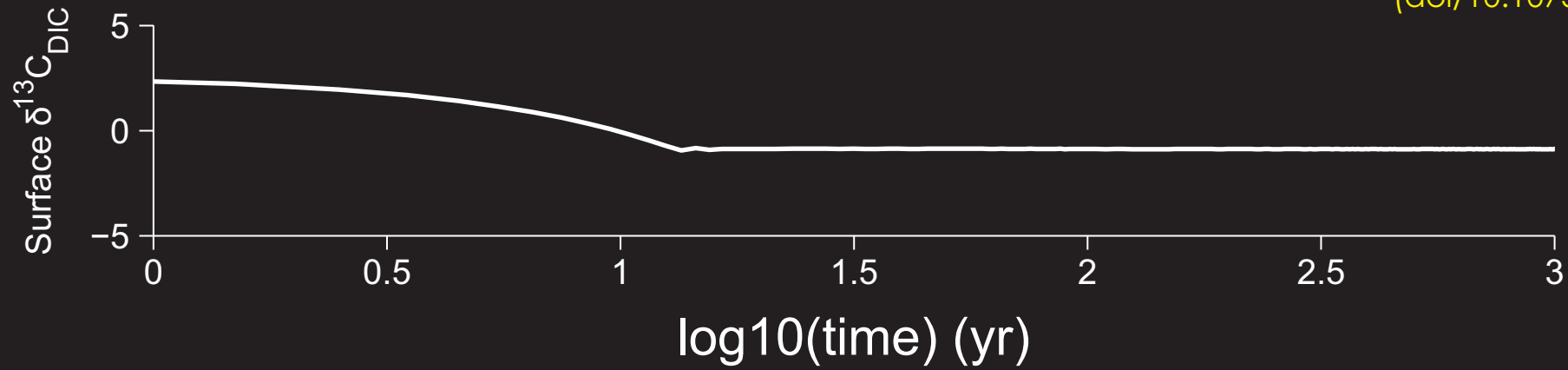
## Slow release of fossil carbon during the Palaeocene-Eocene Thermal Maximum

Ying Cui<sup>1\*</sup>, Lee R. Kump<sup>1</sup>, Andy J. Ridgwell<sup>2</sup>, Adam J. Charles<sup>3</sup>, Christopher K. Junium<sup>1†</sup>, Aaron F. Diefendorf<sup>1†</sup>, Katherine H. Freeman<sup>1</sup>, Nathan M. Urban<sup>1†</sup> and Ian C. Harding<sup>3</sup>

Wright and Schaller [2013]  
(doi/10.1073/pnas.1309188110)

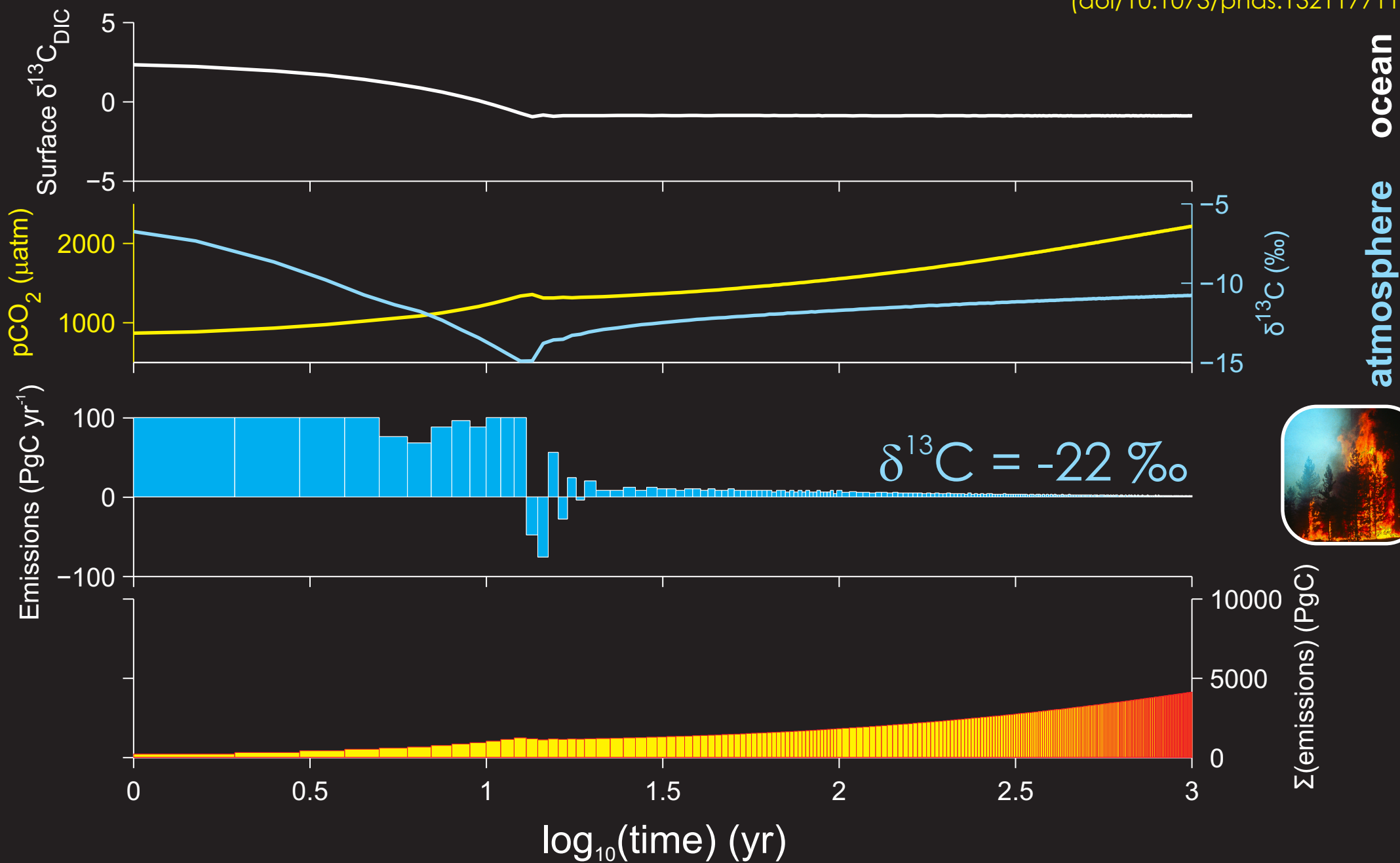


Wright and Schaller [2013]  
(doi/10.1073/pnas.1309188110)



# Simple 'inversions' of isotopic records

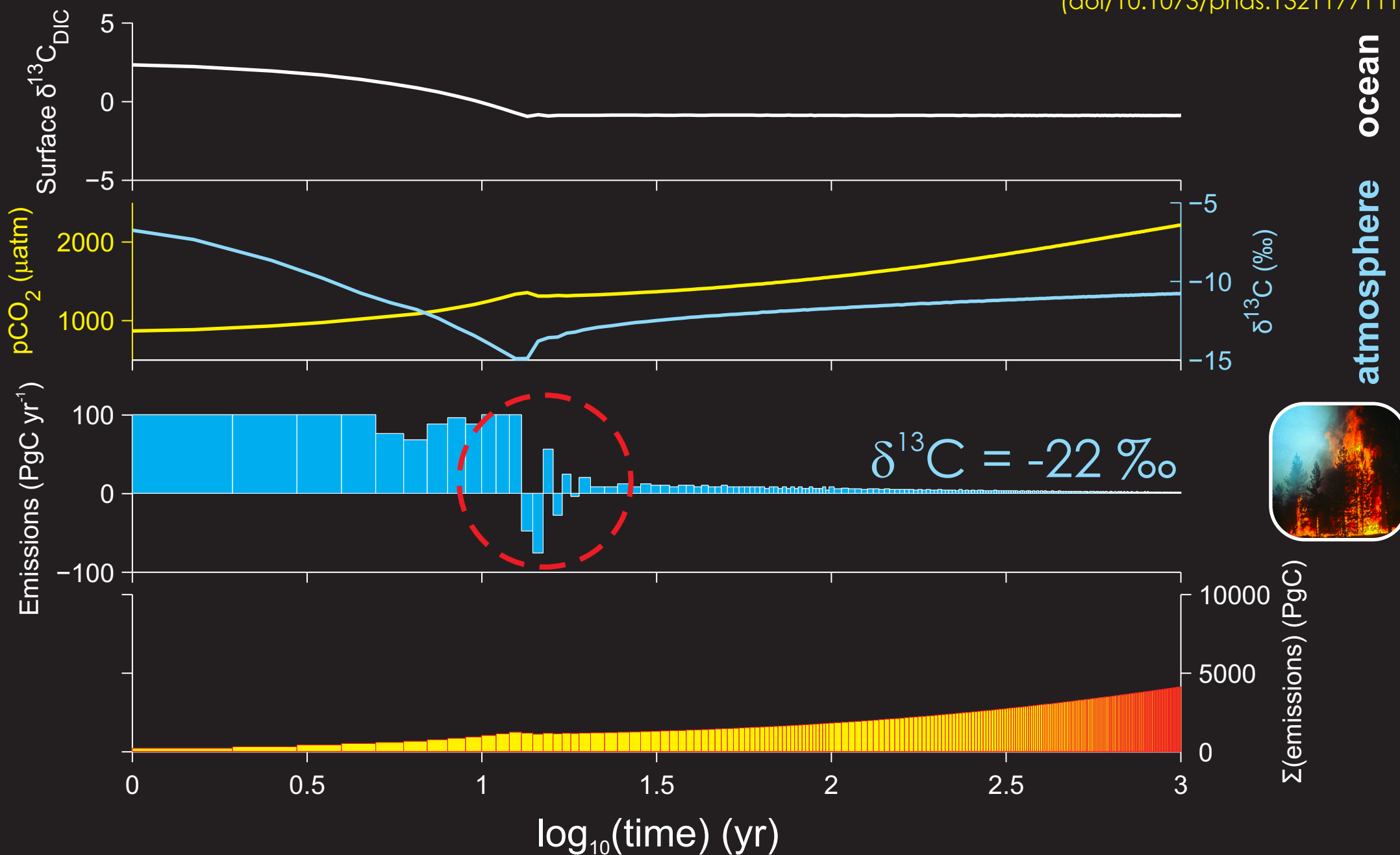
Zeebe et al. [2014]  
(doi/10.1073/pnas.1321177111)



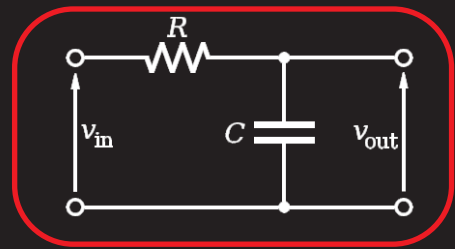


# Simple 'inversions' of isotopic records

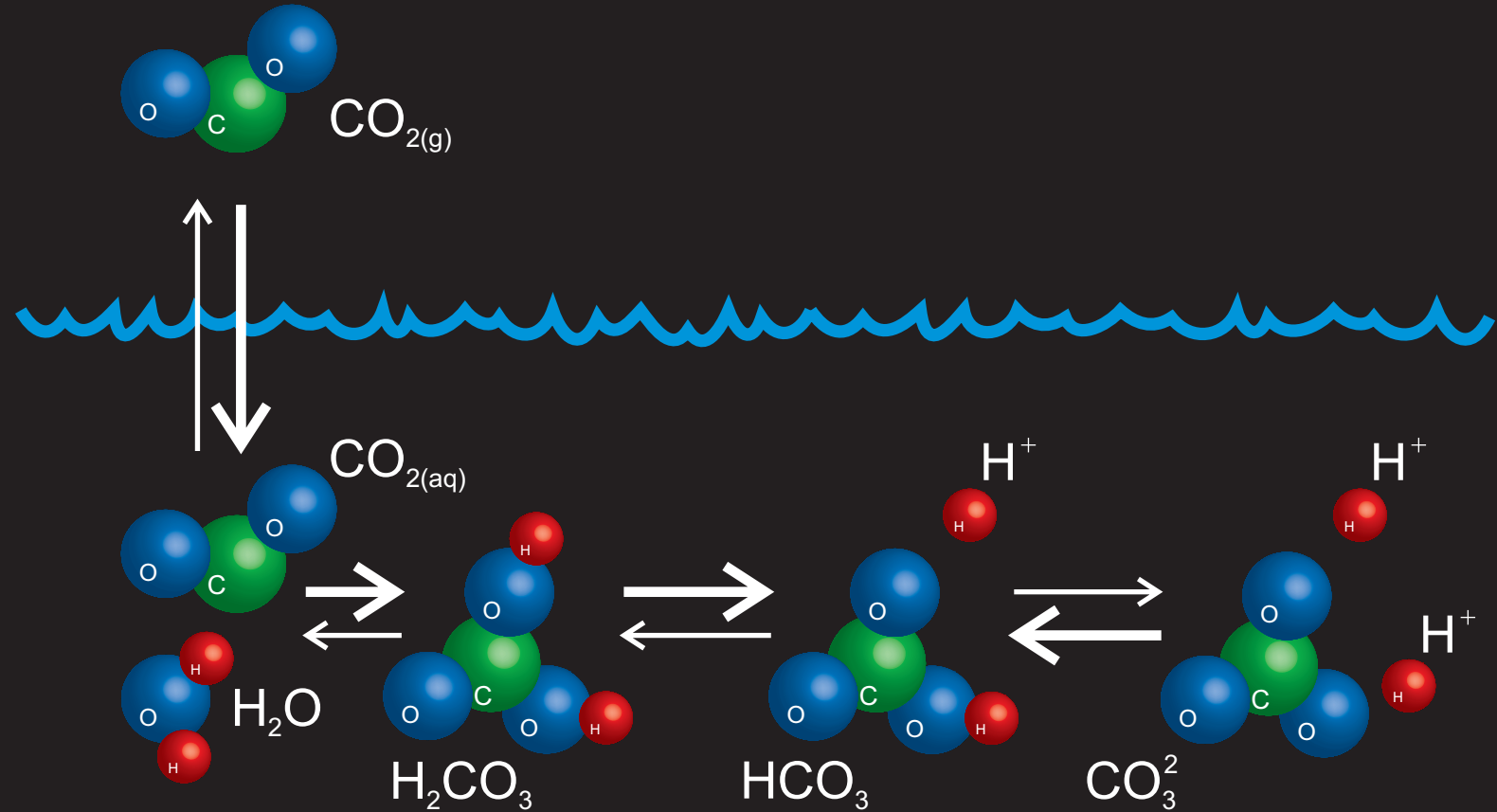
Zeebe et al. [2014]  
(doi/10.1073/pnas.1321177111)



# Simple 'inversions' of isotopic records

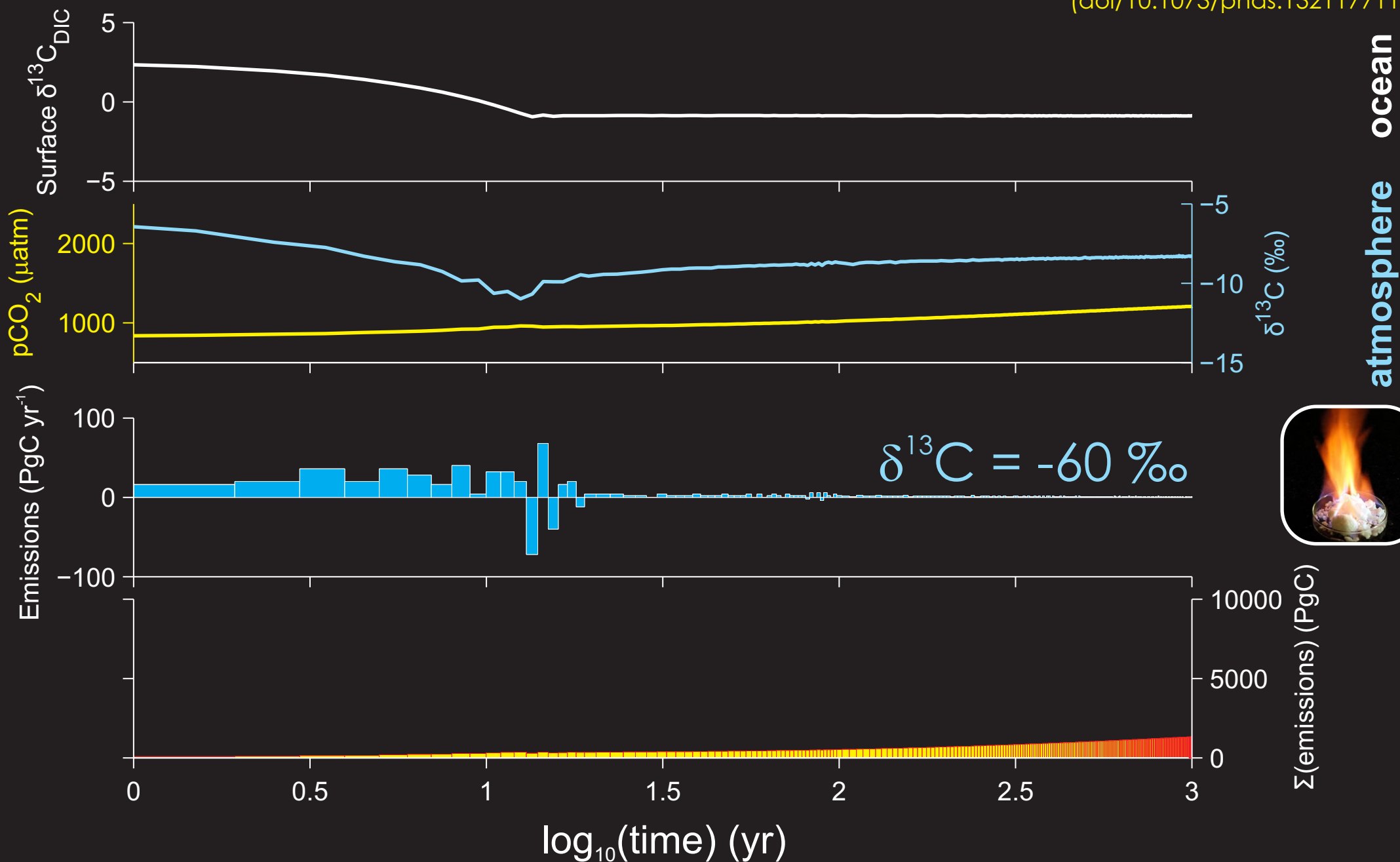


surface ocean  $\delta^{13}\text{C}$  record

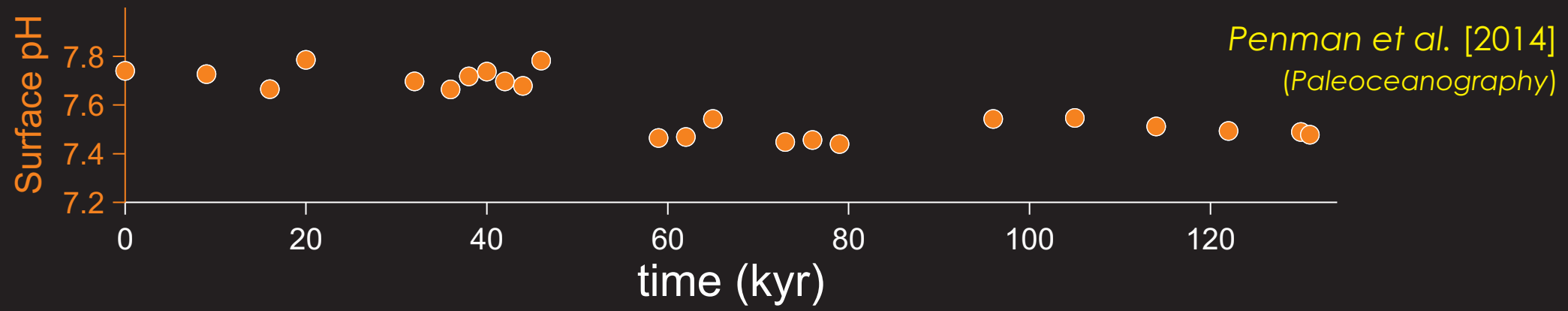


# Simple of 'inversions' of isotopic records

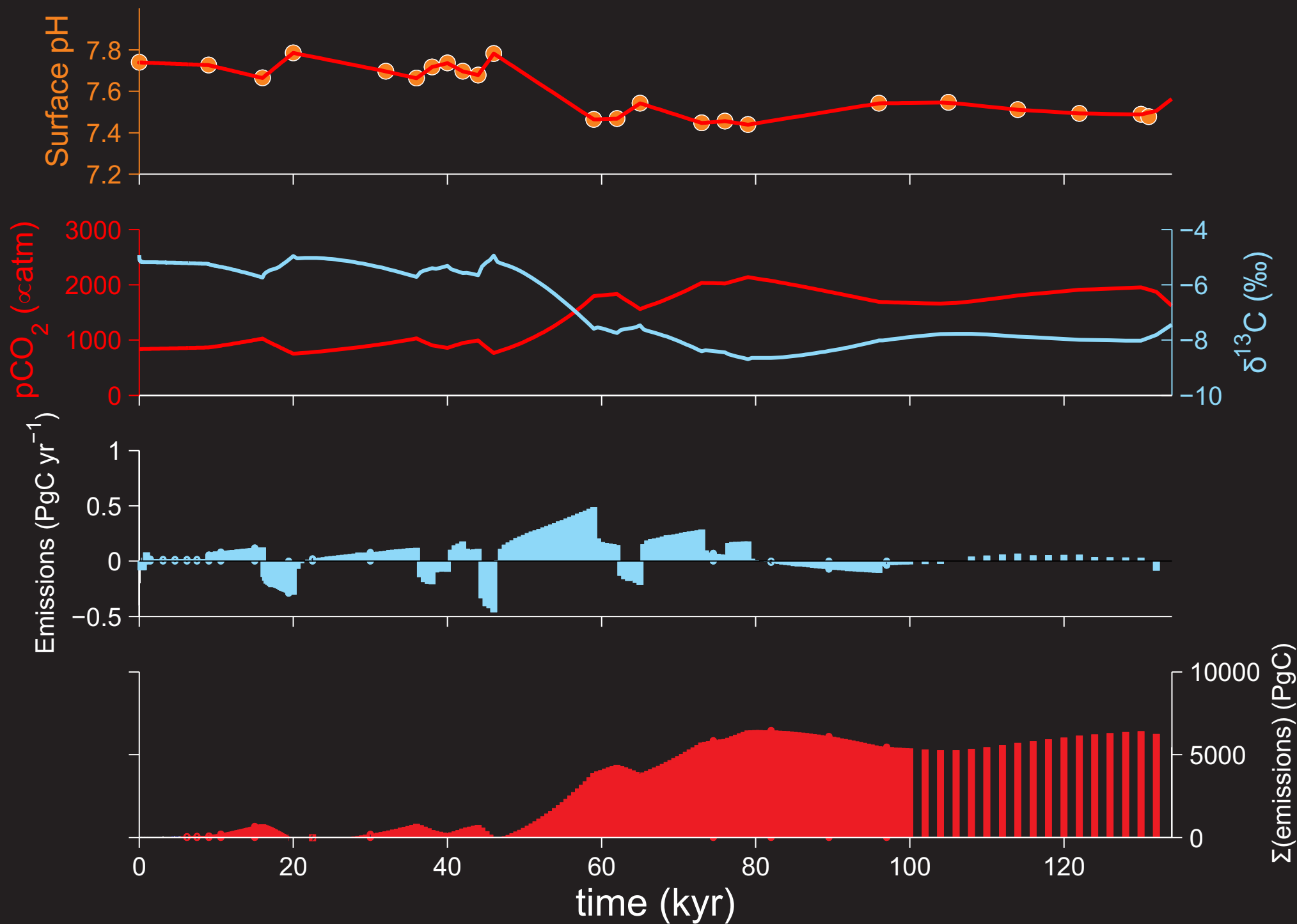
Zeebe et al. [2014]  
(doi/10.1073/pnas.1321177111)



# Simple 'inversions' of isotopic records



# Simple 'inversions' of isotopic records

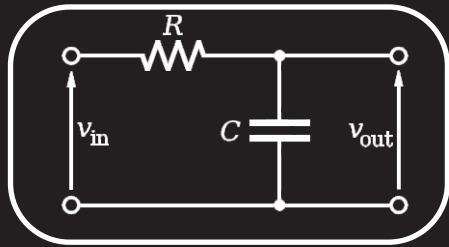




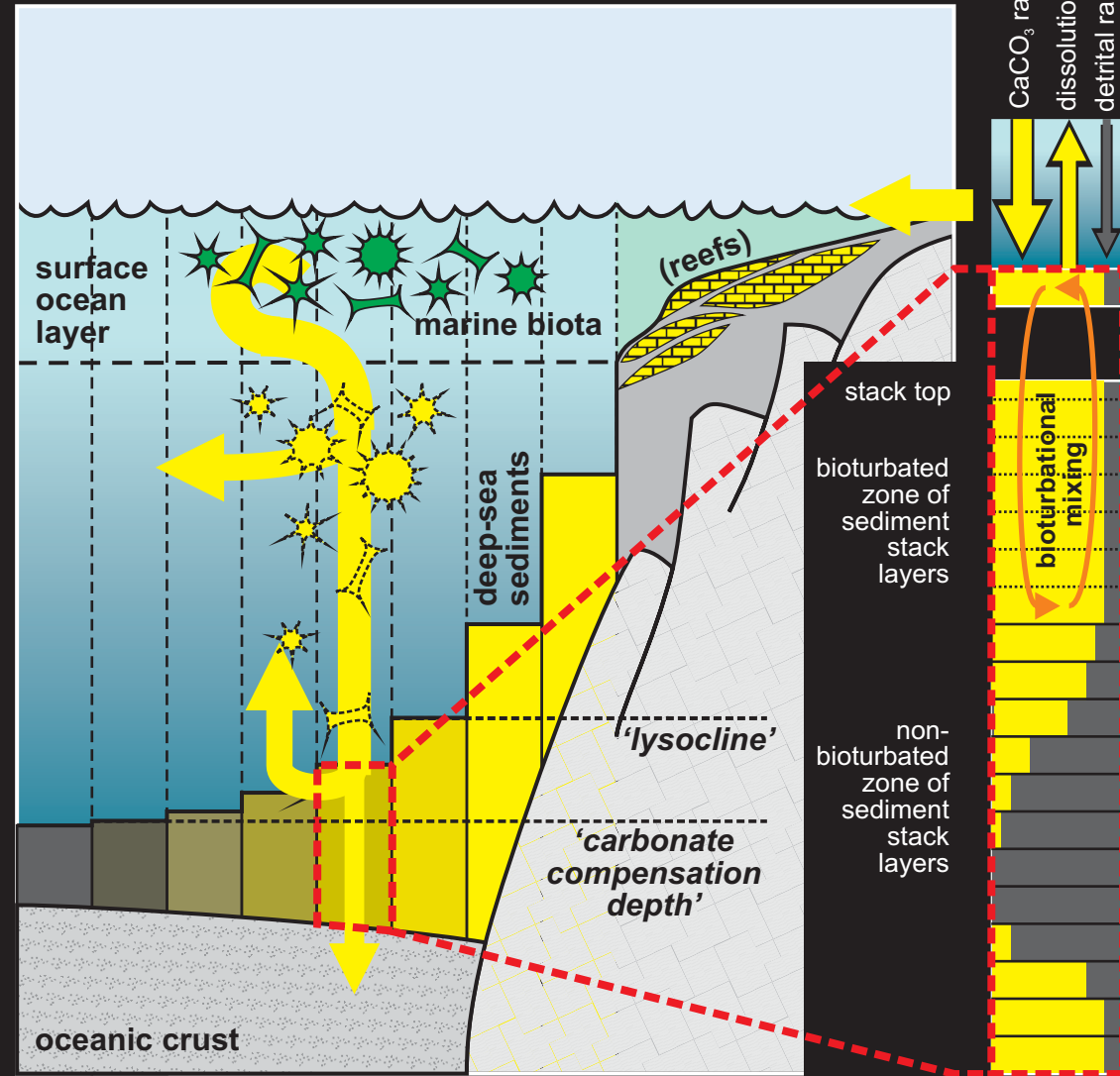
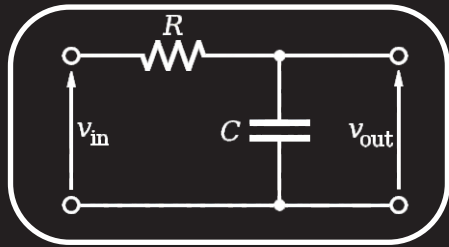
# 'Inverting' isotopic records iteratively



dissolution  
(preservation)



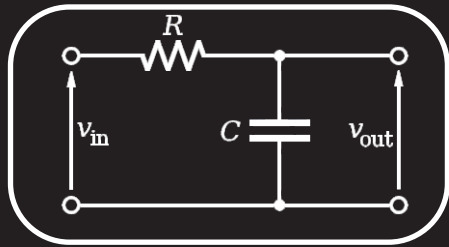
mixing  
(bioturbation)



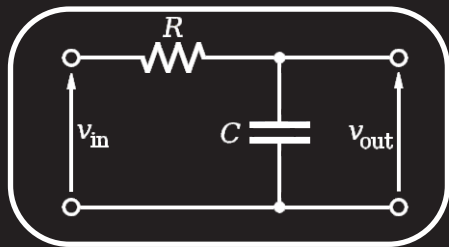
# 'Inverting' isotopic records iteratively



dissolution  
(preservation)



mixing  
(bioturbation)

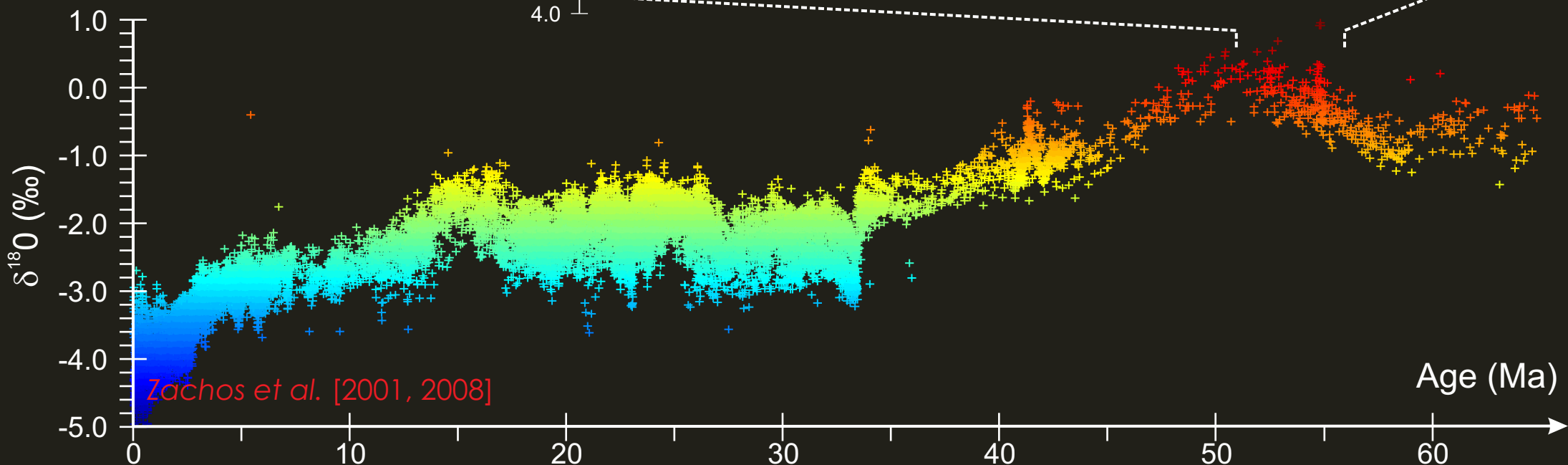
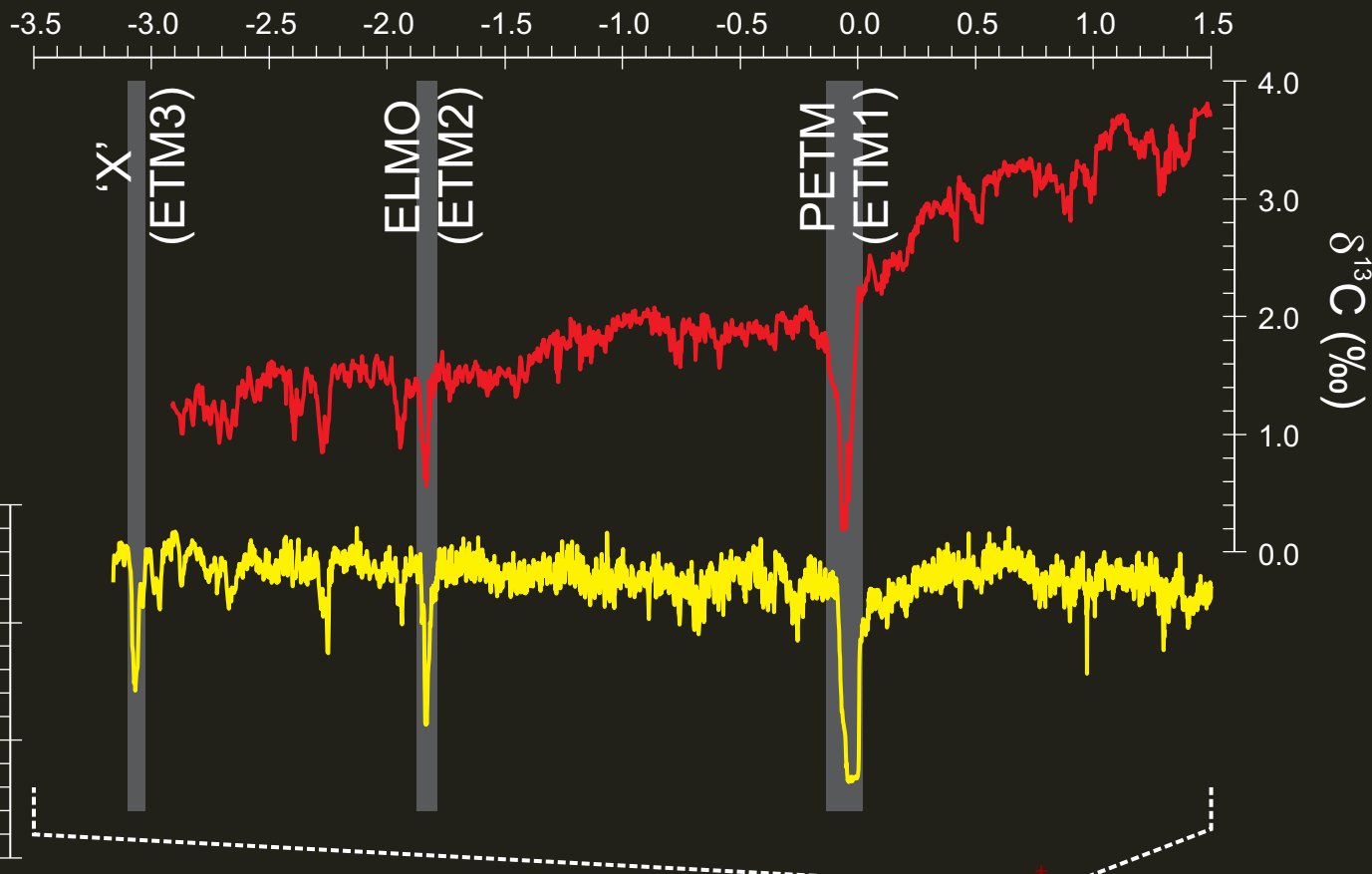


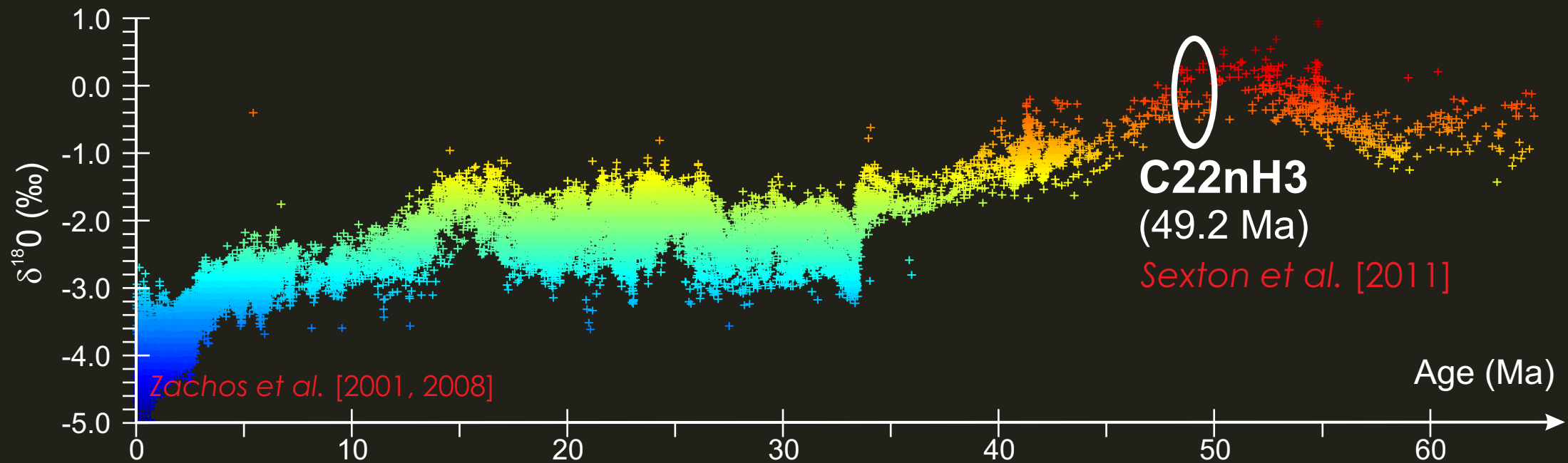
?

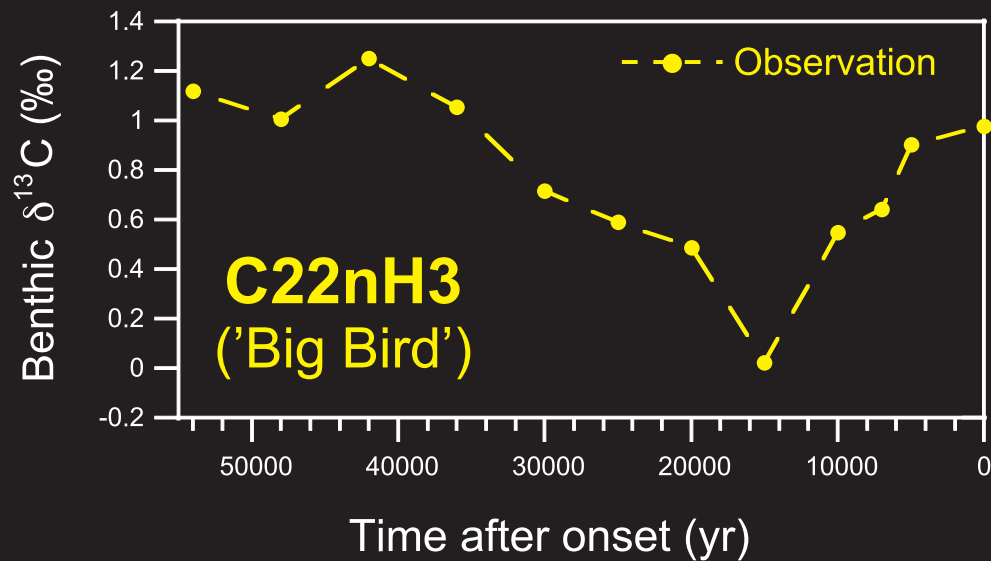


Zachos et al. [2010]  
Lunt et al. [2011]

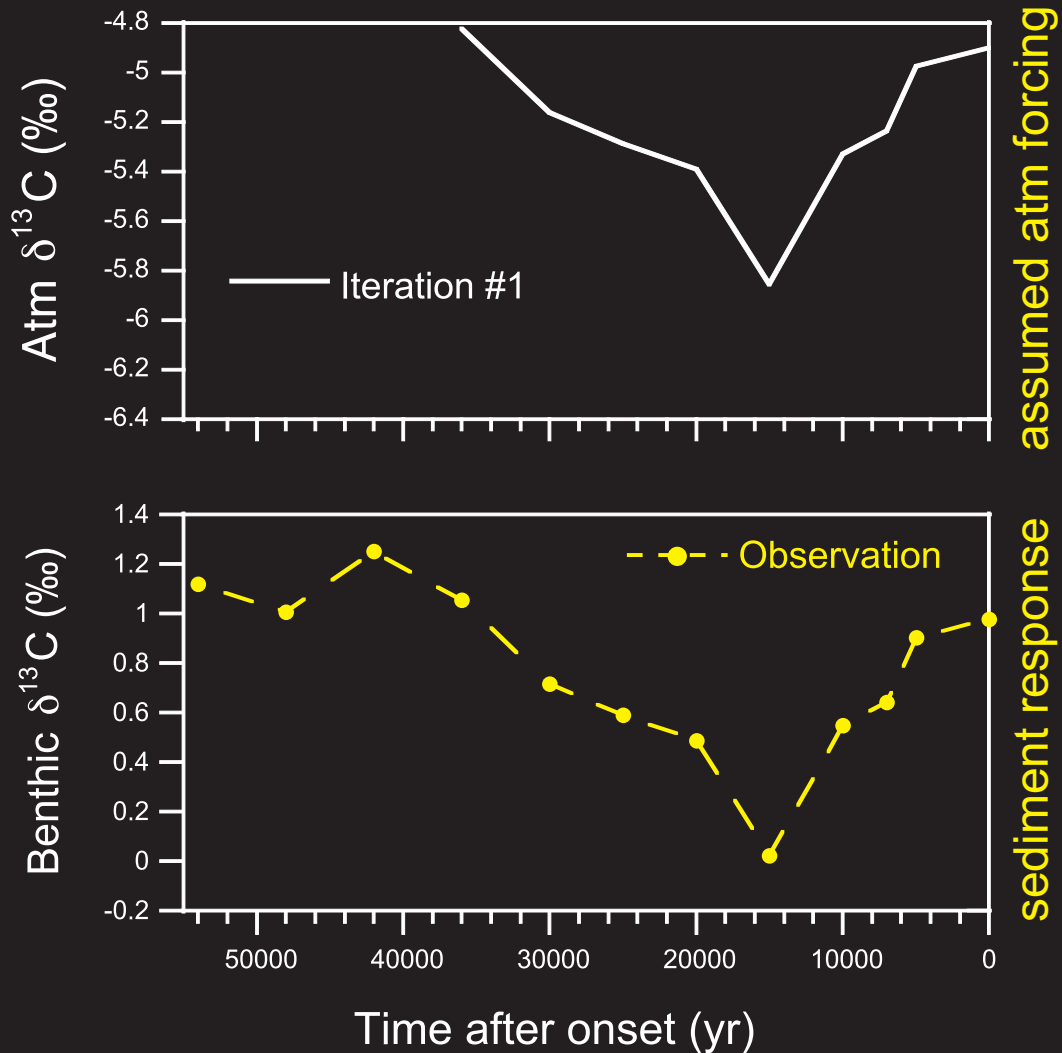
Age relative to the PETM (Ma)



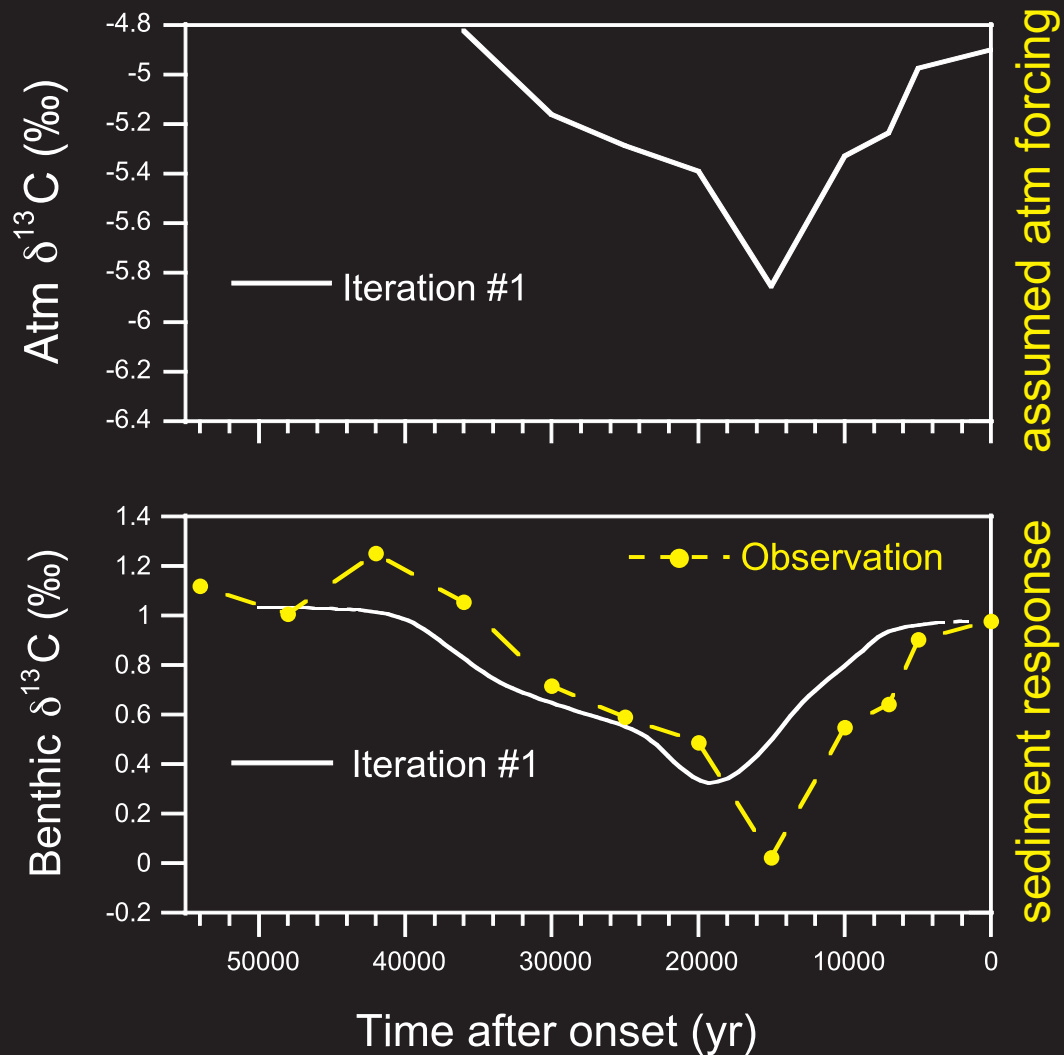




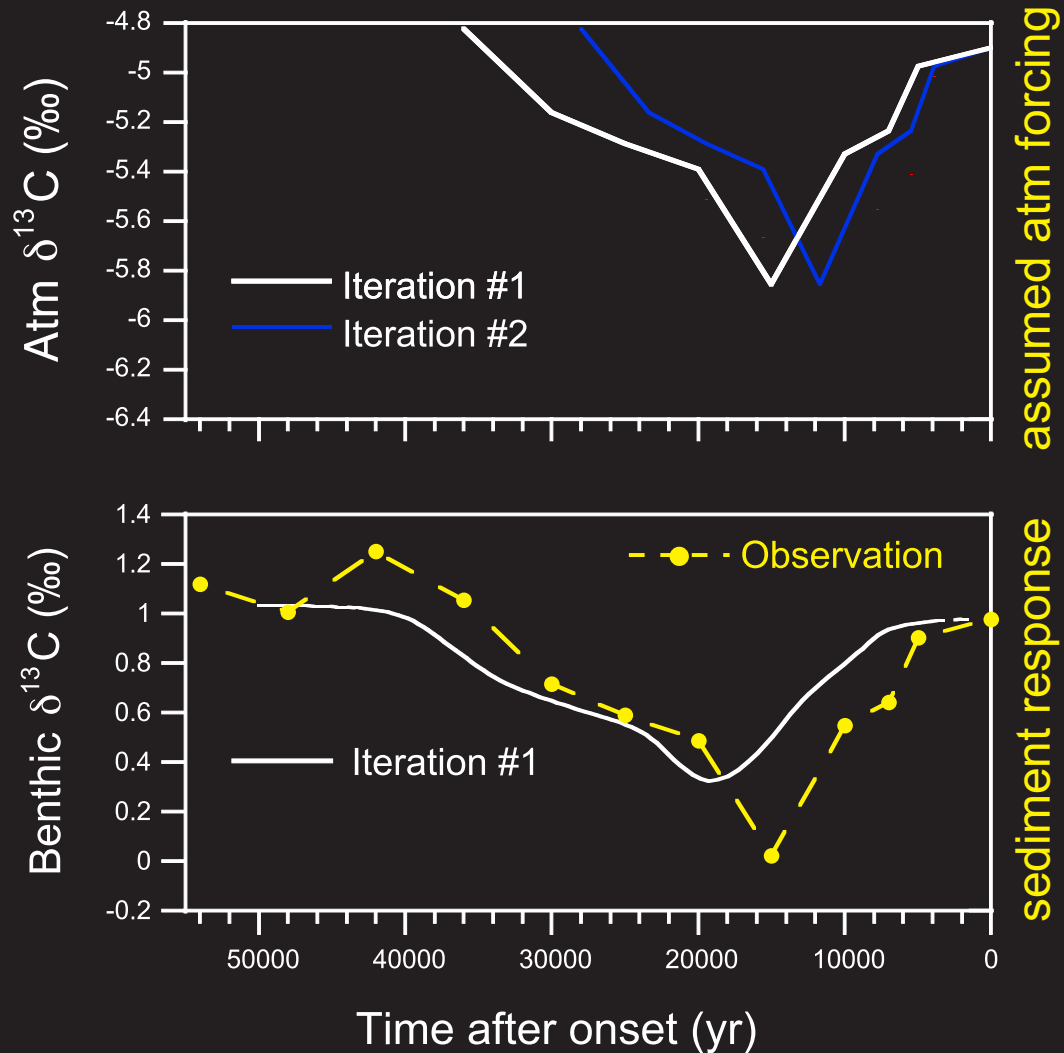




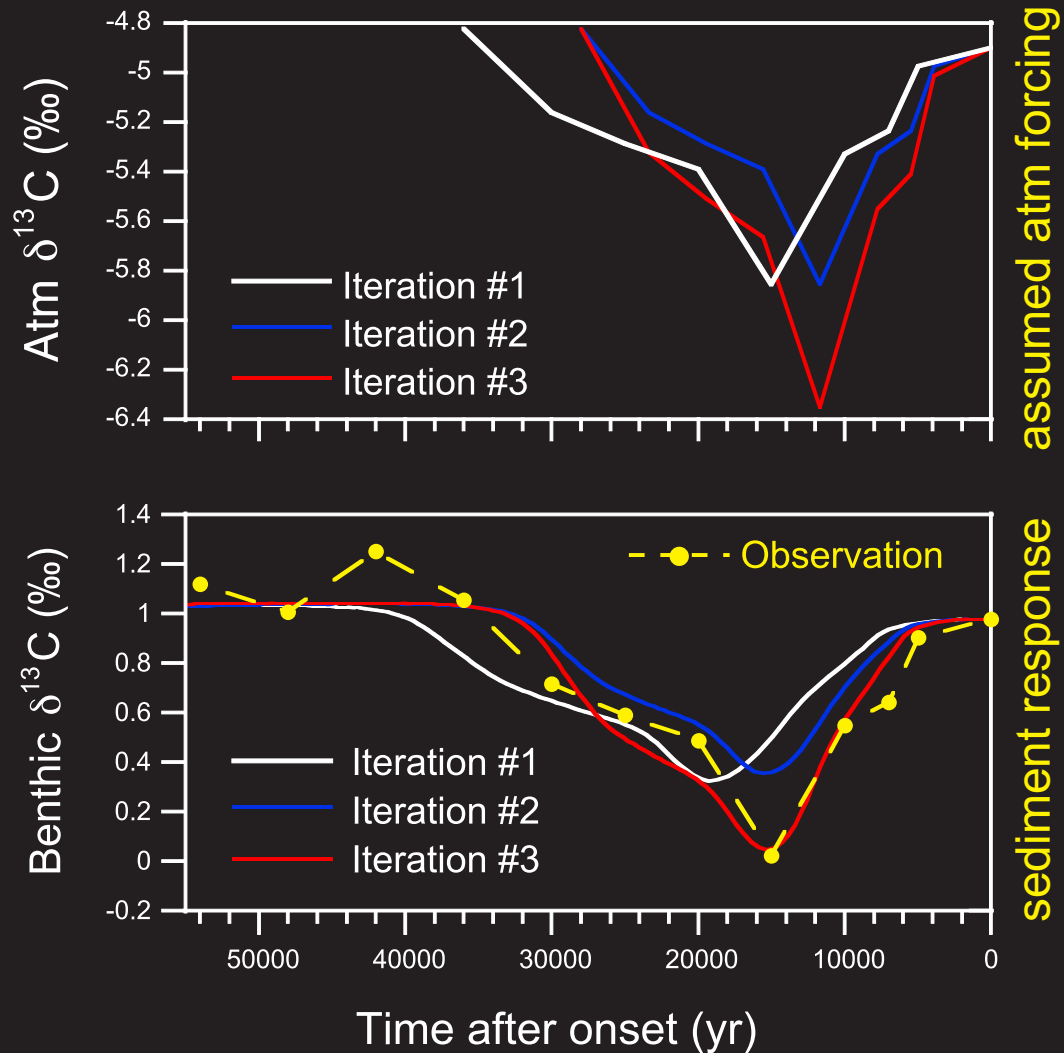
Initial guess:  
observed  $\delta^{13}\text{C}$  record  
==  
the atmospheric forcing



Step #1  
Invert 'guesstimated'  
atmospheric  $\delta^{13}\text{C}$  record and  
calculate sediment expression



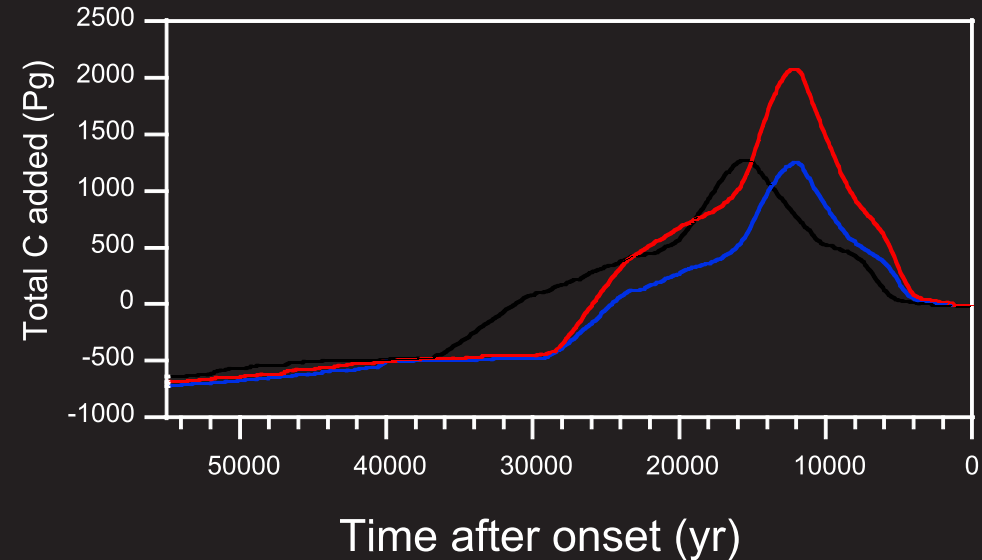
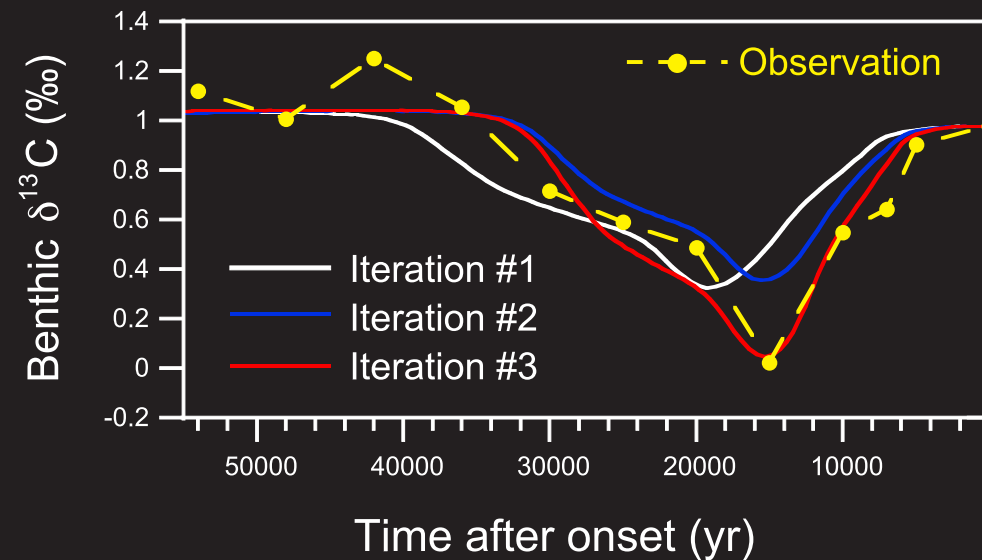
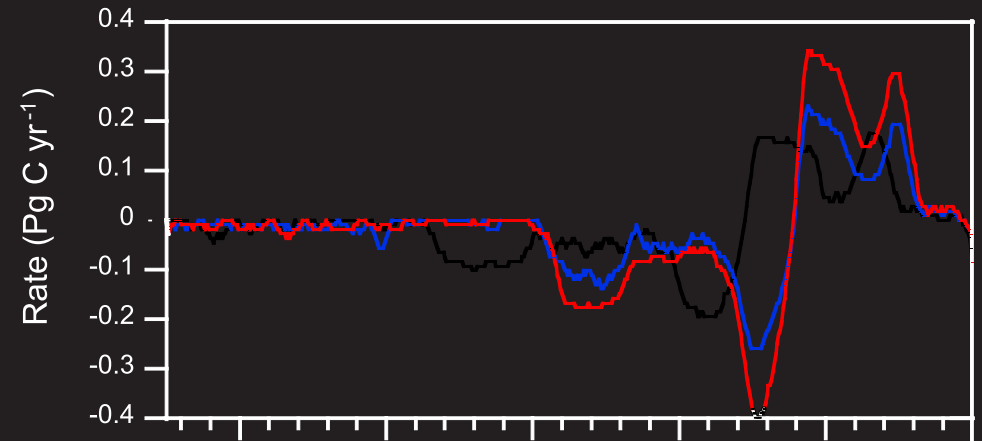
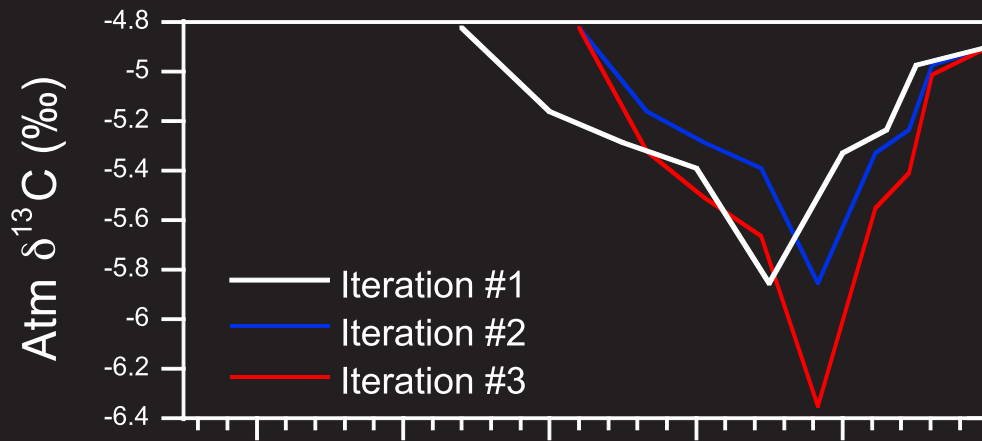
Adjust atmospheric record:  
Correct for distortion in time



Step #2:  
Invert adjusted  
atmospheric  $\delta^{13}\text{C}$  record;  
then adjust forcing magnitude;

Step #3:  
Invert the now twice-adjusted  
atmospheric  $\delta^{13}\text{C}$  record

Recover rates of CO<sub>2</sub> emissions

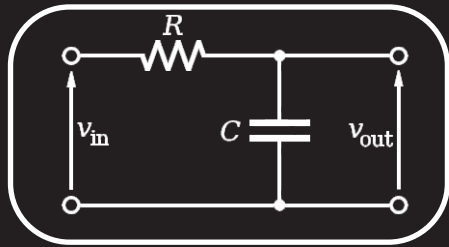




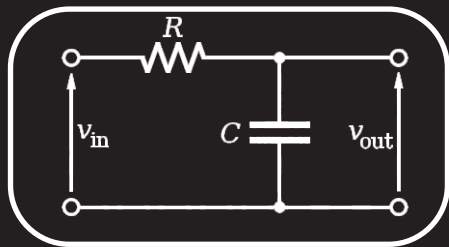
# 'Inverting' isotopic records iteratively



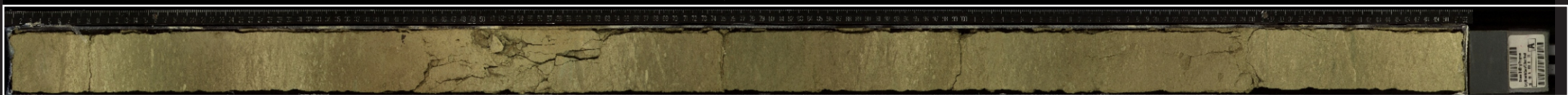
dissolution  
(preservation)



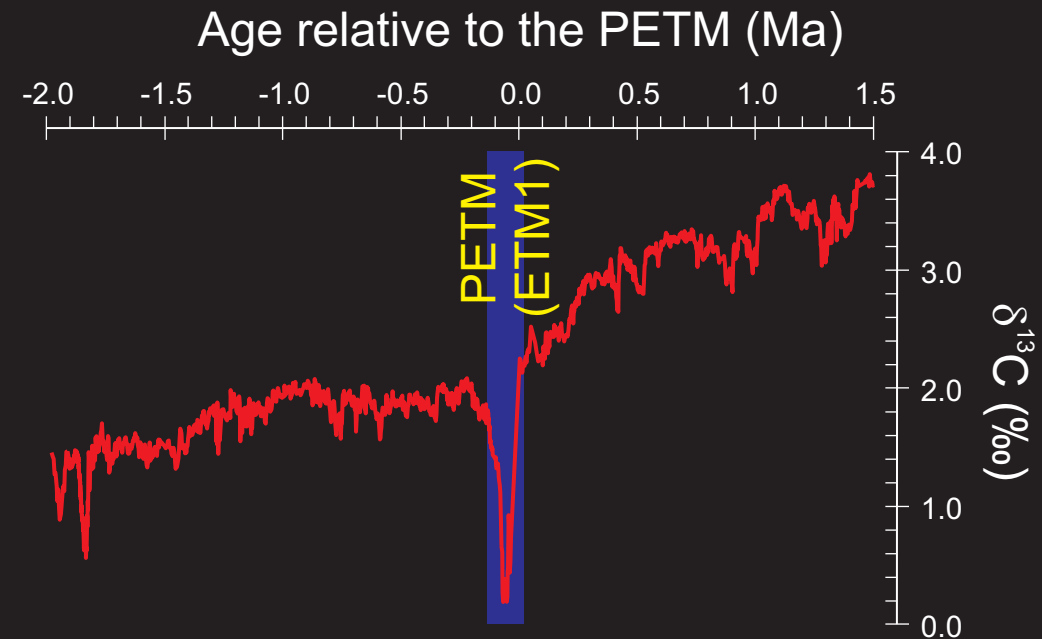
mixing  
(bioturbation)

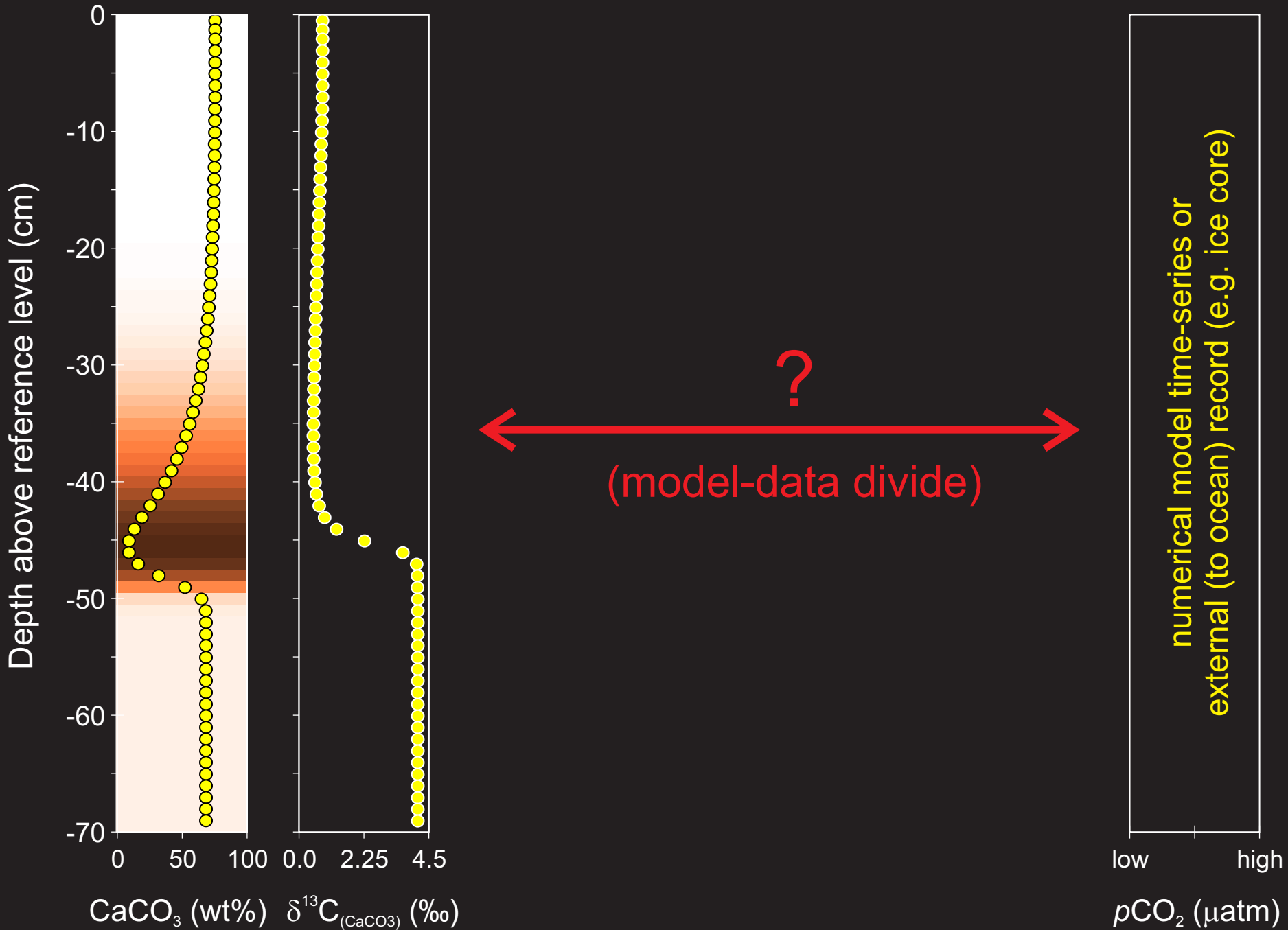


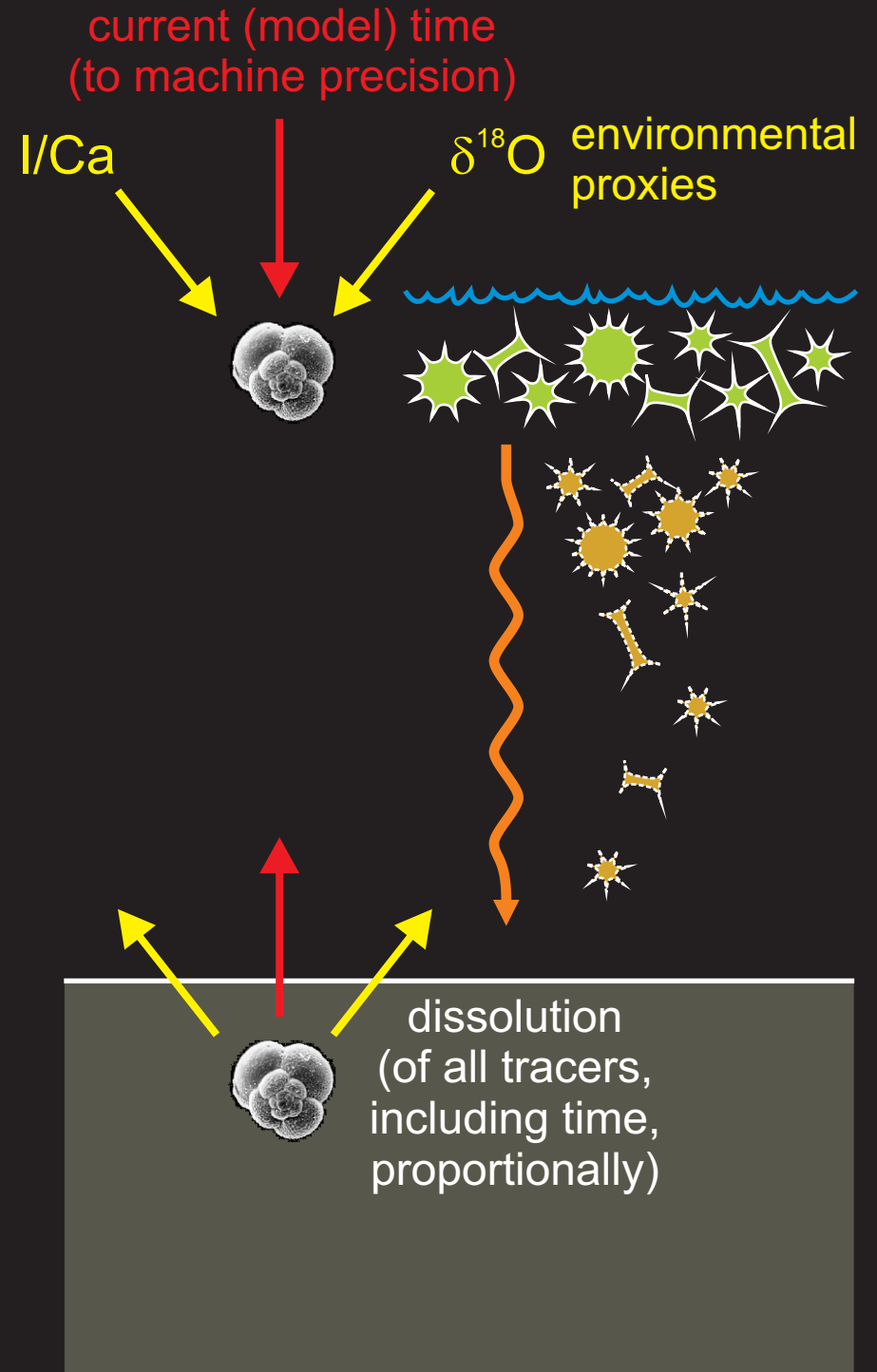
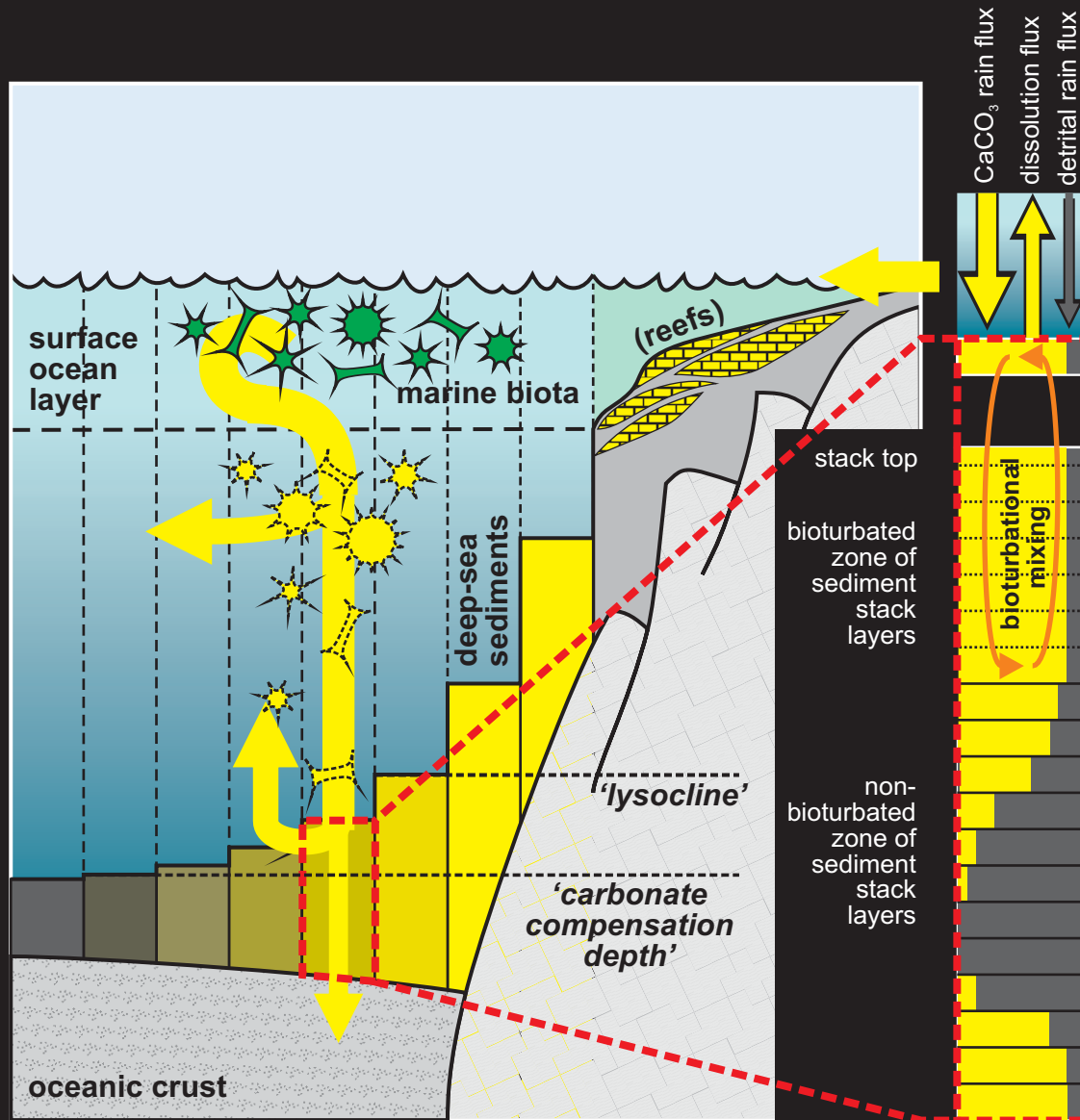
**C22nH3**  
(**'Big Bird'**)

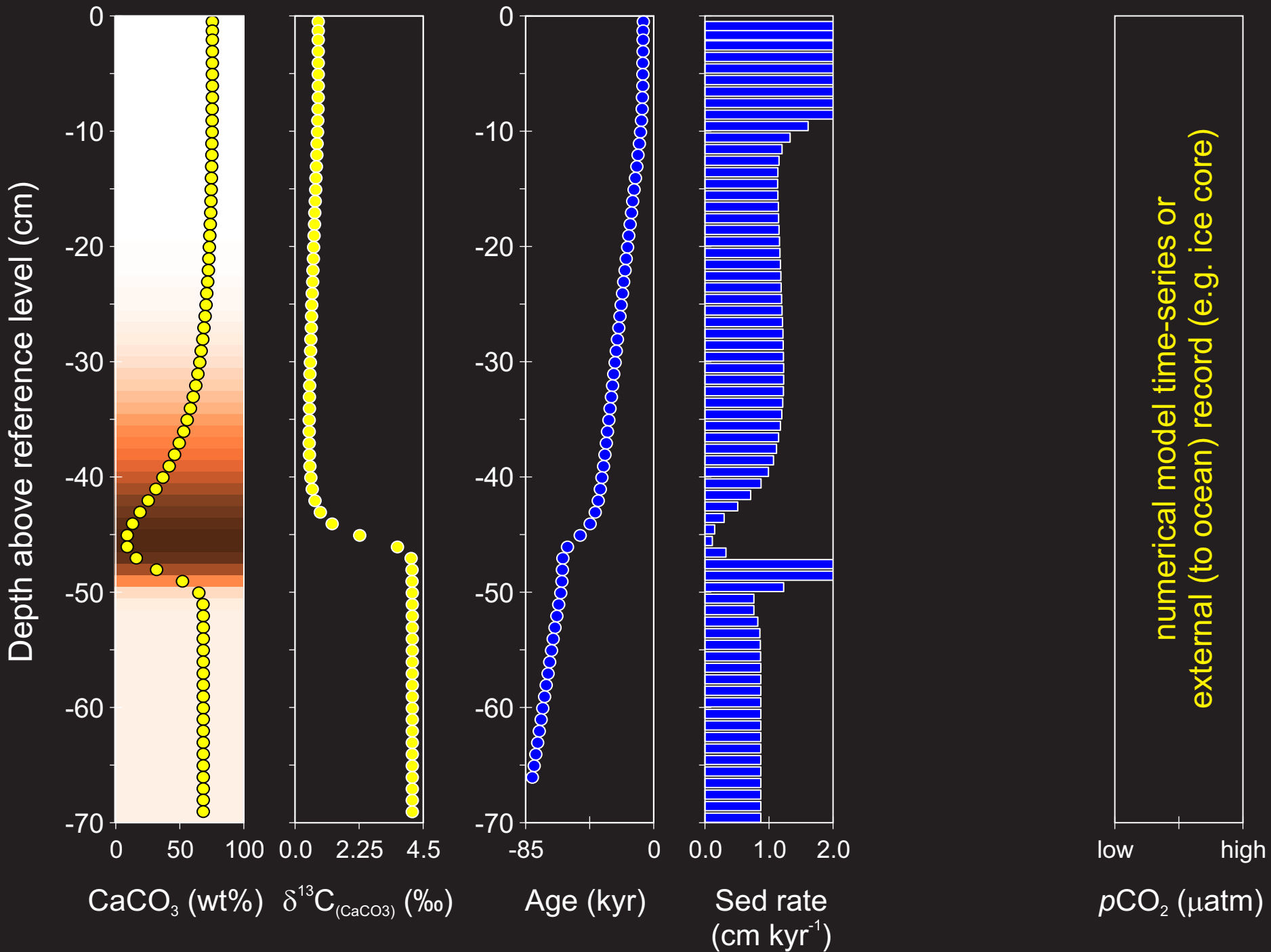


Consider: An event characterized by a (severe) reduction in carbonate preservation

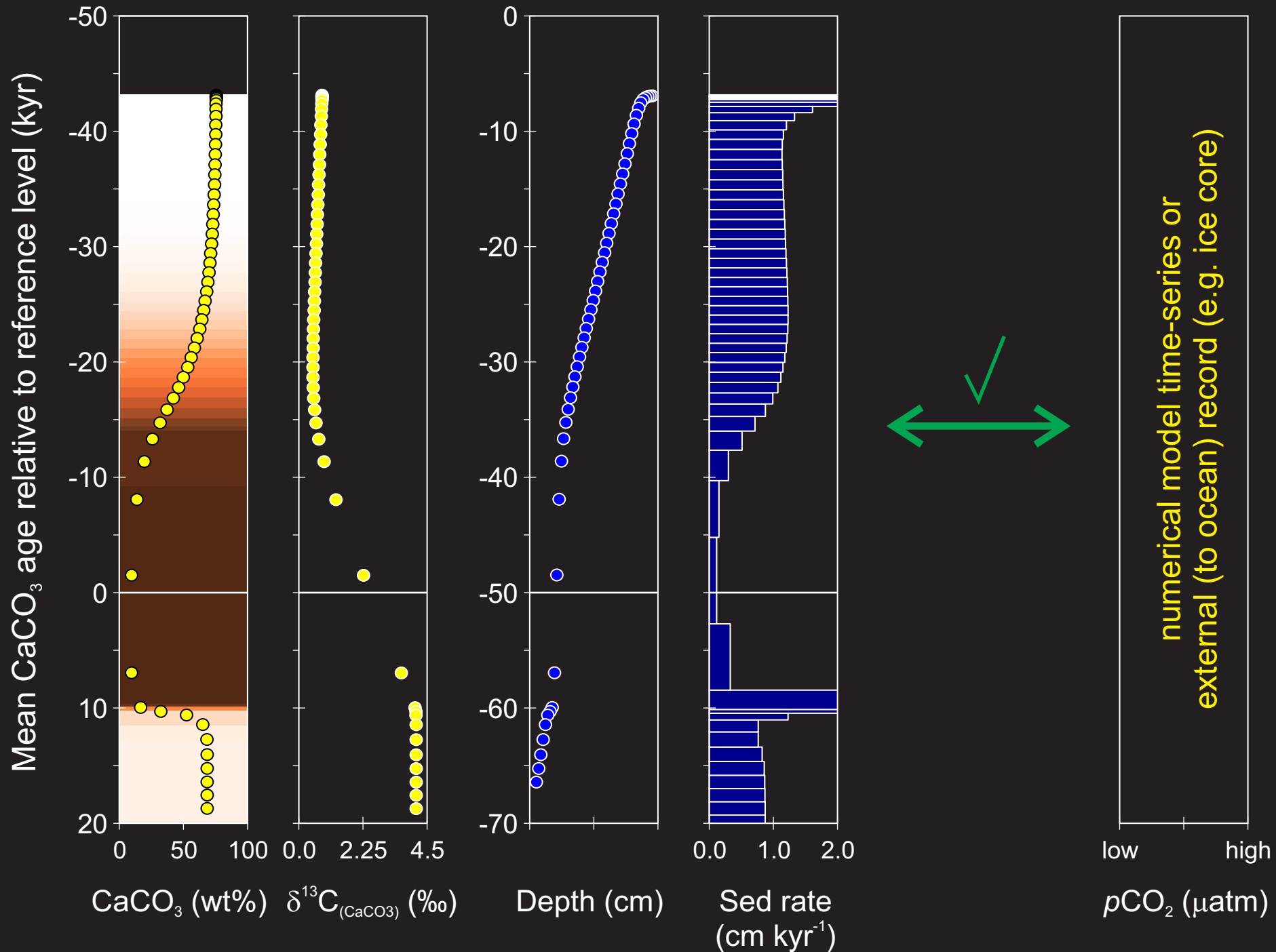


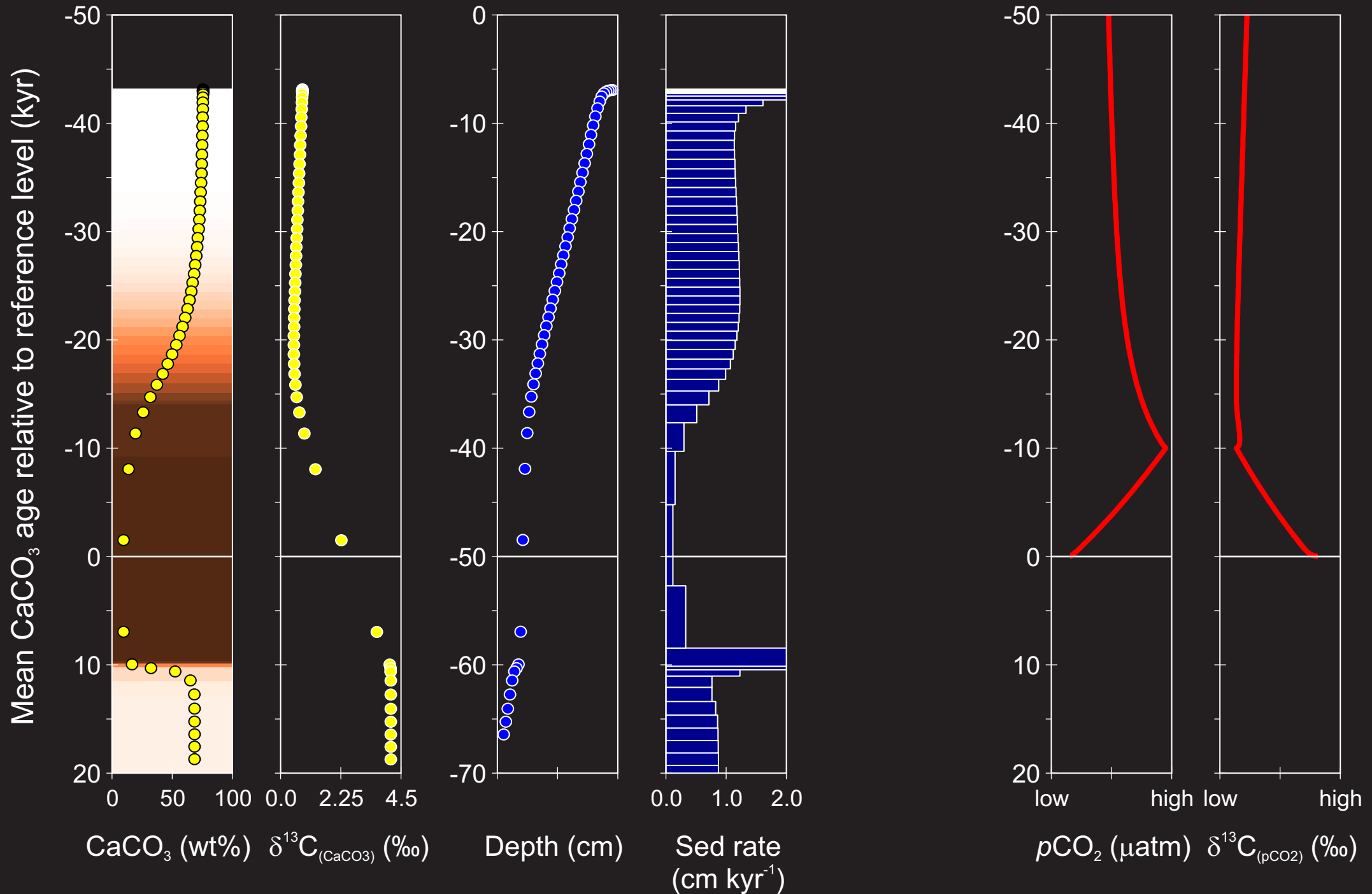




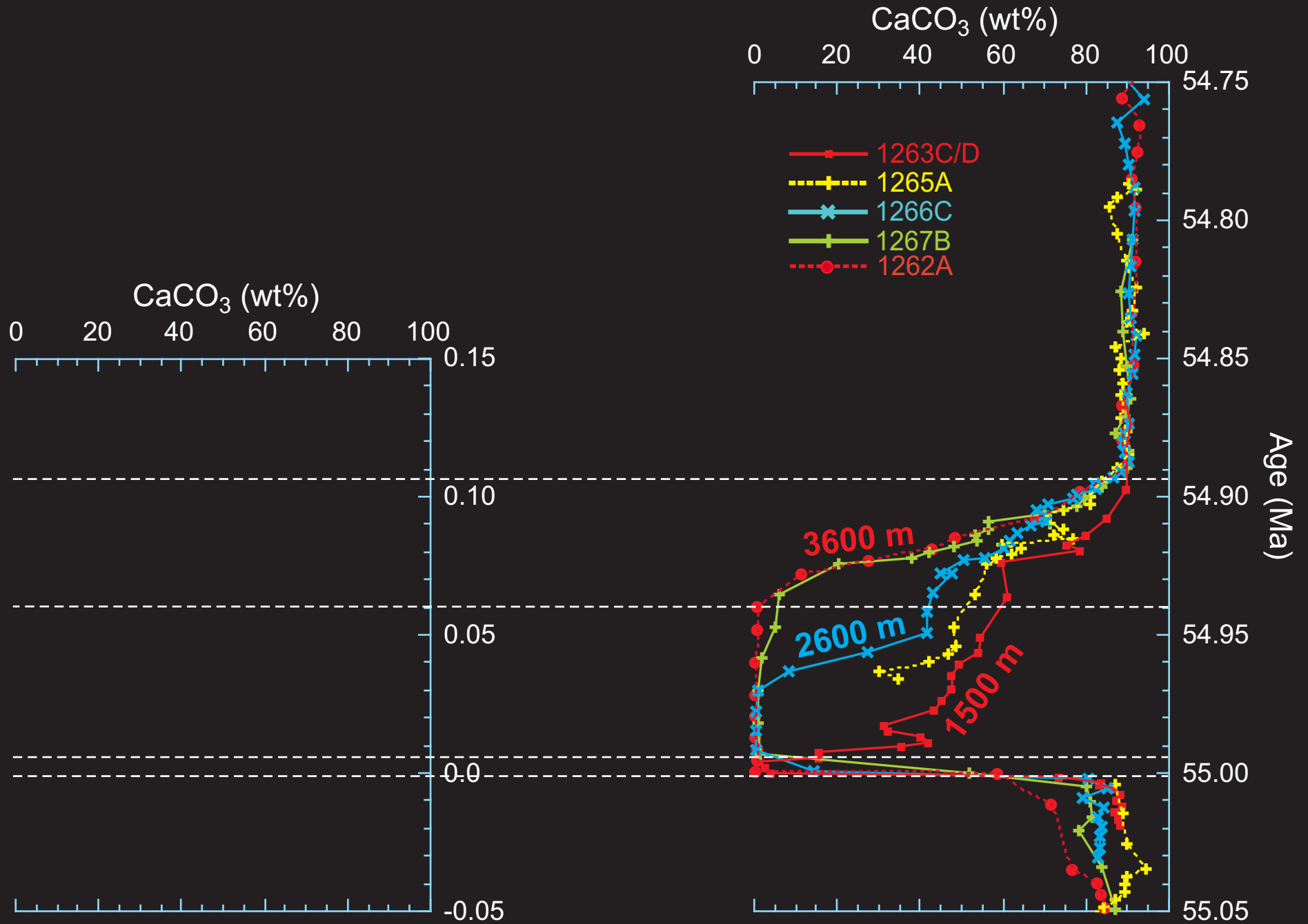








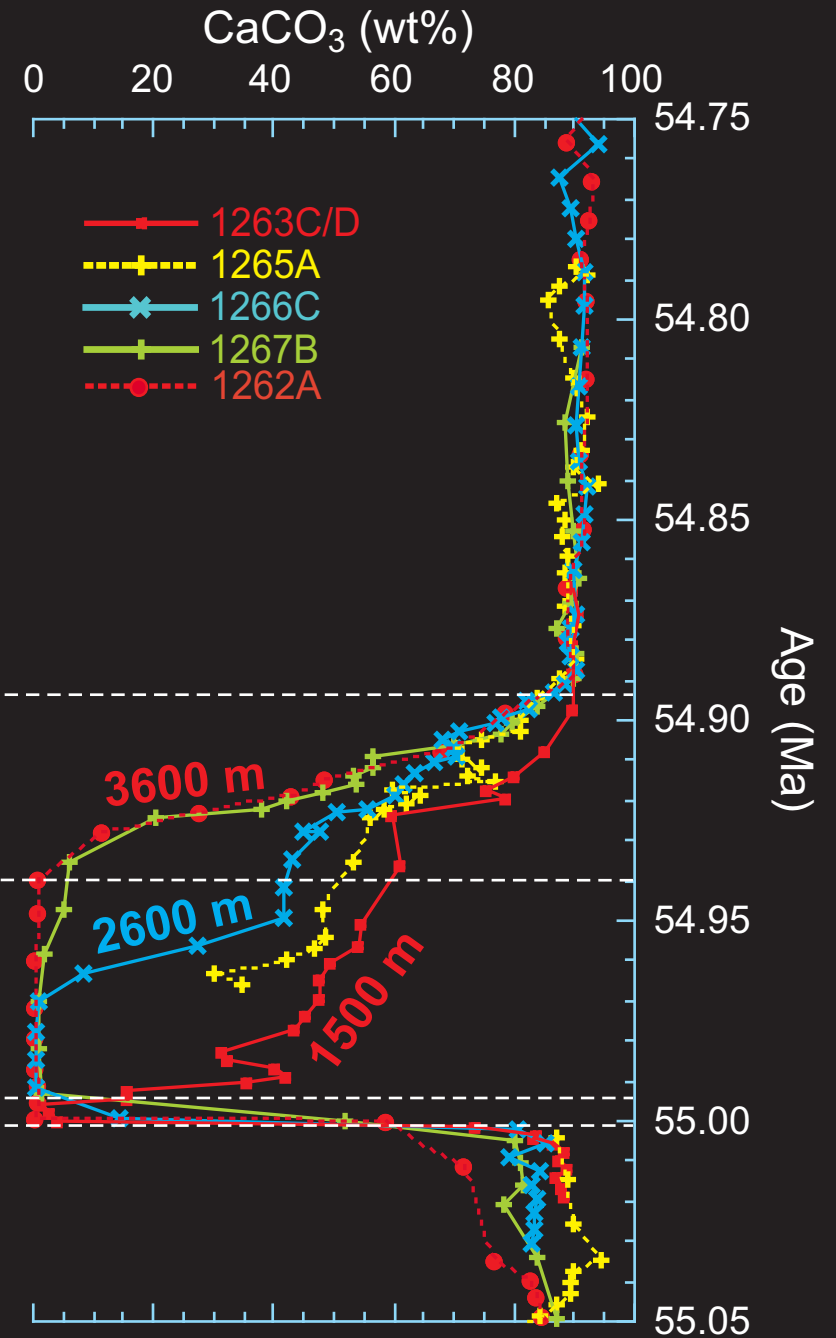
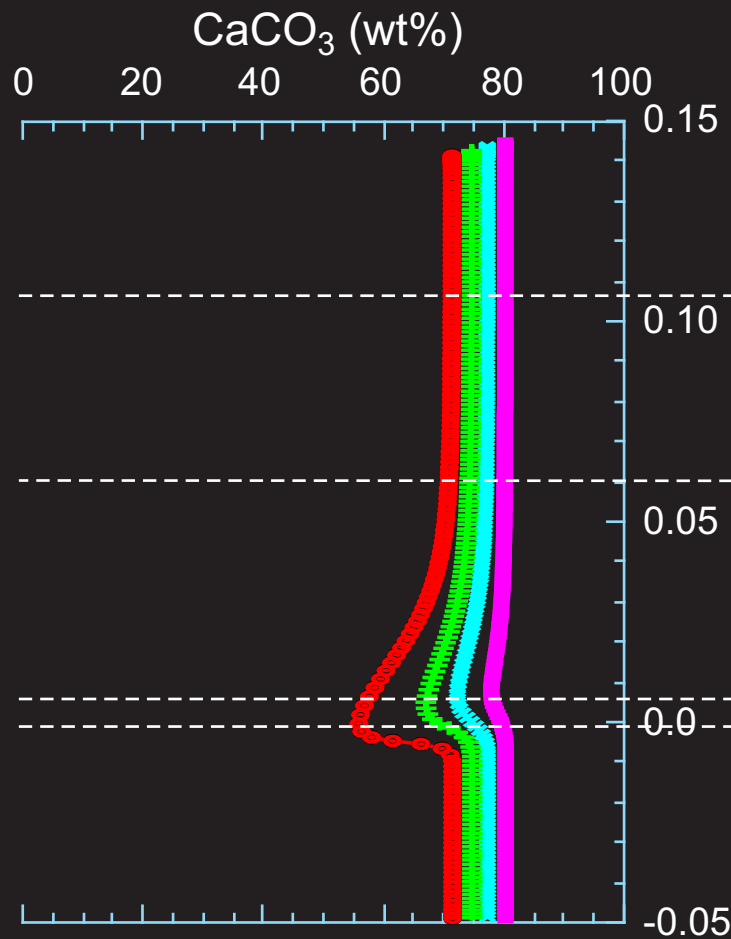
Model-generated synthetic sediment core response [Ridgwell, 2007]



Bulk sediment wt% CaCO<sub>3</sub> content [Zachos et al., 2005]

2000 PgC CO<sub>2</sub> perturbation

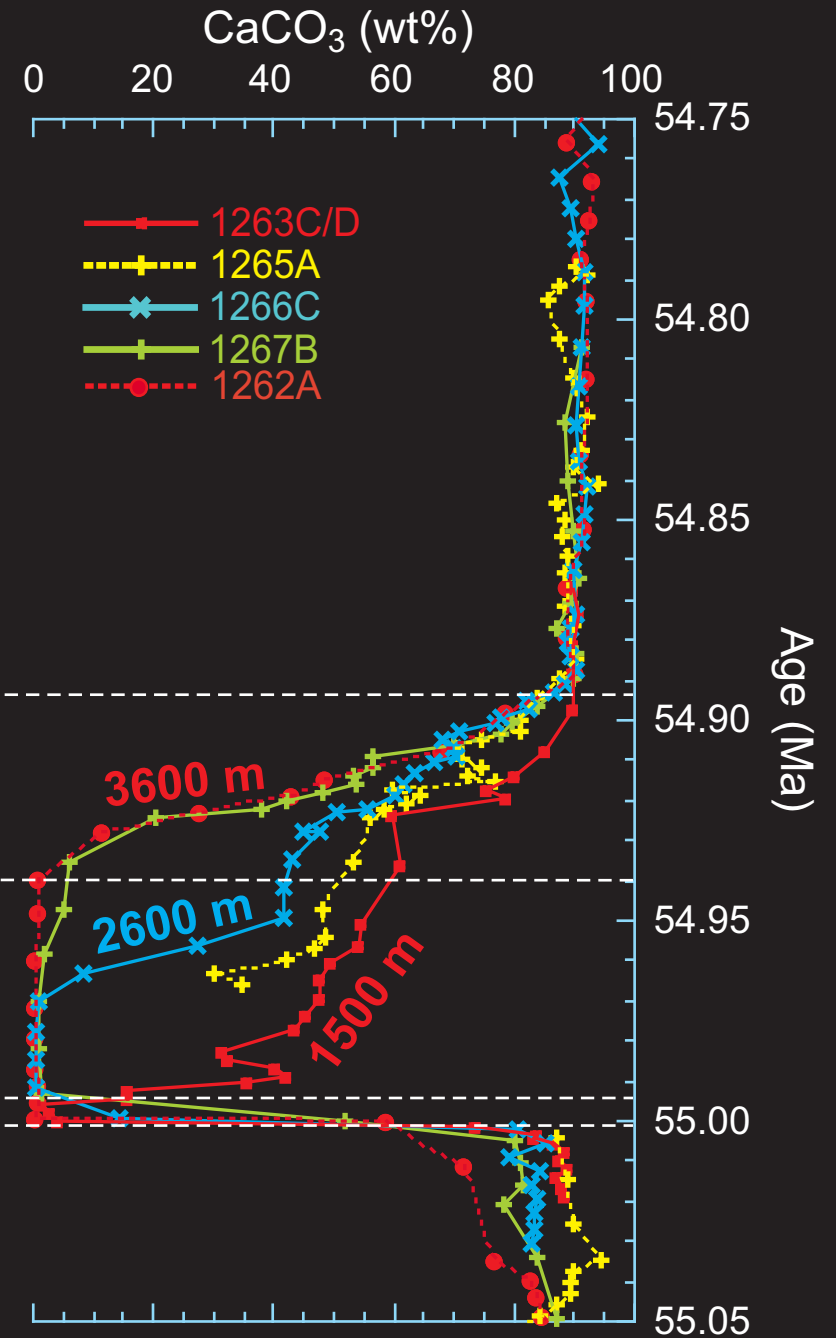
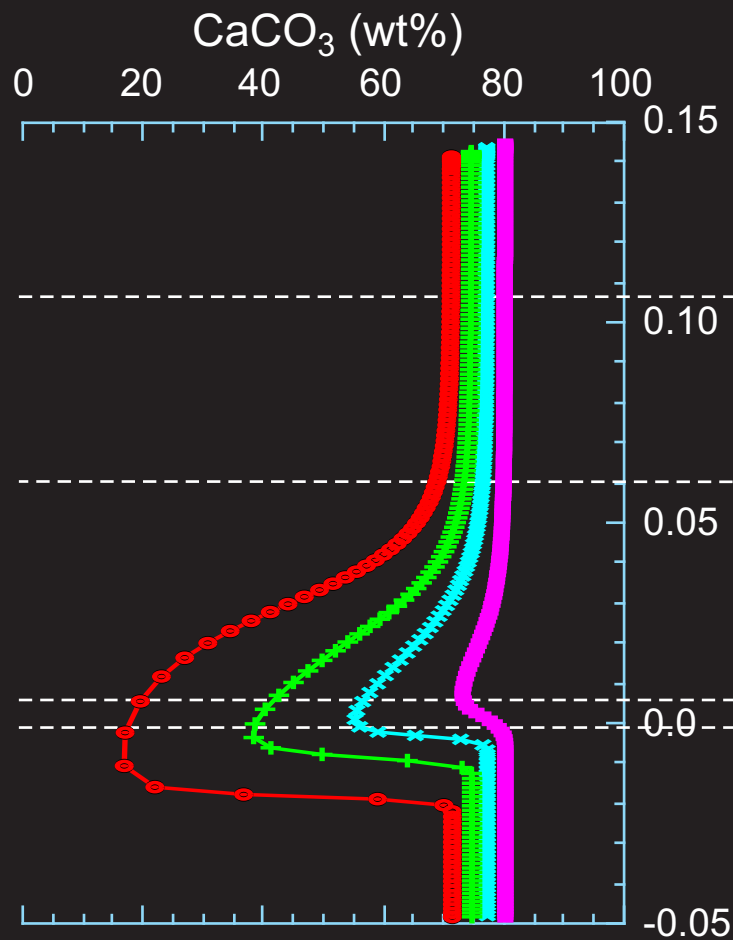
Model-generated synthetic sediment core response [Ridgwell, 2007]



Bulk sediment wt% CaCO<sub>3</sub> content [Zachos et al., 2005]

4000 PgC CO<sub>2</sub> perturbation

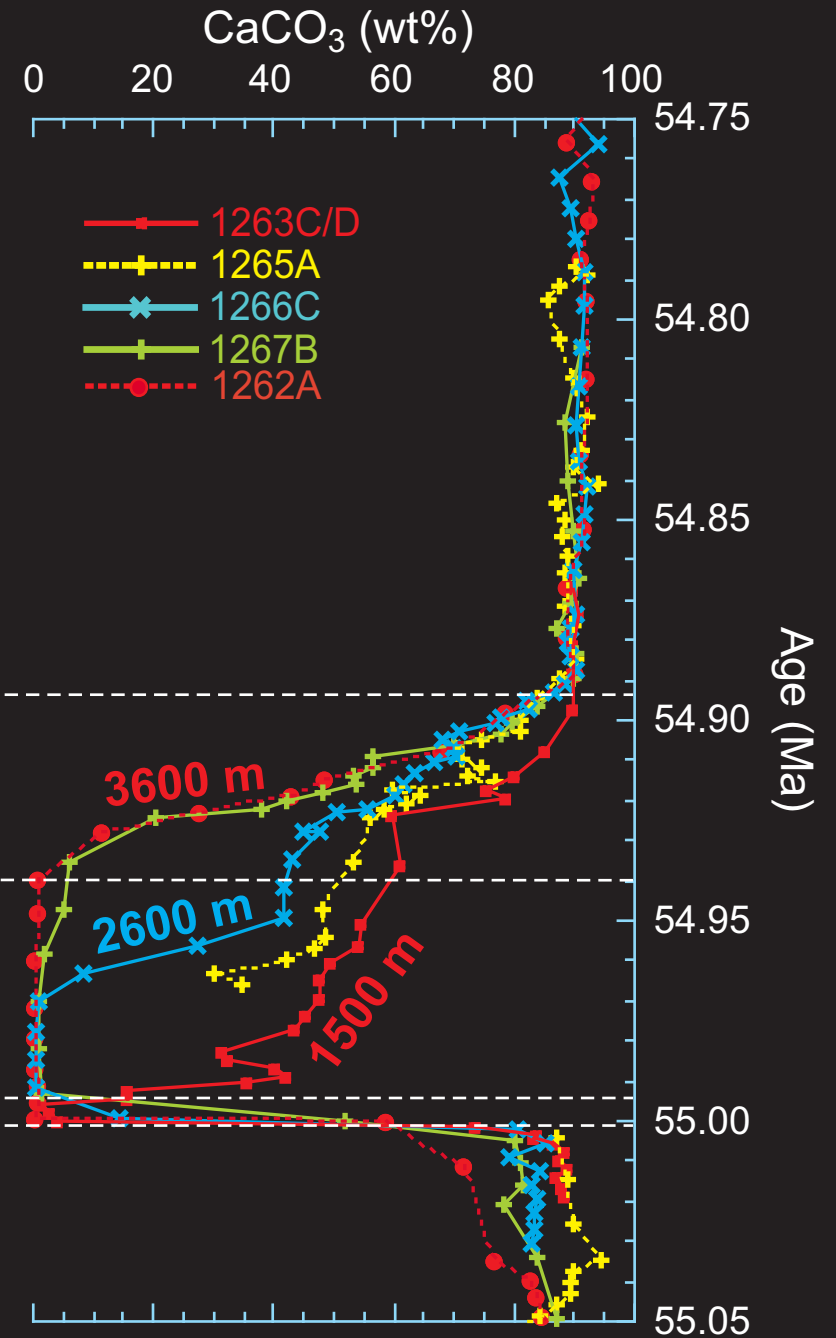
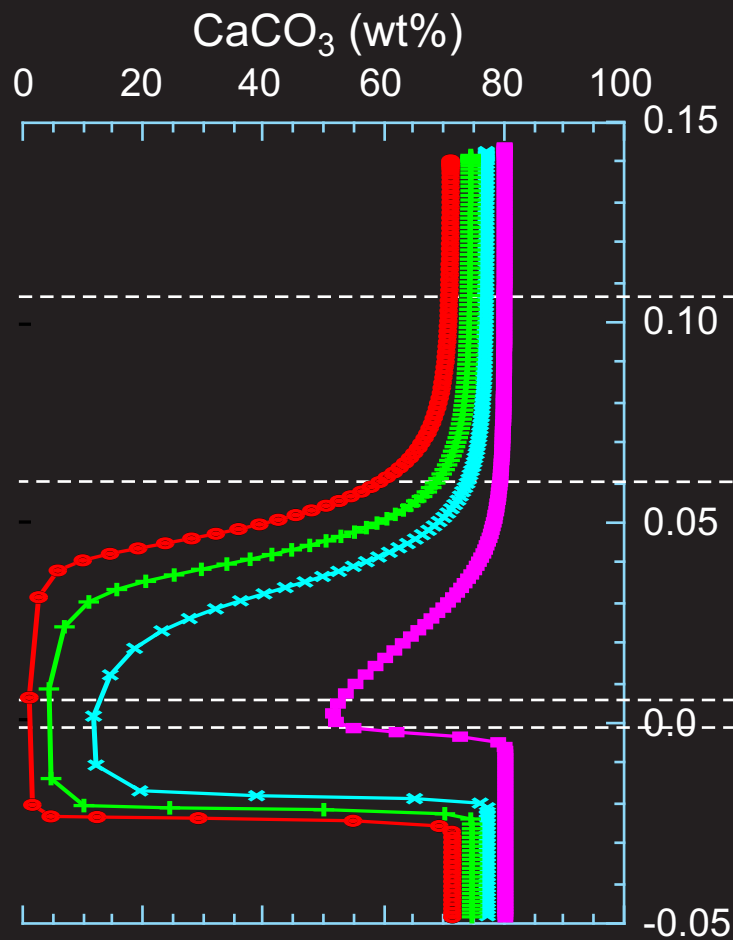
Model-generated synthetic sediment core response [Ridgwell, 2007]



Bulk sediment wt% CaCO<sub>3</sub> content [Zachos et al., 2005]

6000 PgC CO<sub>2</sub> perturbation

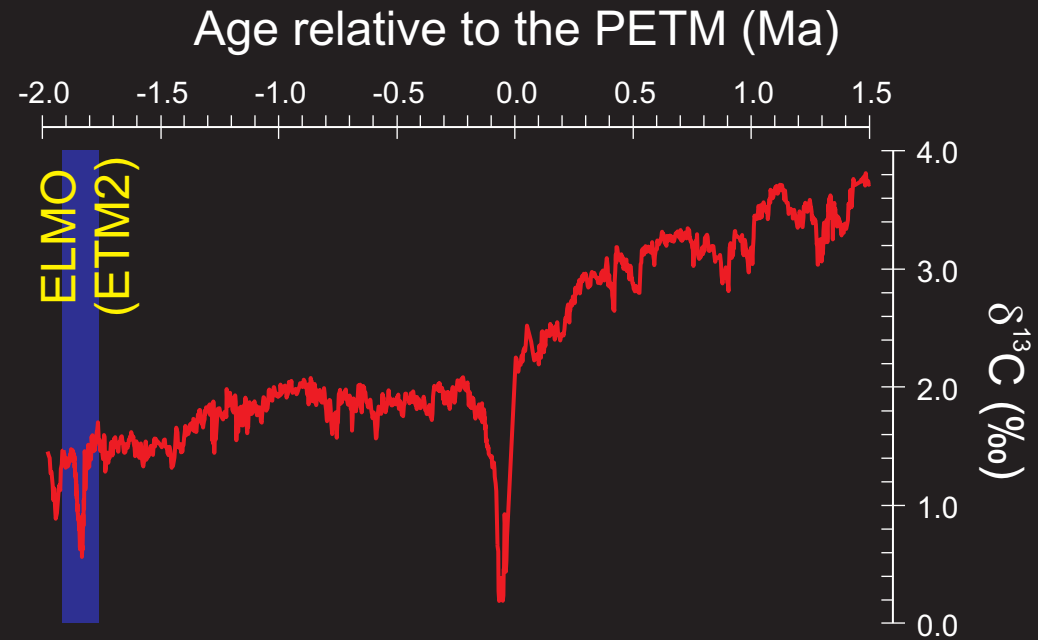
Model-generated synthetic sediment core response [Ridgwell, 2007]

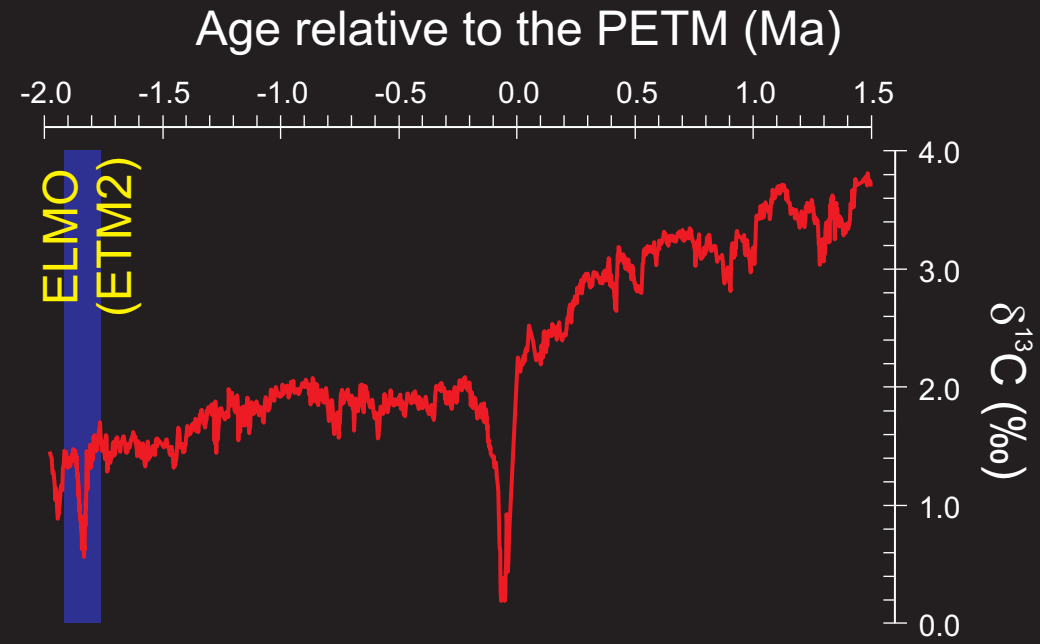
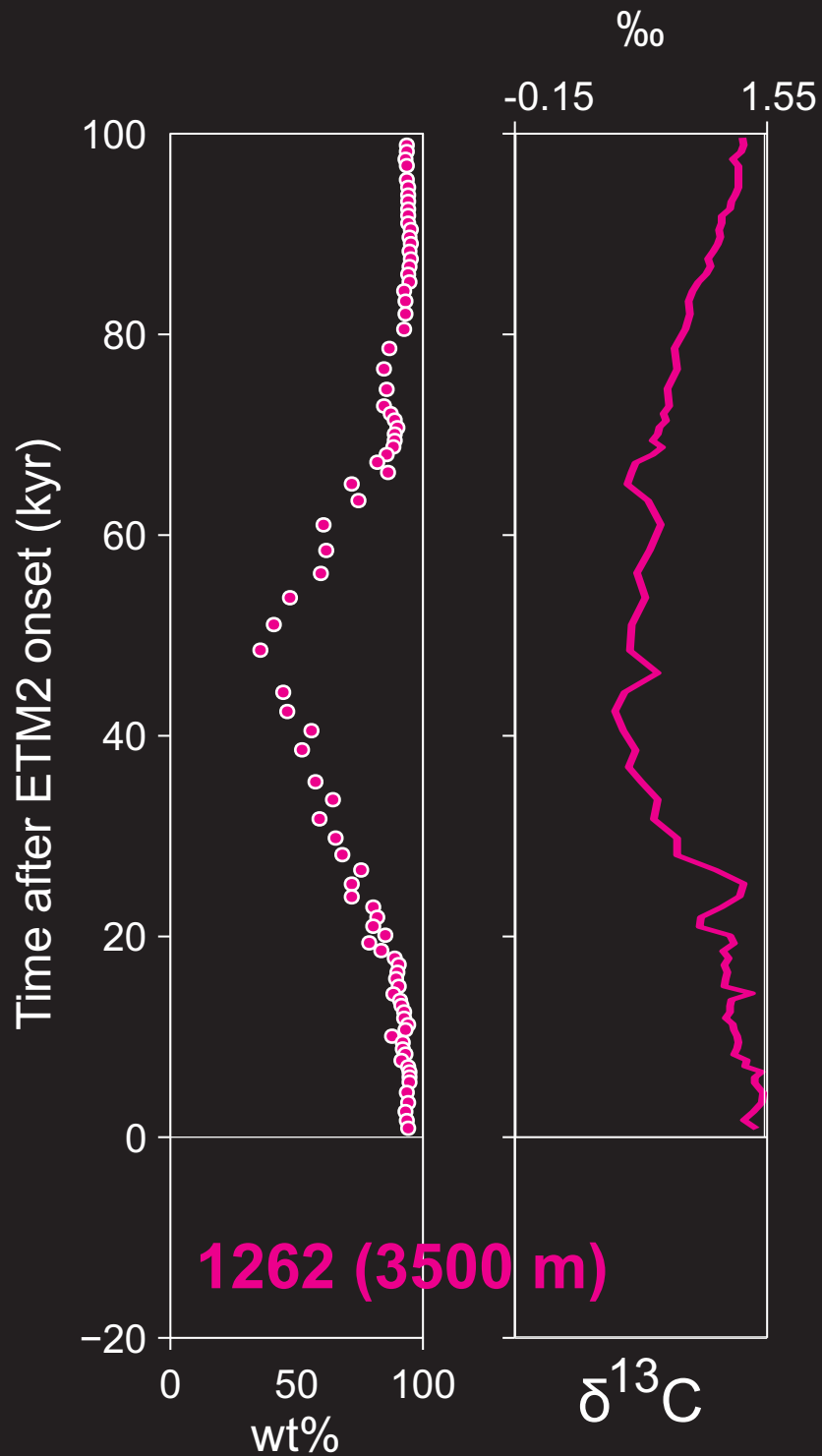


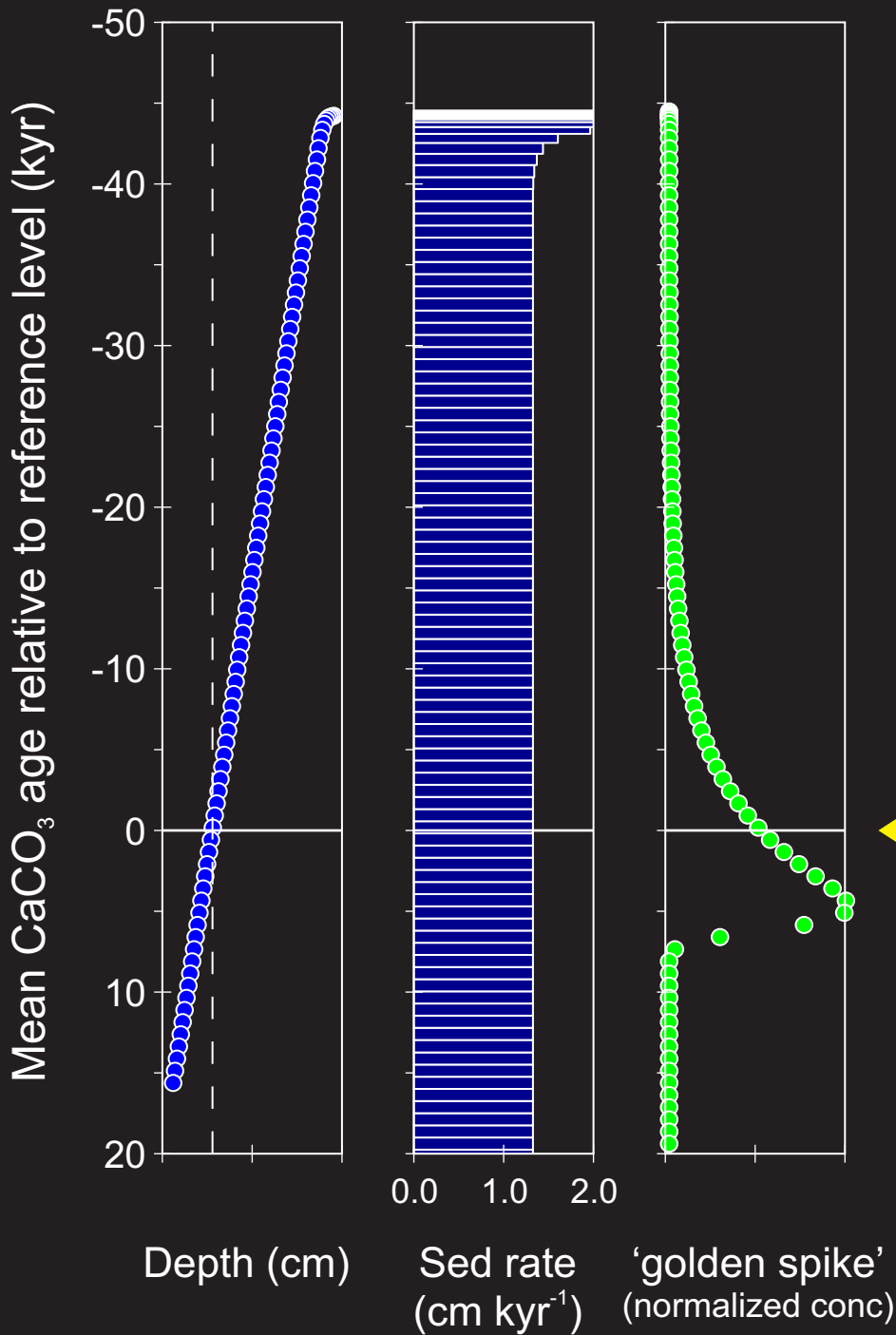
Bulk sediment wt% CaCO<sub>3</sub> content [Zachos et al., 2005]



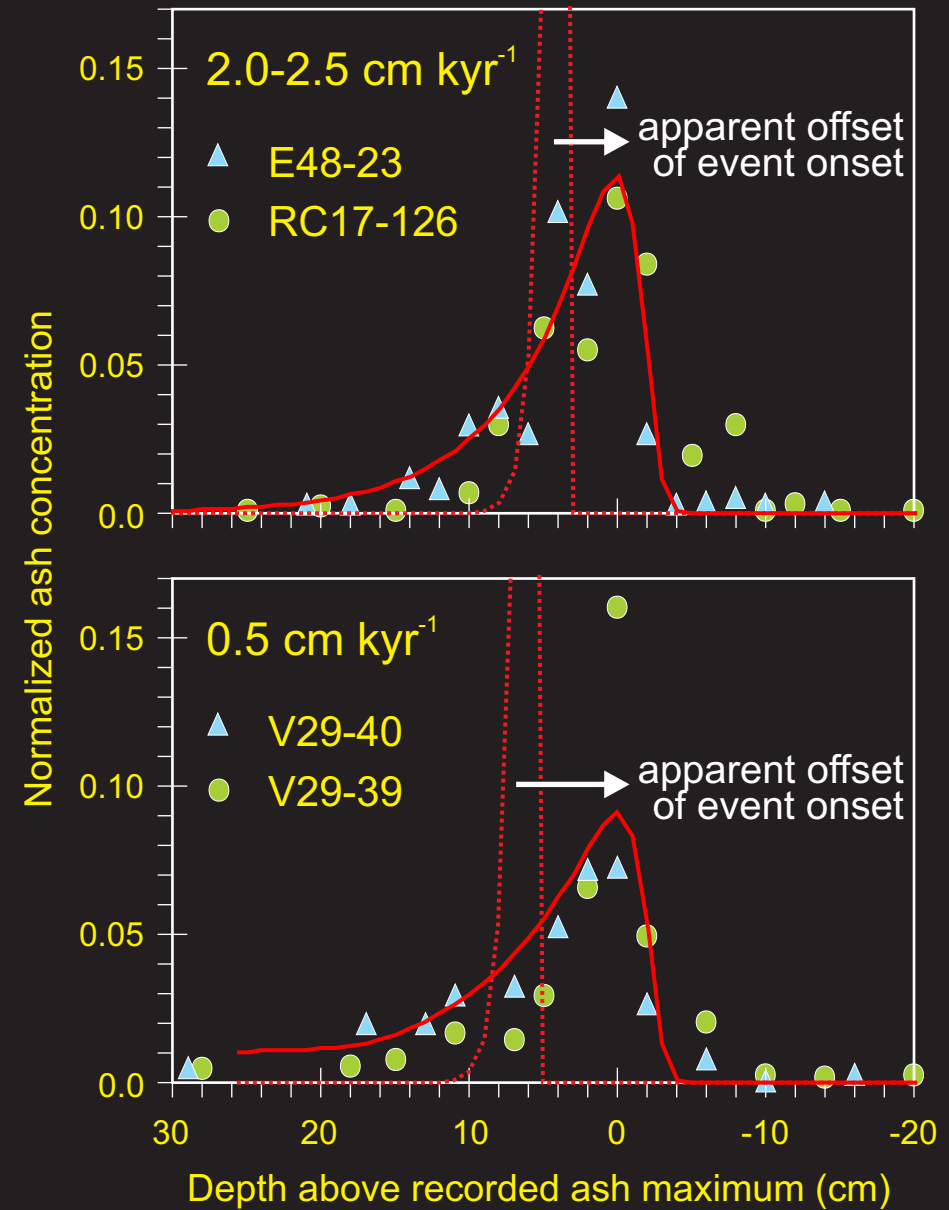
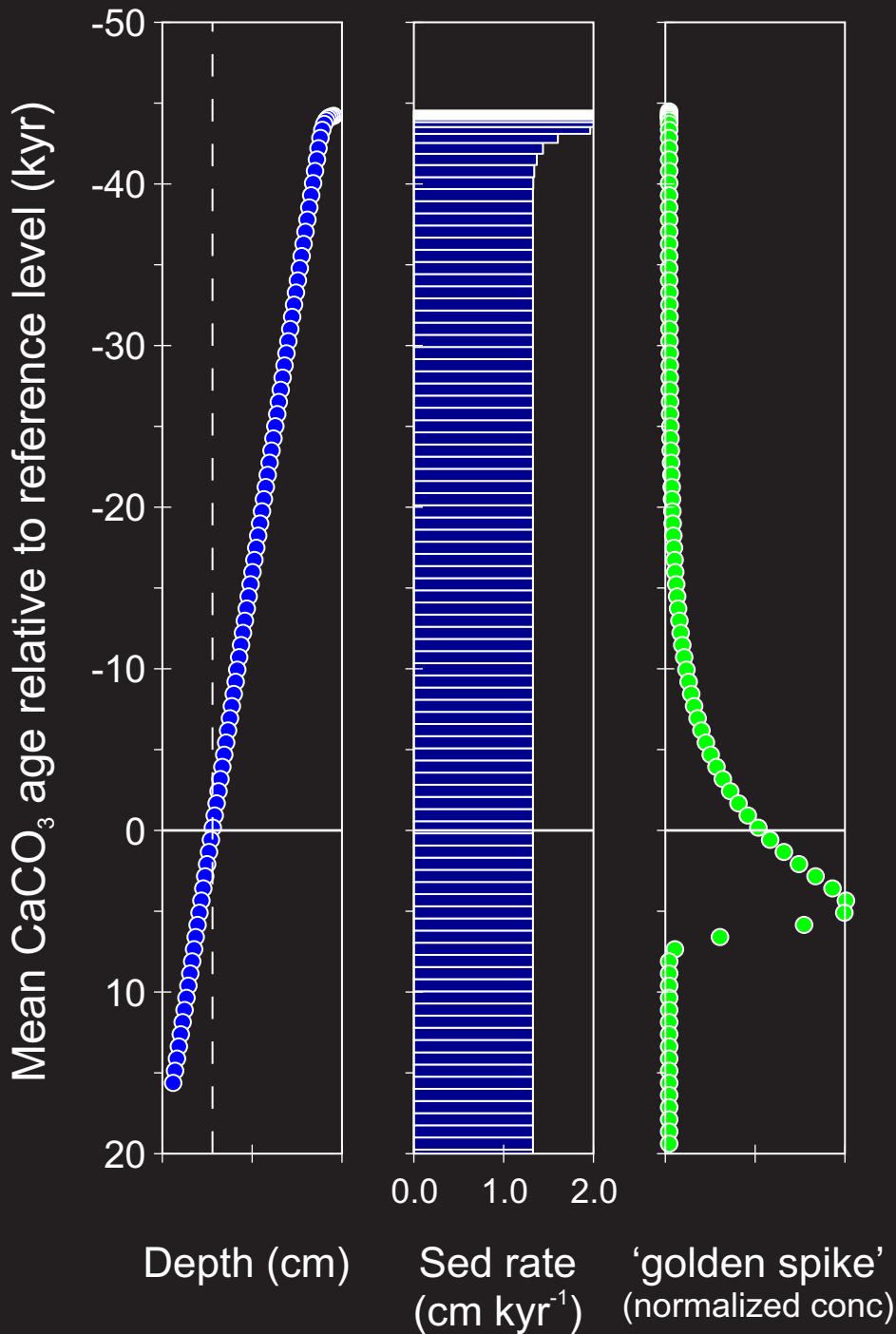
Consider: An event characterized by a (mild) reduction in carbonate preservation

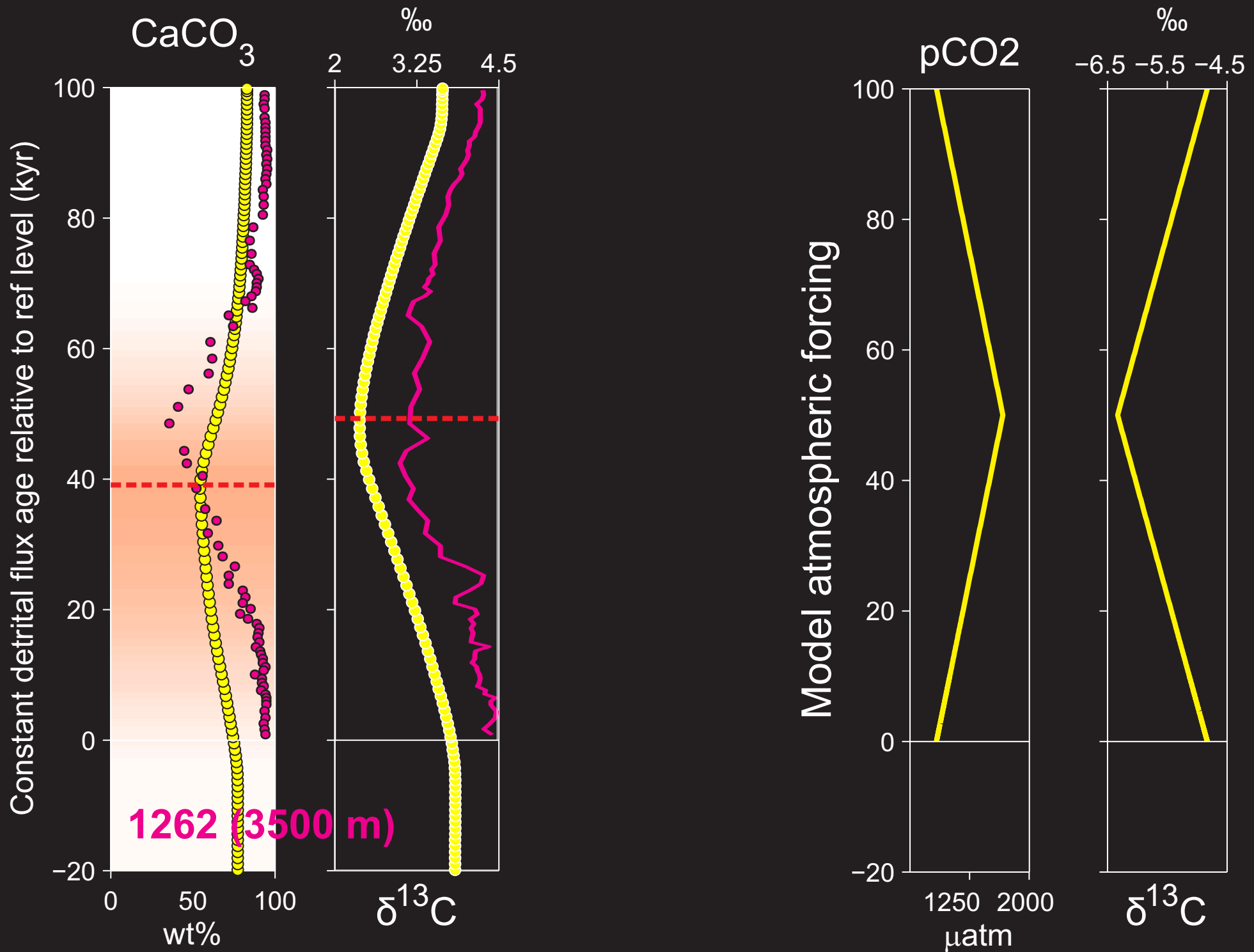


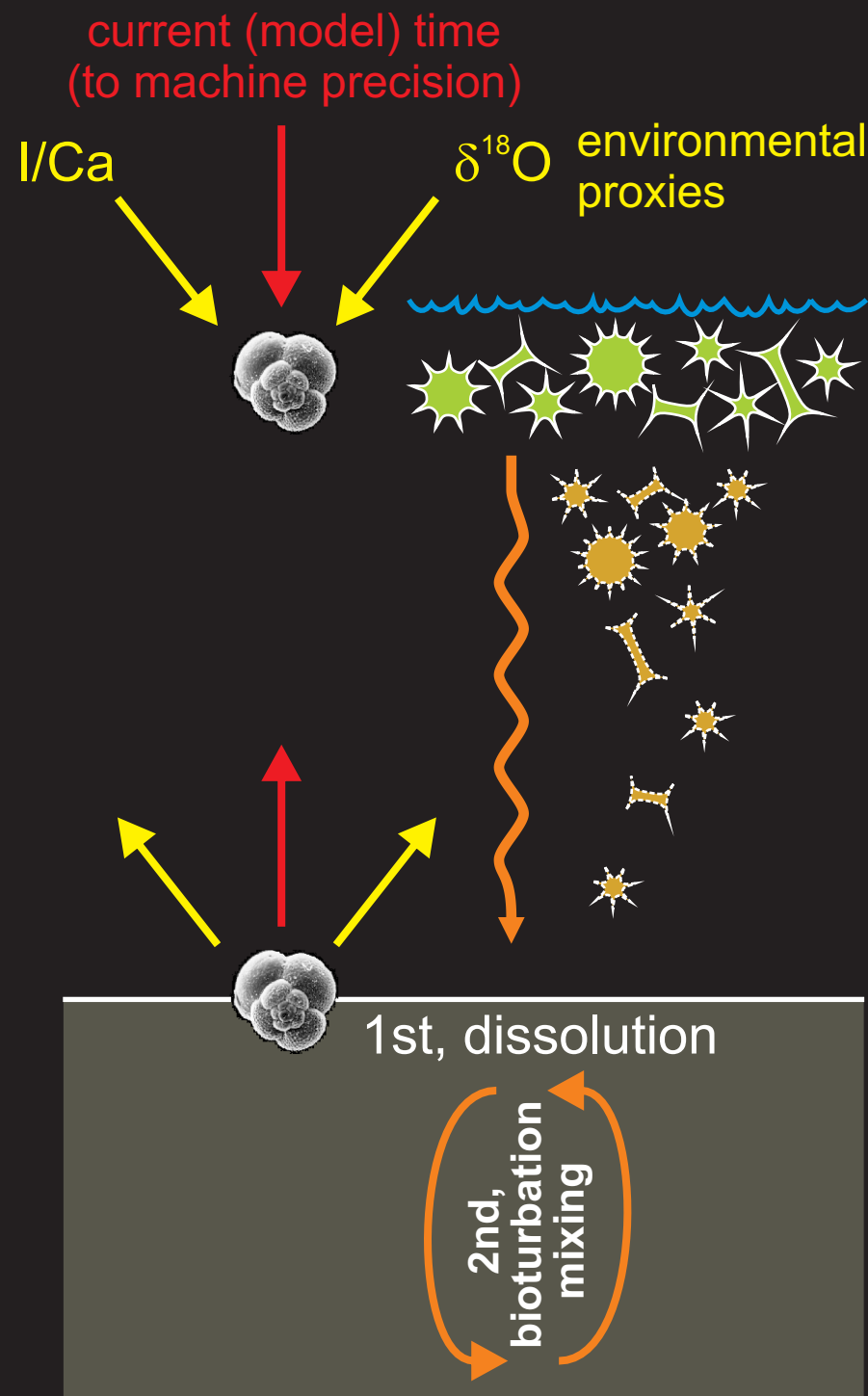
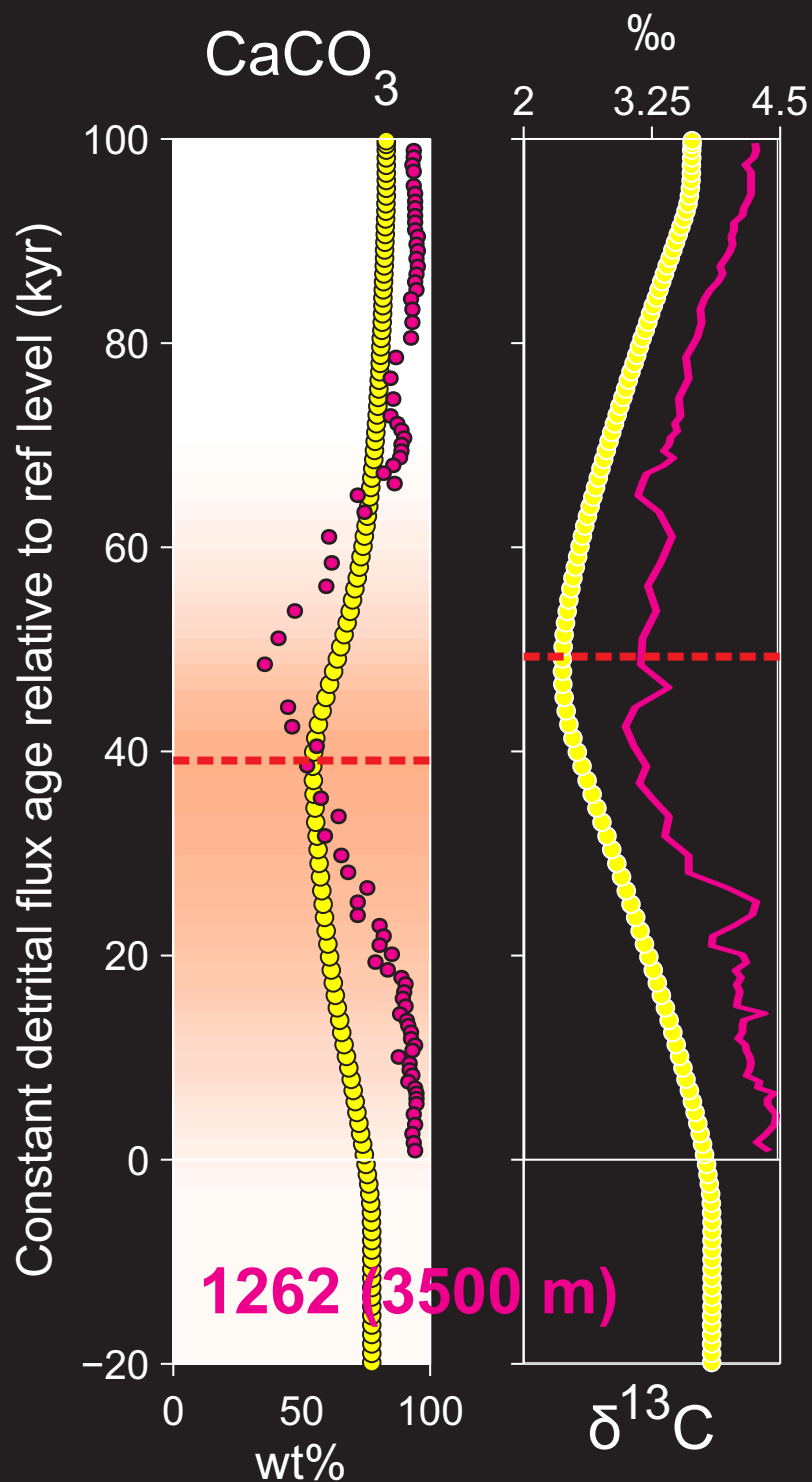


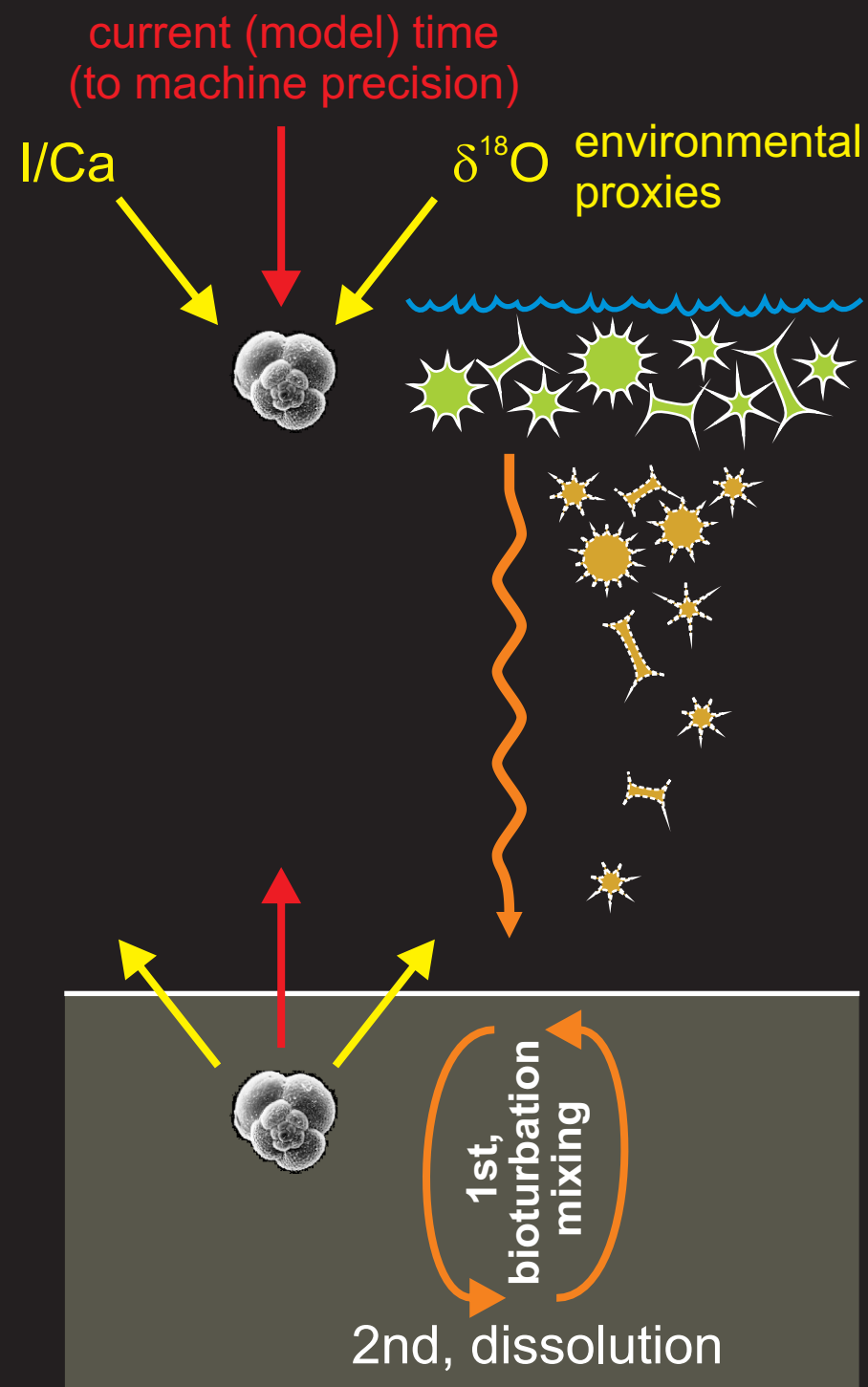


model experiment starts, surface sediments 'tagged' (instantaneous pulse of inert, conservative, numerical tracer)



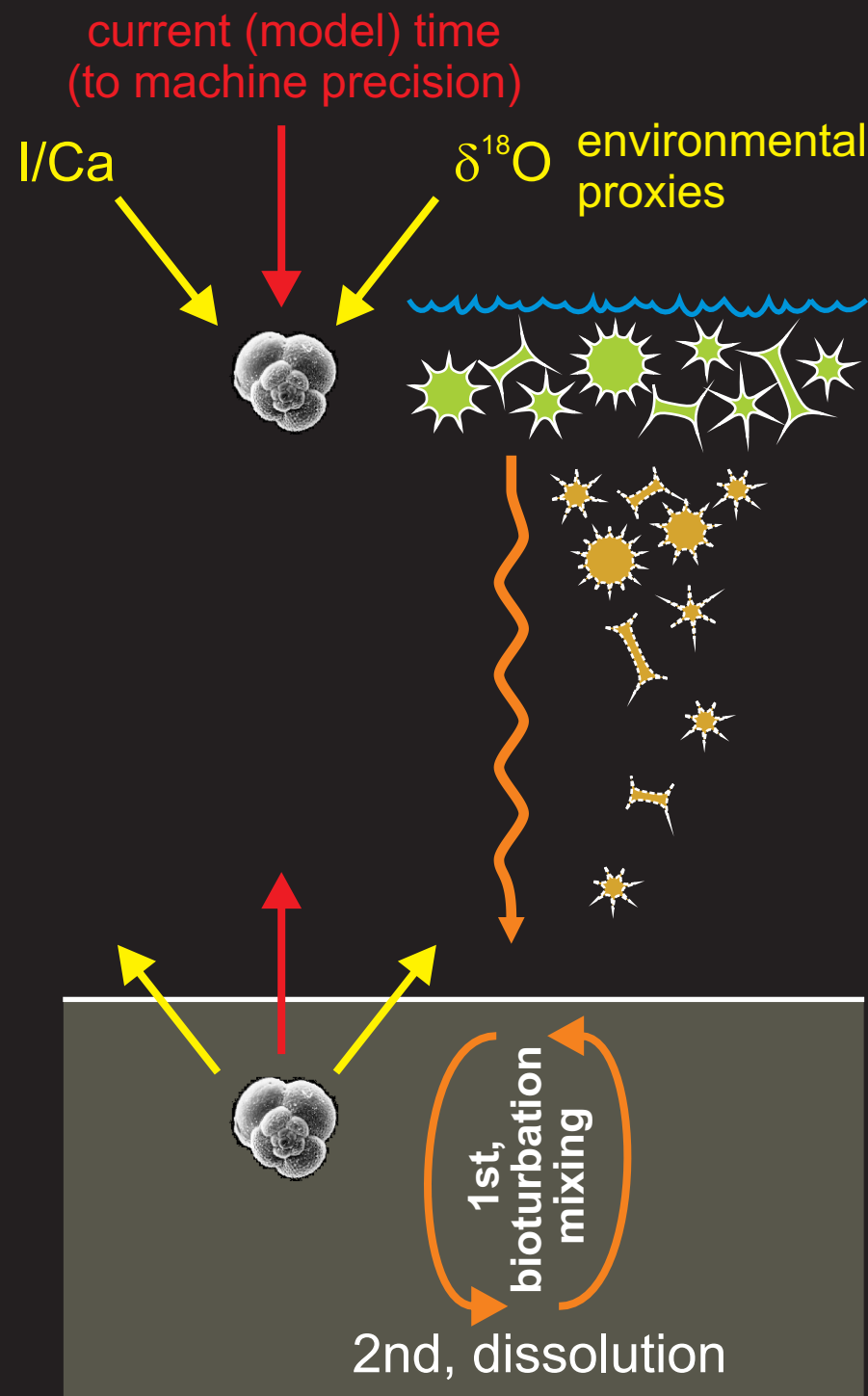
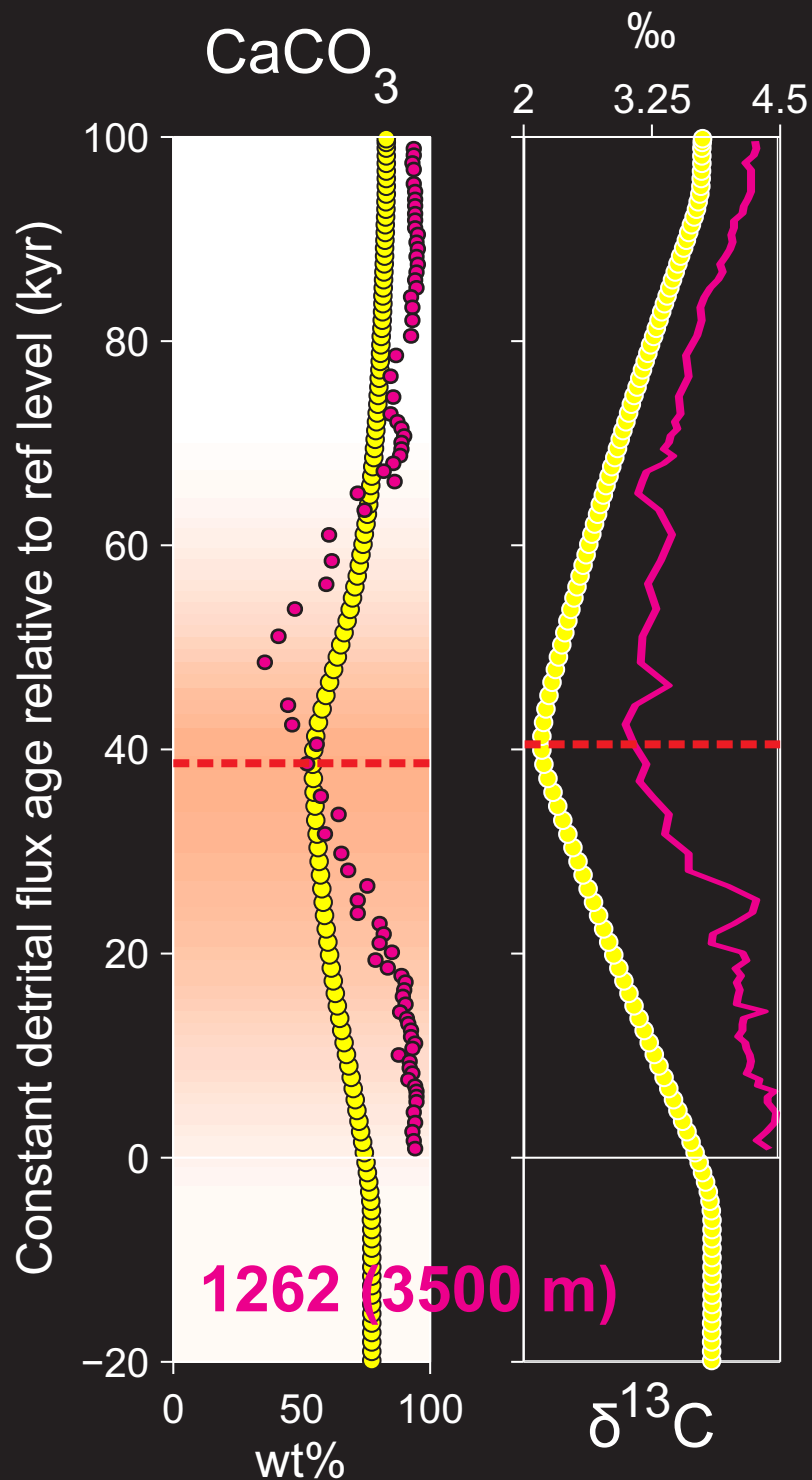








# 'Homogeneous' $\text{CaCO}_3$ dissolution



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