

The Global Weathering Thermostat in the Anthropocene

(numerical fun with silicate weathering and other climate control knobs)

Andy Ridgwell



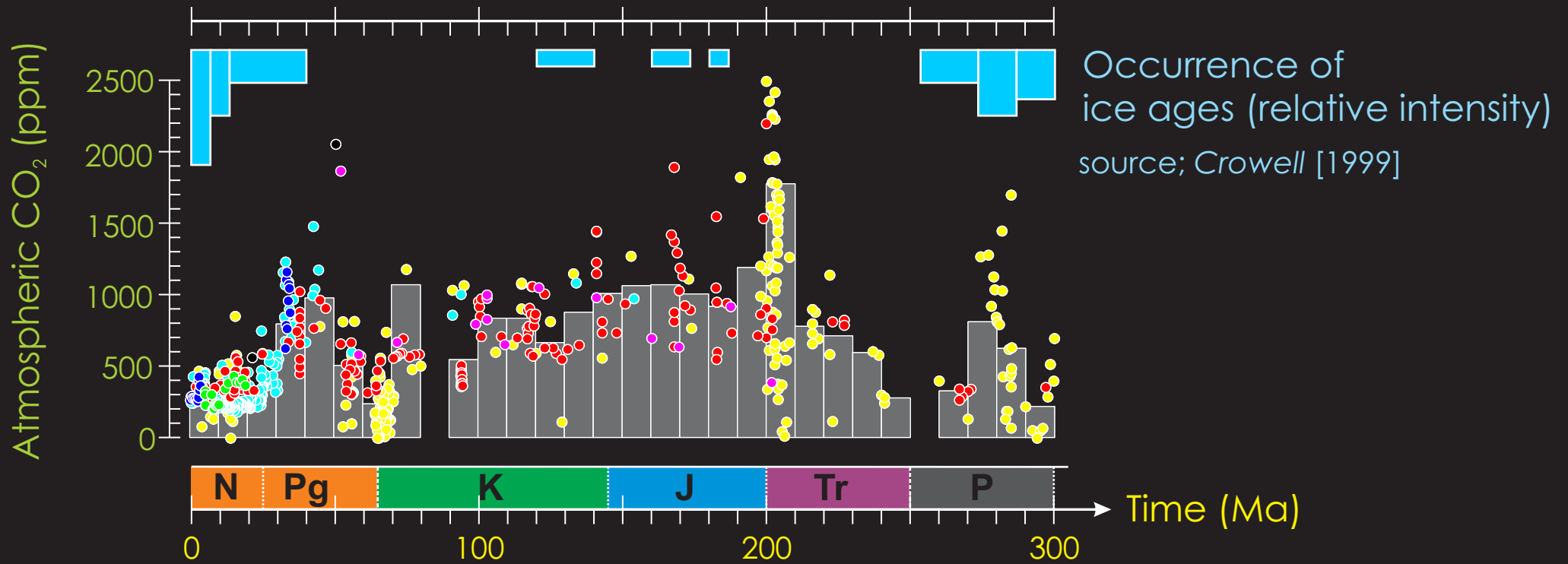
Regulation of global climate



Regulation of global climate



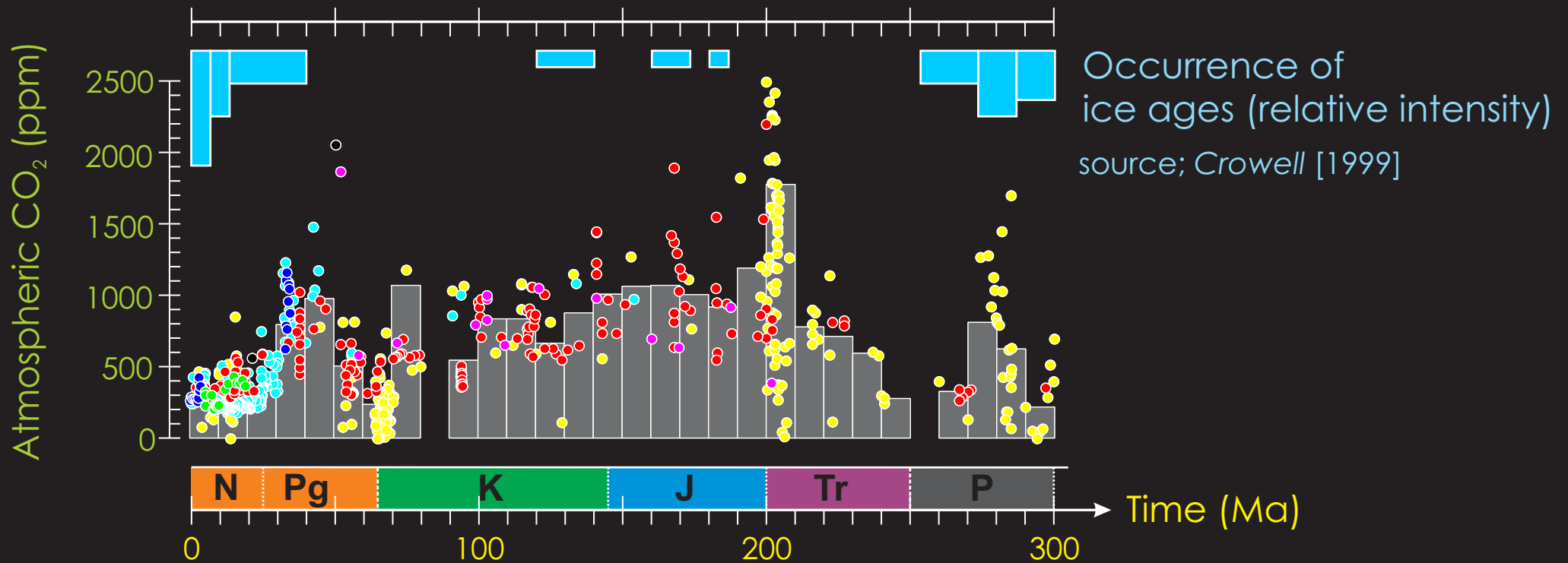
From: *Hönisch et al. [2012]*



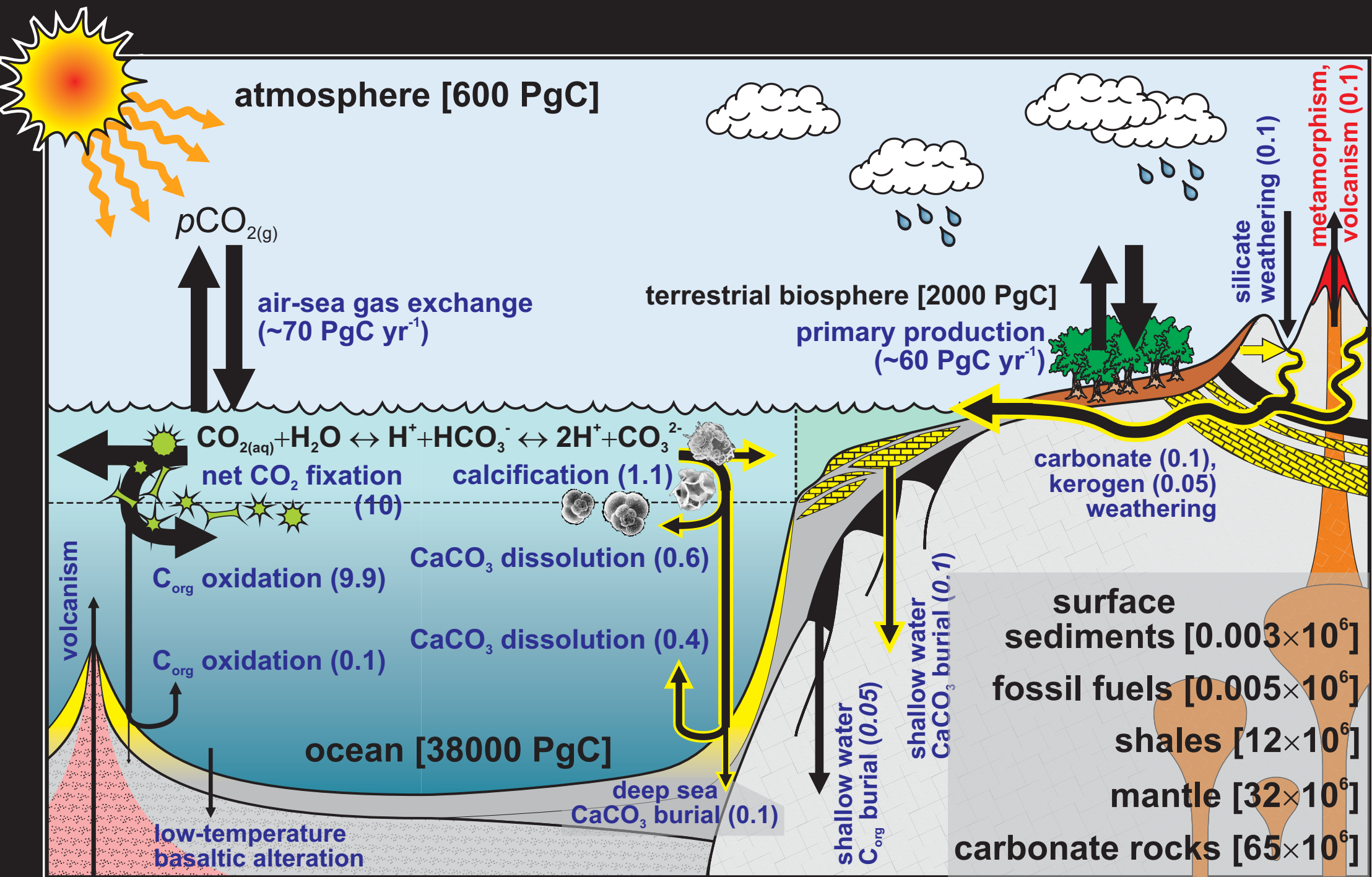
Regulation of global ~~climate~~ carbon cycling



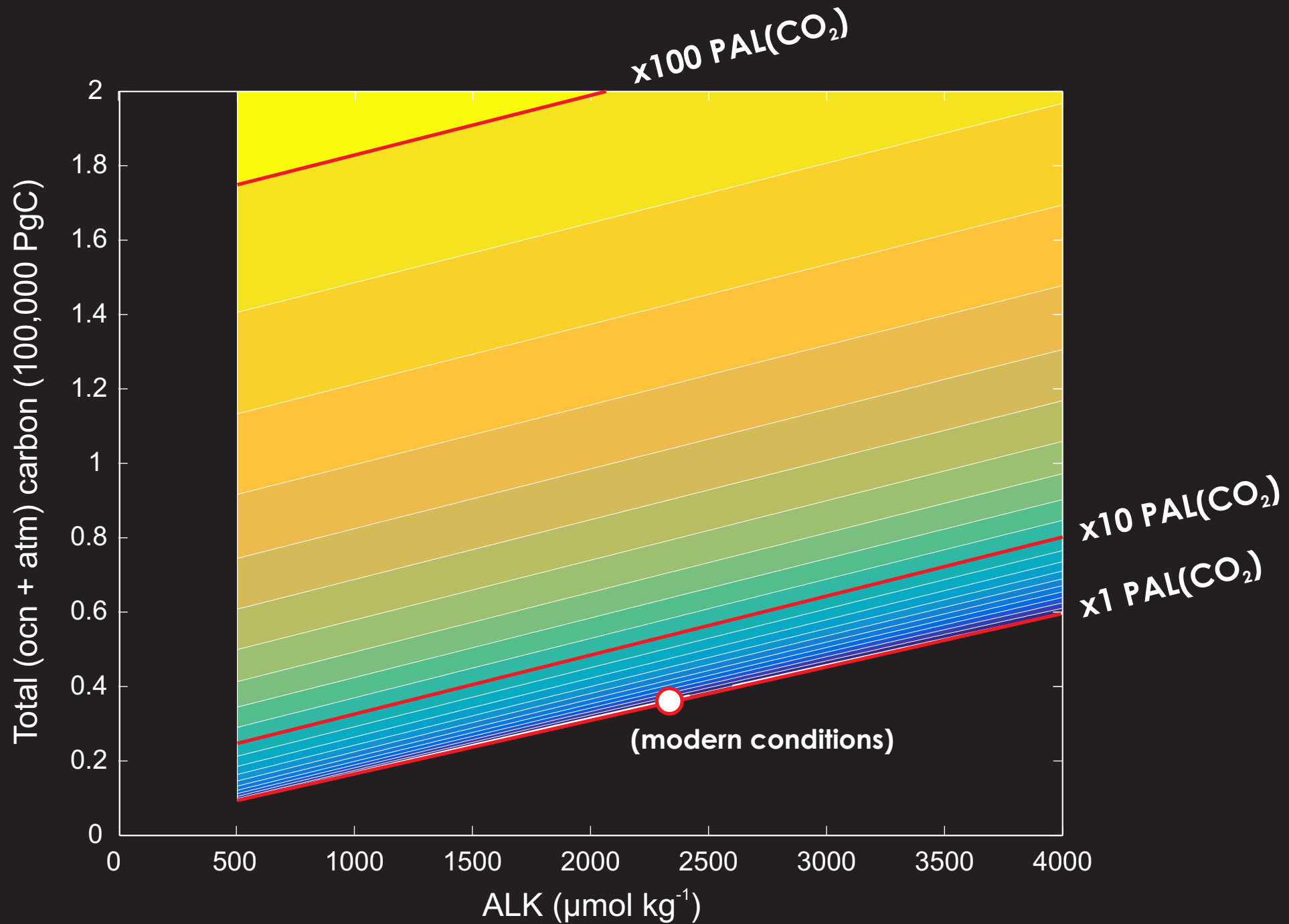
From: *Hönisch et al. [2012]*



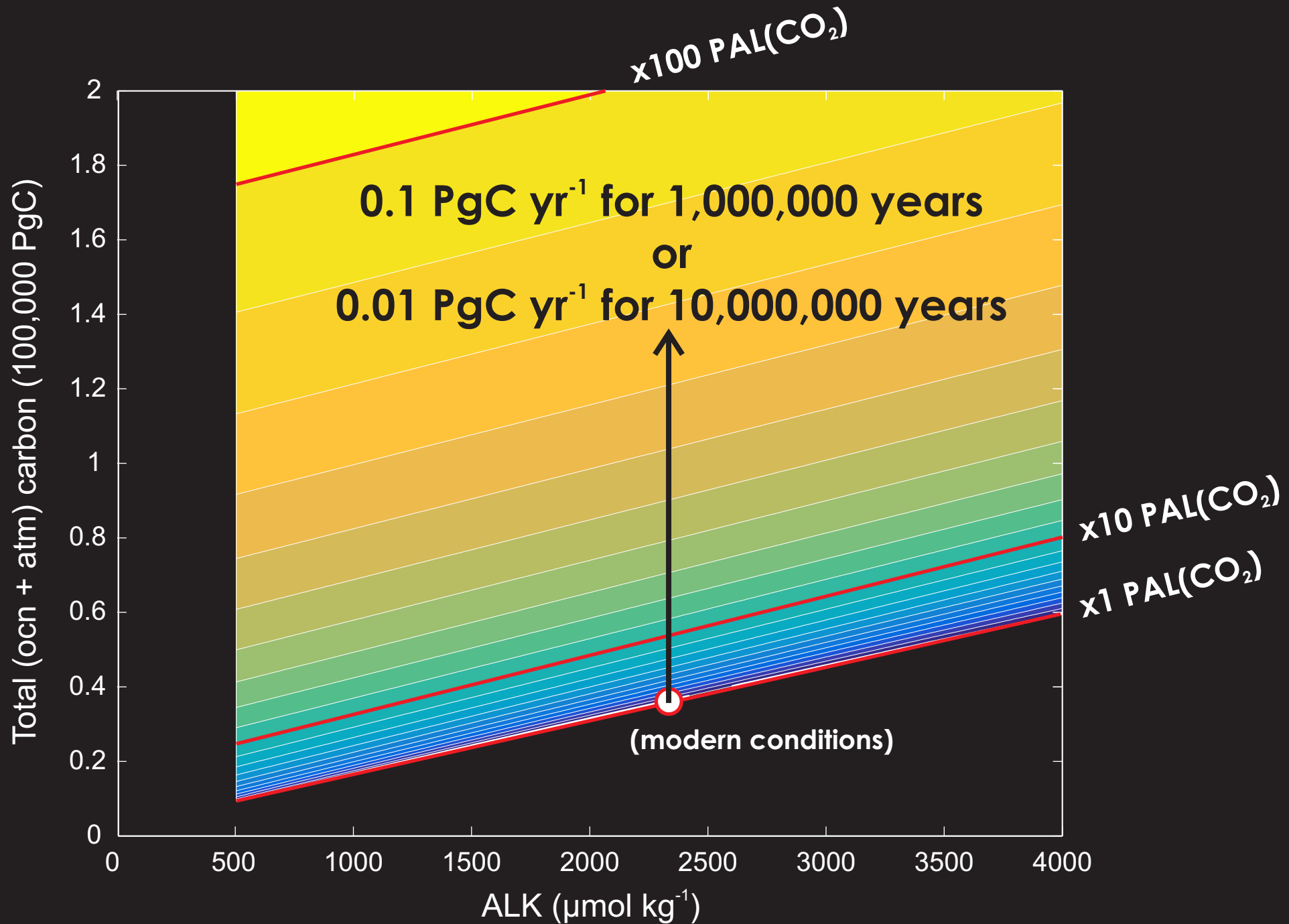
Regulation of global ~~climate~~ carbon cycling



Regulation of global ~~climate~~ carbon cycling



Regulation of global ~~climate~~ carbon cycling



Regulation of global ~~climate~~ carbon cycling





Terrestrial weathering can be (approximately equally) divided into carbonate (CaCO_3) and calcium-silicate (' CaSiO_3 ') weathering:



Ultimately, the (alkalinity: Ca^{2+}) weathering products must be removed through carbonate precipitation and burial in marine sediments:



It can be seen that in (2) + (3), that the CO_2 removed (from the atmosphere) during weathering, is returned upon carbonate precipitation (and burial). In (1) + (3) (silicate weathering) CO_2 is permanently removed to the geological reservoir. This CO_2 must be balanced by mantle (/volcanic) out-gassing on the very long term.

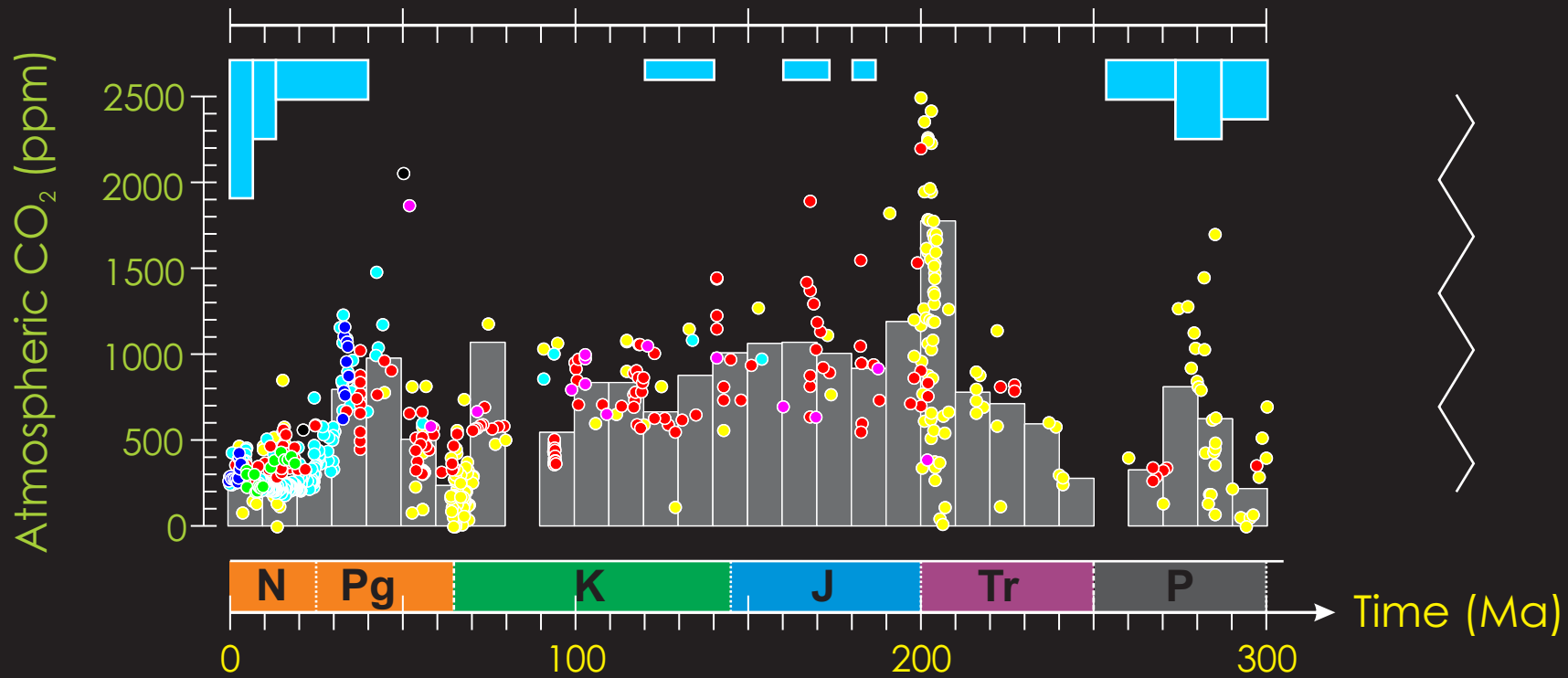
Regulation of global ~~climate~~ carbon cycling



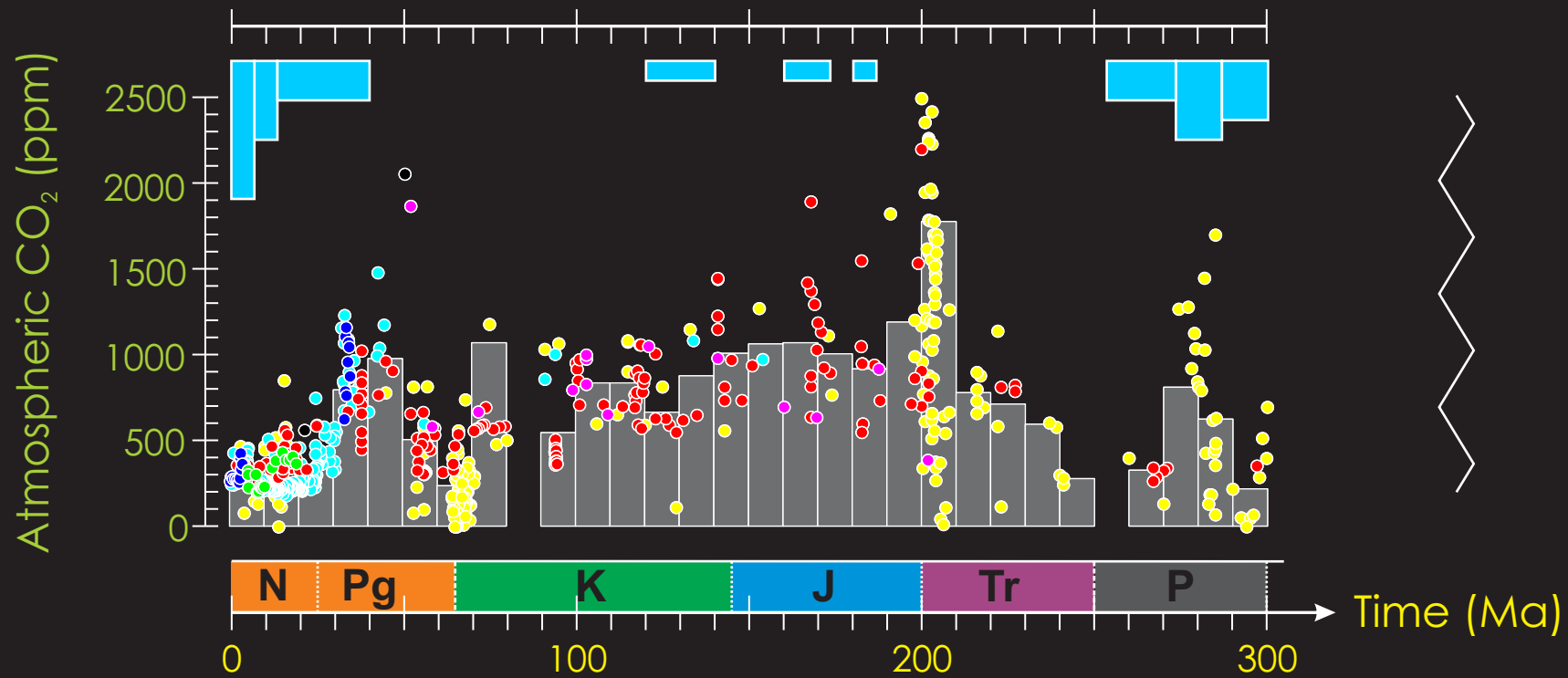
Furthermore, the rate of silicate weathering should scale with climate. Hence the **silicate weathering feedback** is formed:

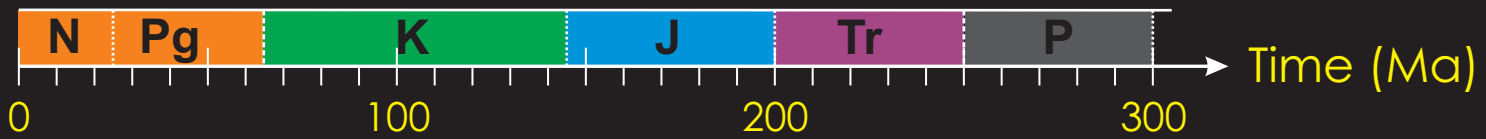
higher $p\text{CO}_2$ \rightarrow higher temperatures (& rainfall) \rightarrow higher weathering rates \rightarrow lower $p\text{CO}_2$

Regulation of global ~~climate~~ carbon cycling



Regulation of global ~~climate~~ carbon cycling





lies, damn lies, and computer models



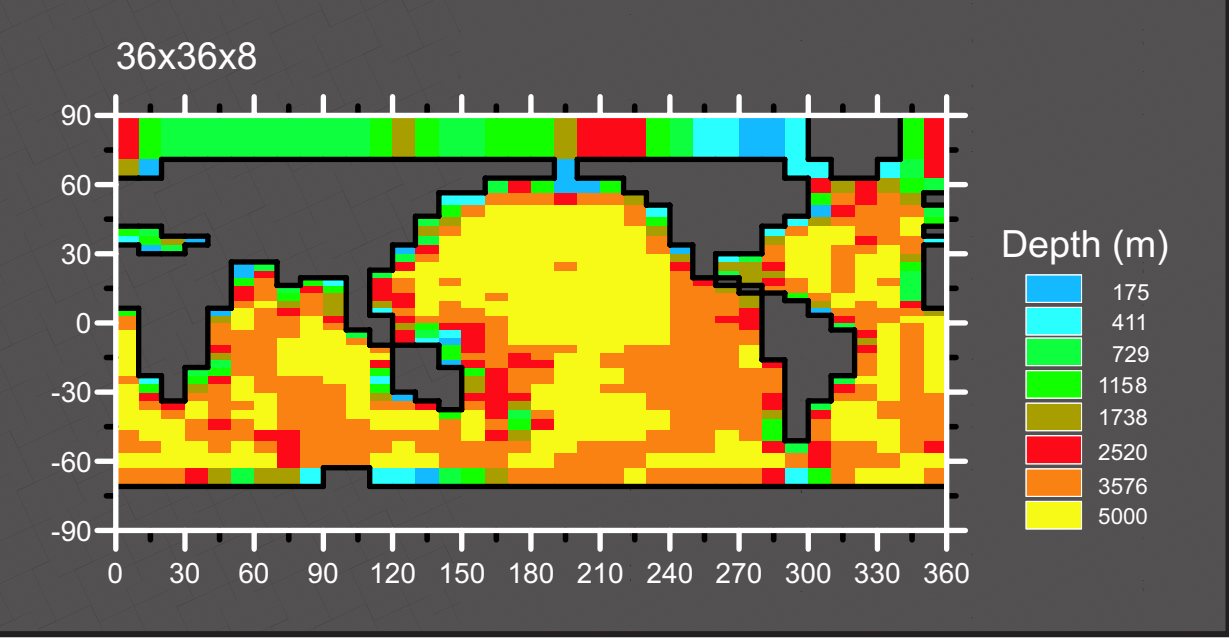
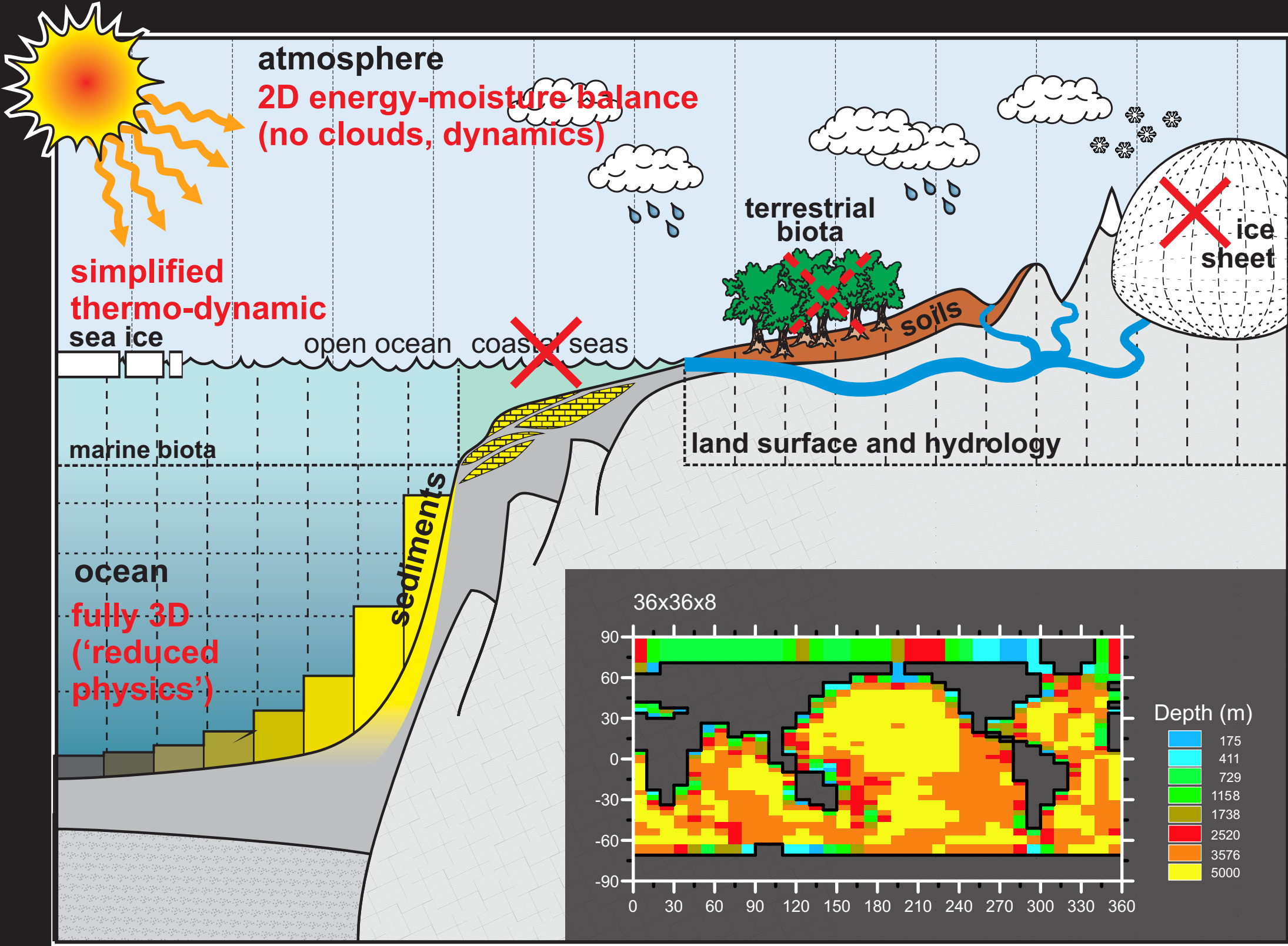
```
! calculate carbonate alkalinity
loc_ALK_DIC = dum_ALK &
& - loc_H4BO4 - loc_OH - loc_HPO4 - 2.0*loc_PO4 - loc_H3SiO4 - loc_NH3 - loc_HS &
& + loc_H + loc_HSO4 + loc_HF + loc_H3PO4

! estimate the partitioning between the aqueous carbonate species
loc_zed = ( &
& (4.0*loc_ALK_DIC + dum_DIC*dum_carbconst(icc_k) -
loc_ALK_DIC*dum_carbconst(icc_k))**2 + &
& 4.0*(dum_carbconst(icc_k) - 4.0)*loc_ALK_DIC**2 &
& )**0.5      loc_conc_HCO3 = (dum_DIC*dum_carbconst(icc_k) -
loc_zed)/(dum_carbconst(icc_k) - 4.0)

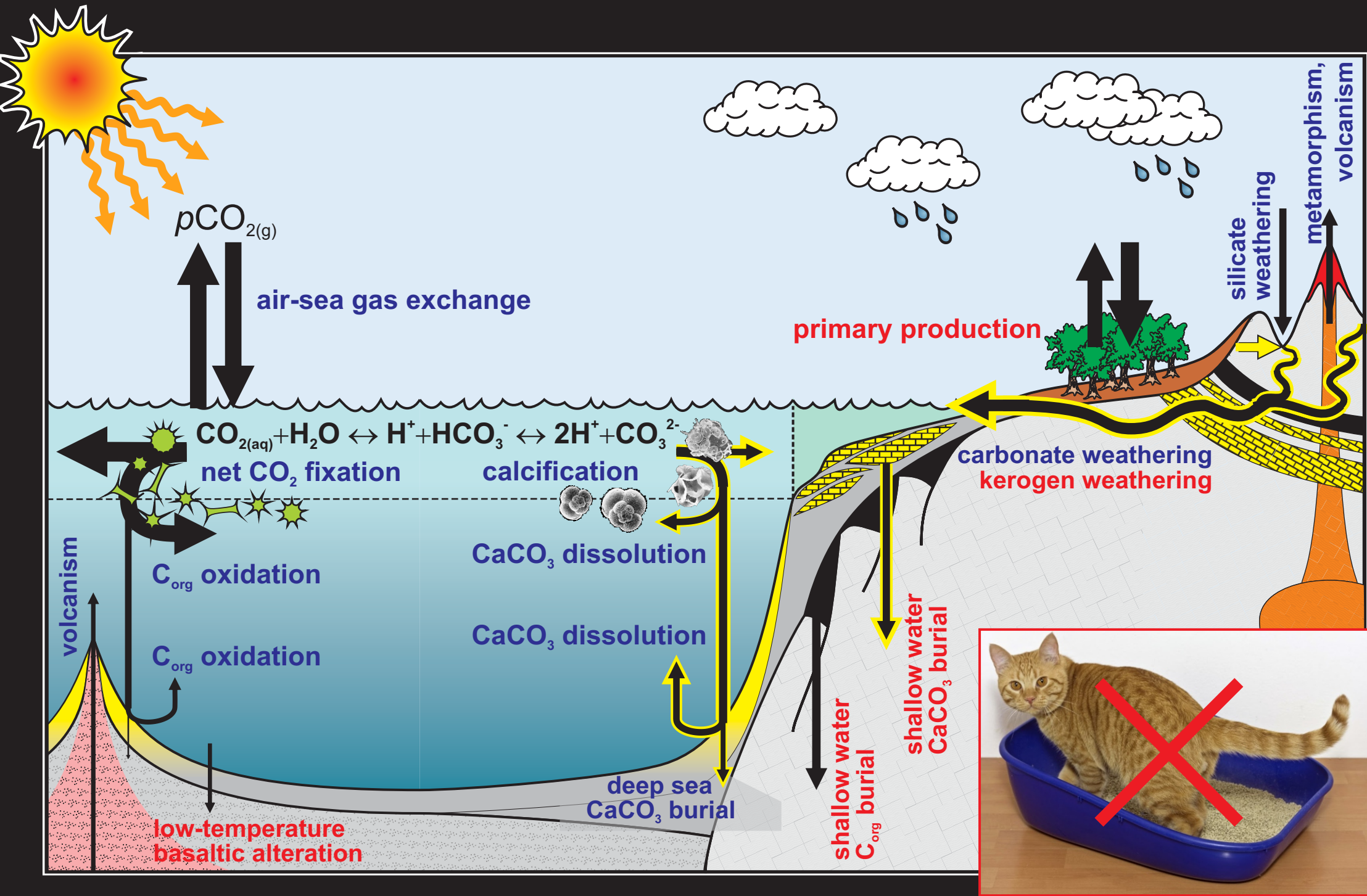
loc_conc_CO3 = &
& ( &
& loc_ALK_DIC*dum_carbconst(icc_k) - dum_DIC*dum_carbconst(icc_k) - &
& 4.0*loc_ALK_DIC + loc_zed &
& ) &
& /(2.0*(dum_carbconst(icc_k) - 4.0))

loc_conc_CO2 = dum_DIC - loc_ALK_DIC + &
& ( &
& loc_ALK_DIC*dum_carbconst(icc_k) - dum_DIC*dum_carbconst(icc_k) - &
& 4.0*loc_ALK_DIC + loc_zed &
& ) &
& /(2.0*(dum_carbconst(icc_k) - 4.0))

loc_H1 = dum_carbconst(icc_k1)*loc_conc_CO2/loc_conc_HCO3
loc_H2 = dum_carbconst(icc_k2)*loc_conc_HCO3/loc_conc_CO3
```

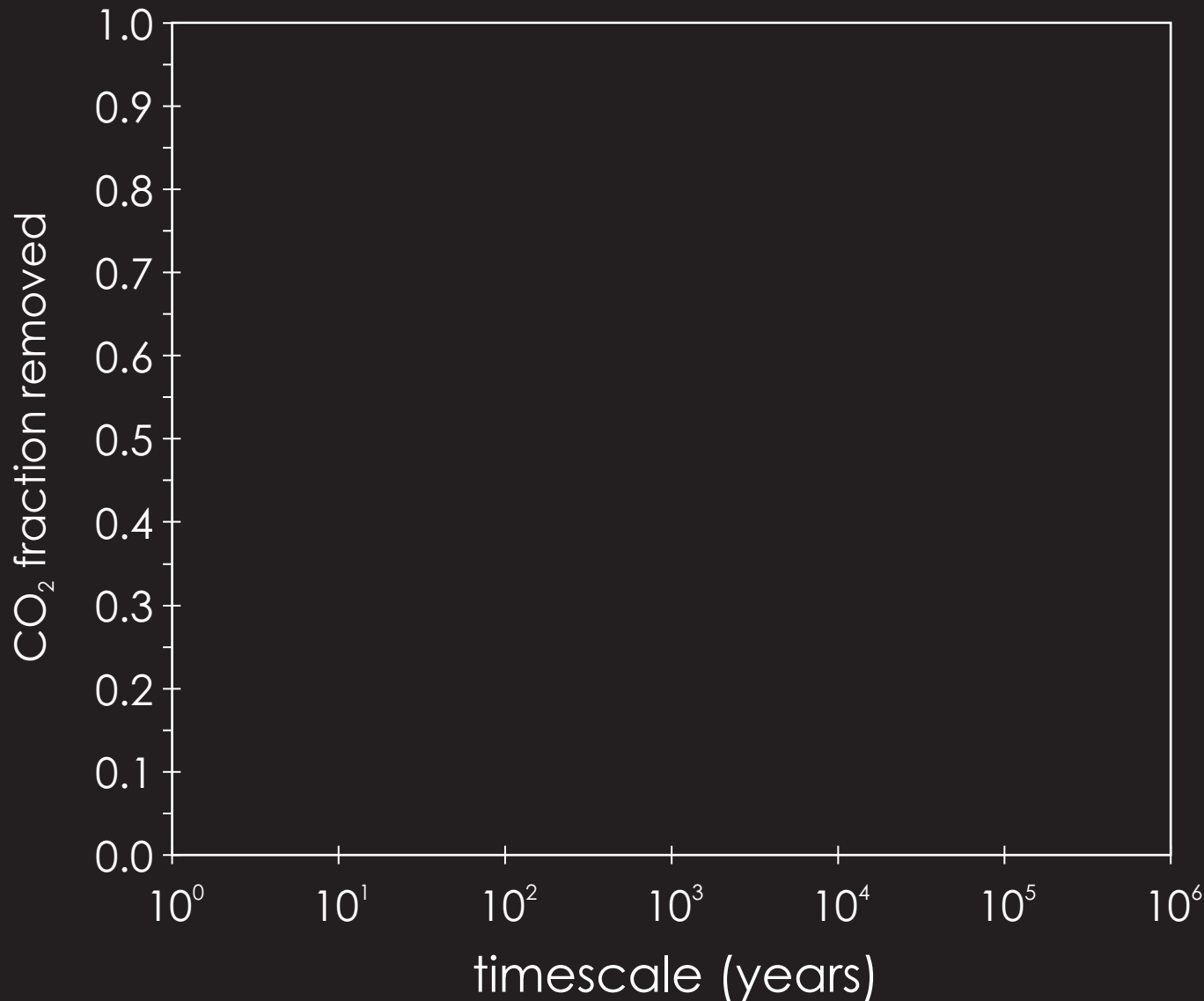


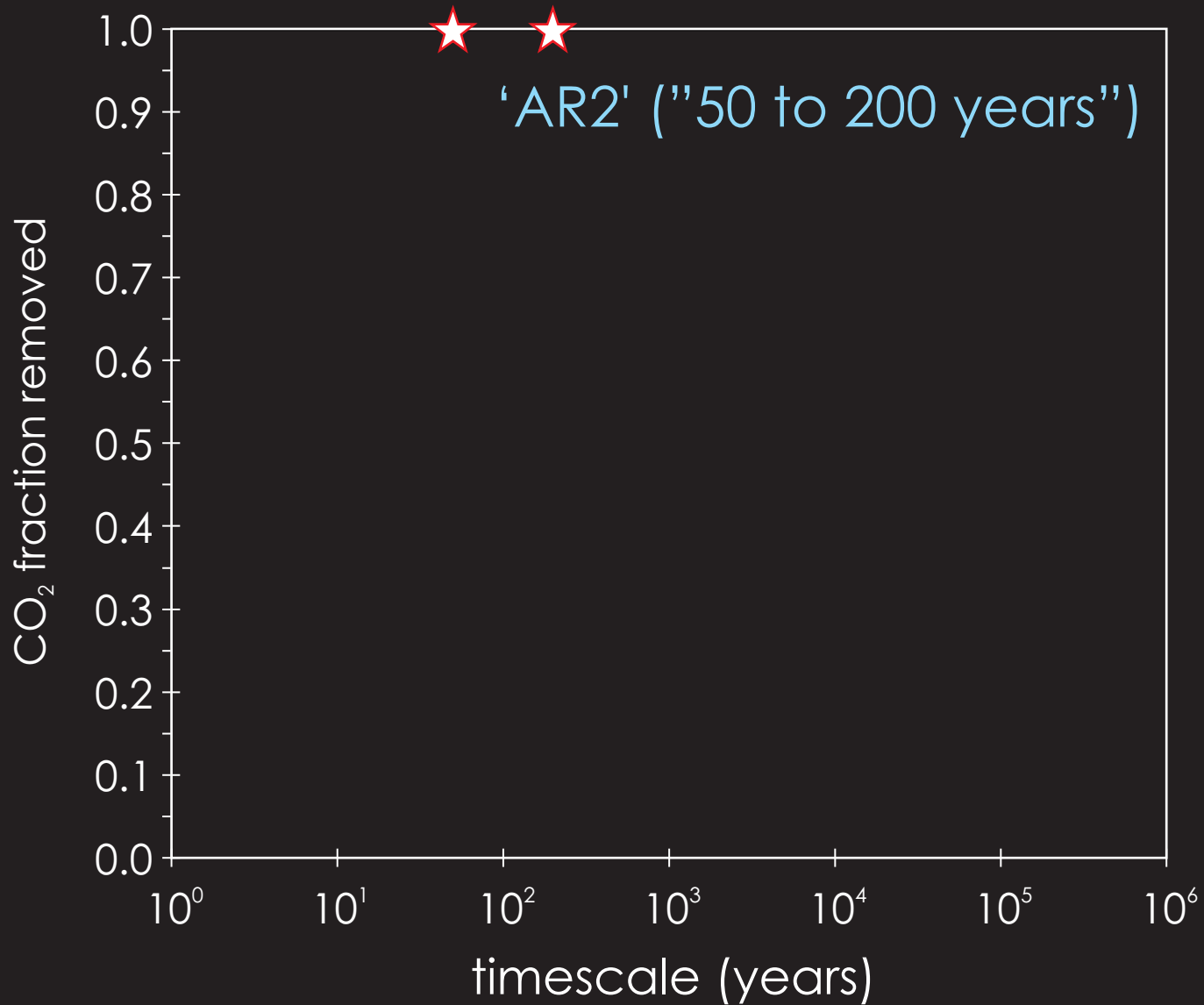
lies, damn lies, and computer models



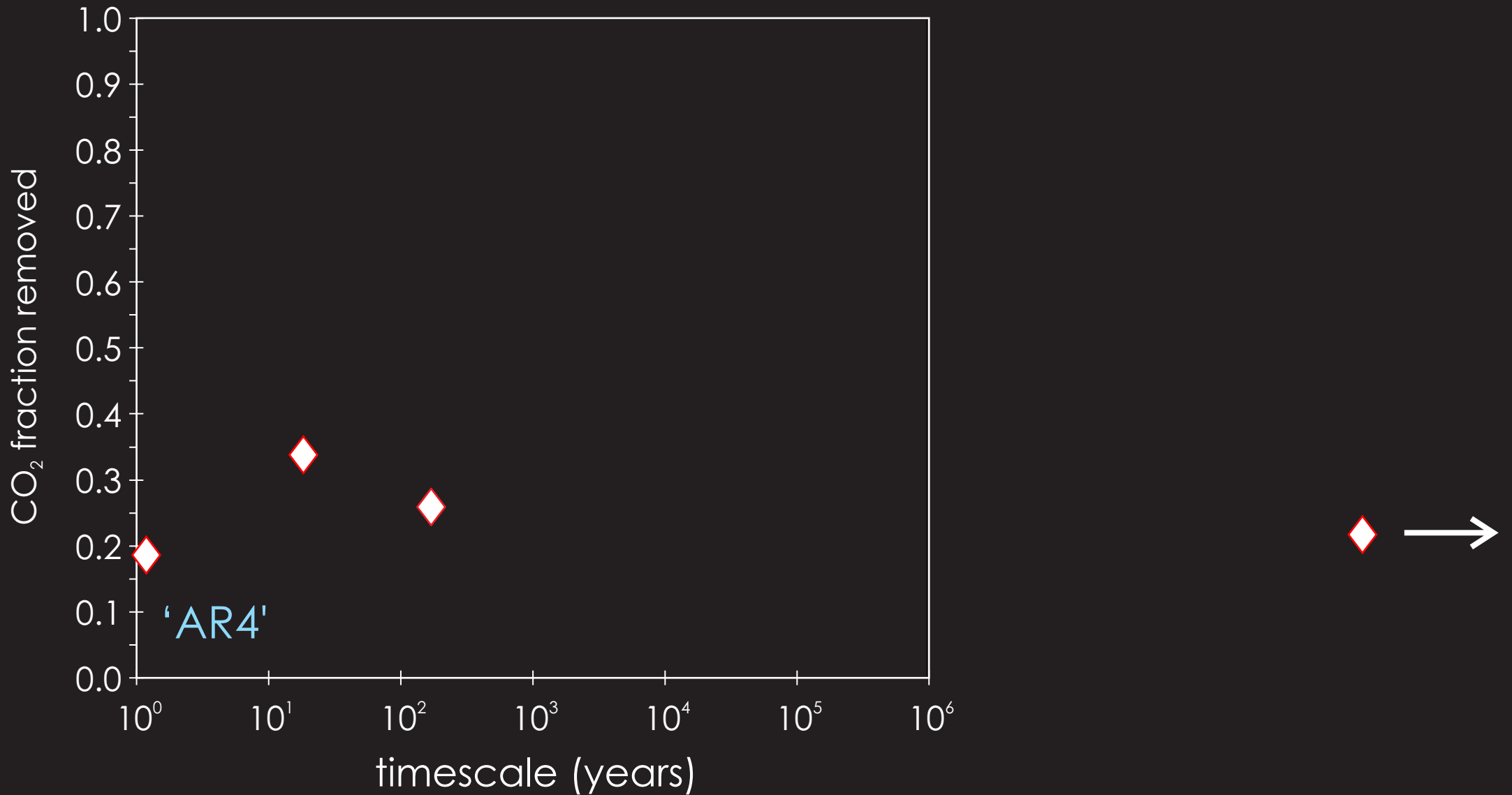


Cross-plot of the fraction of total CO_2 emissions to the atmosphere removed by a particular process (carbon sink), vs. the characteristic (e-folding) time-scale of that process (\log_{10} scale).





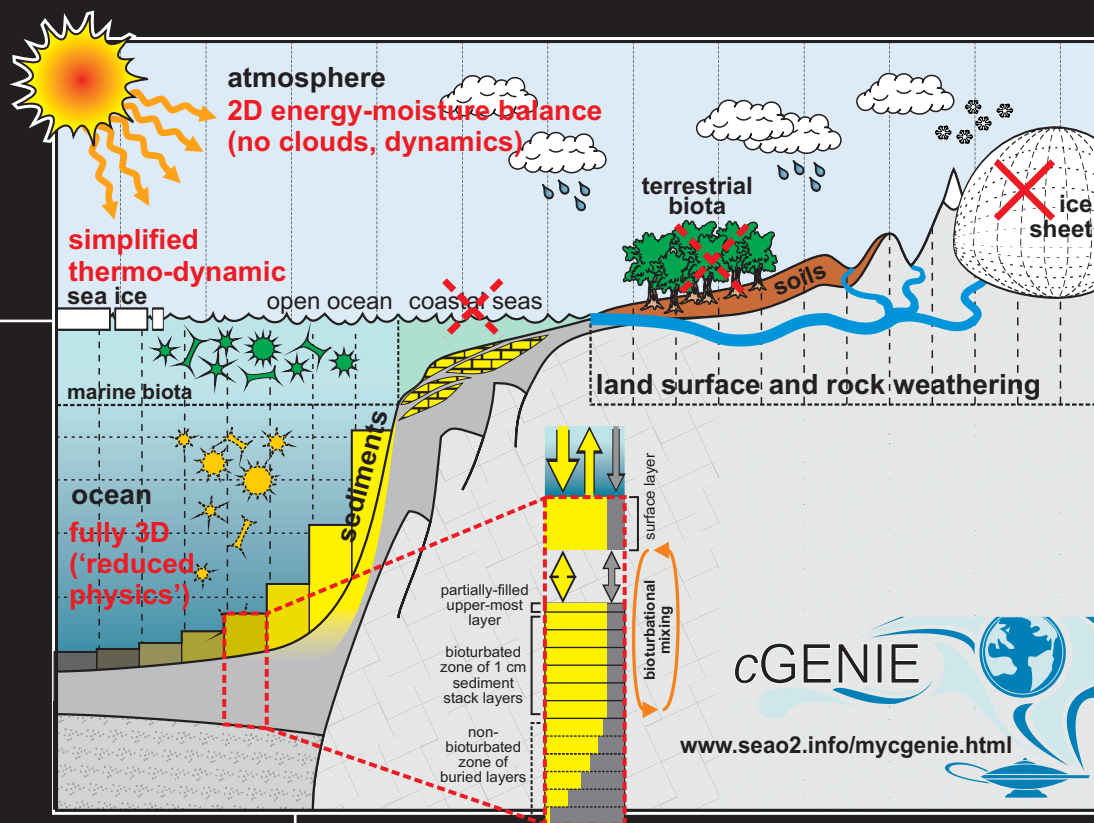
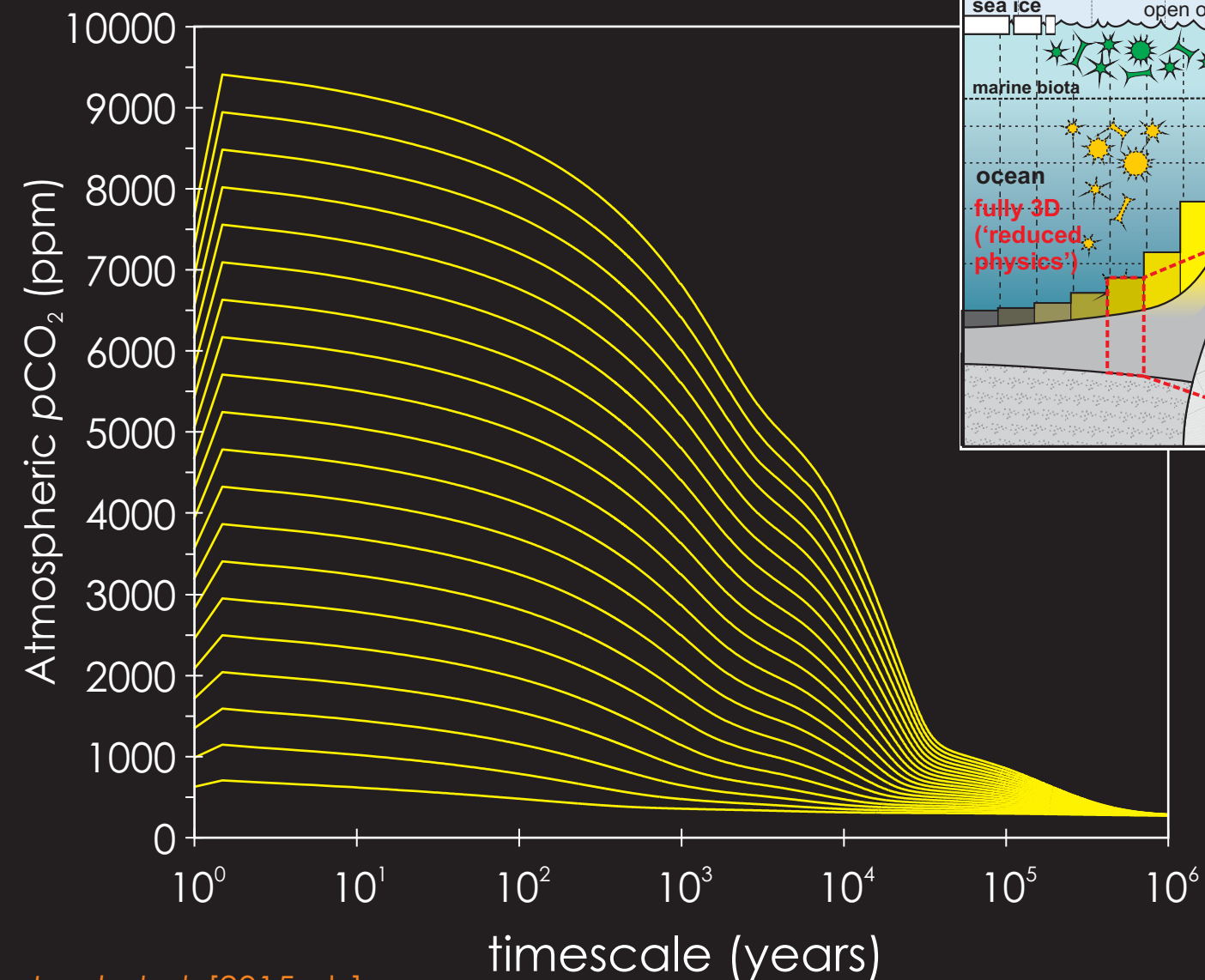
Impulse response function analysis of the 'long tail' of $\text{CO}_{2(\text{excess})}$



Impulse response function analysis of the 'long tail' of CO_{2(excess)}



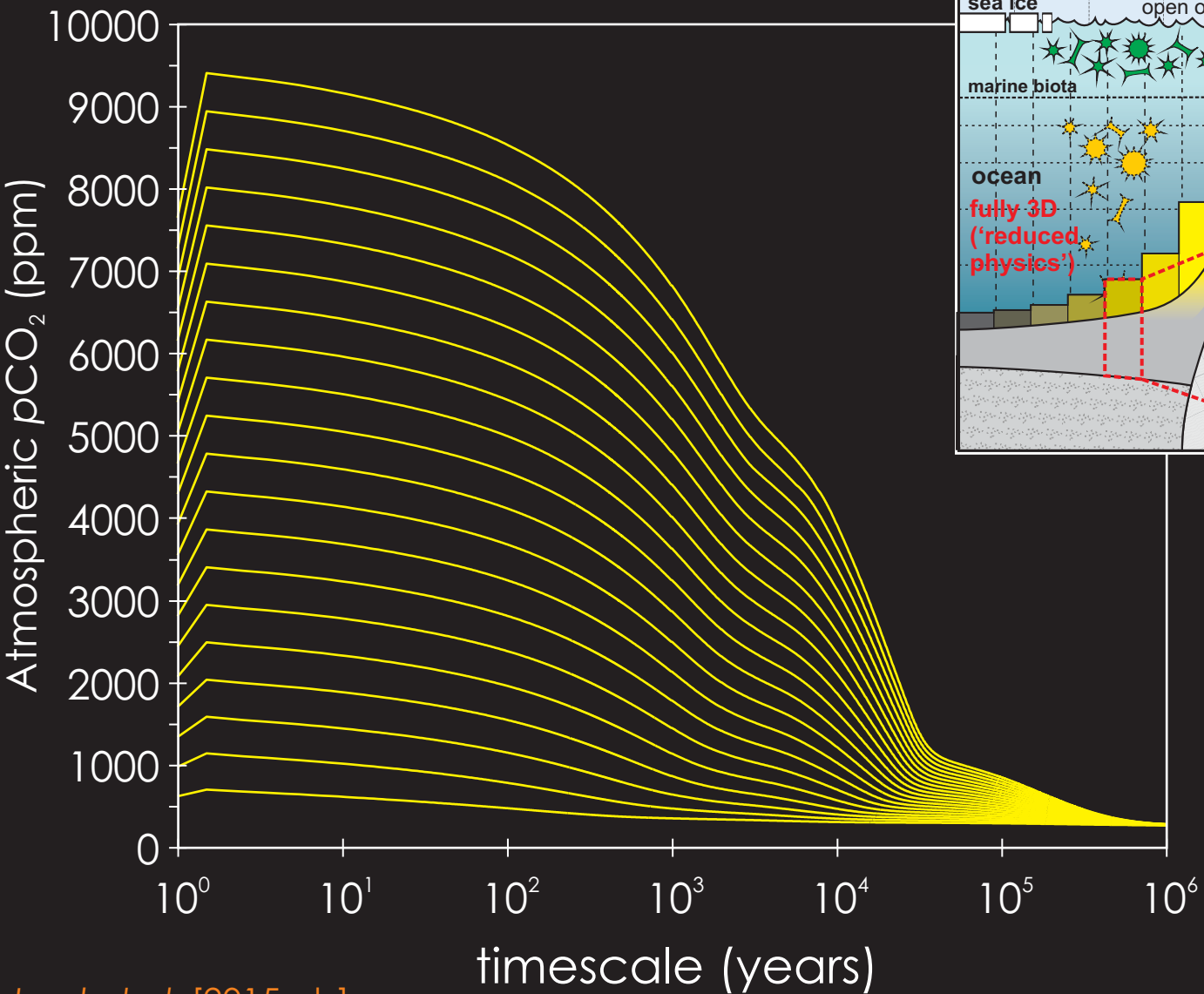
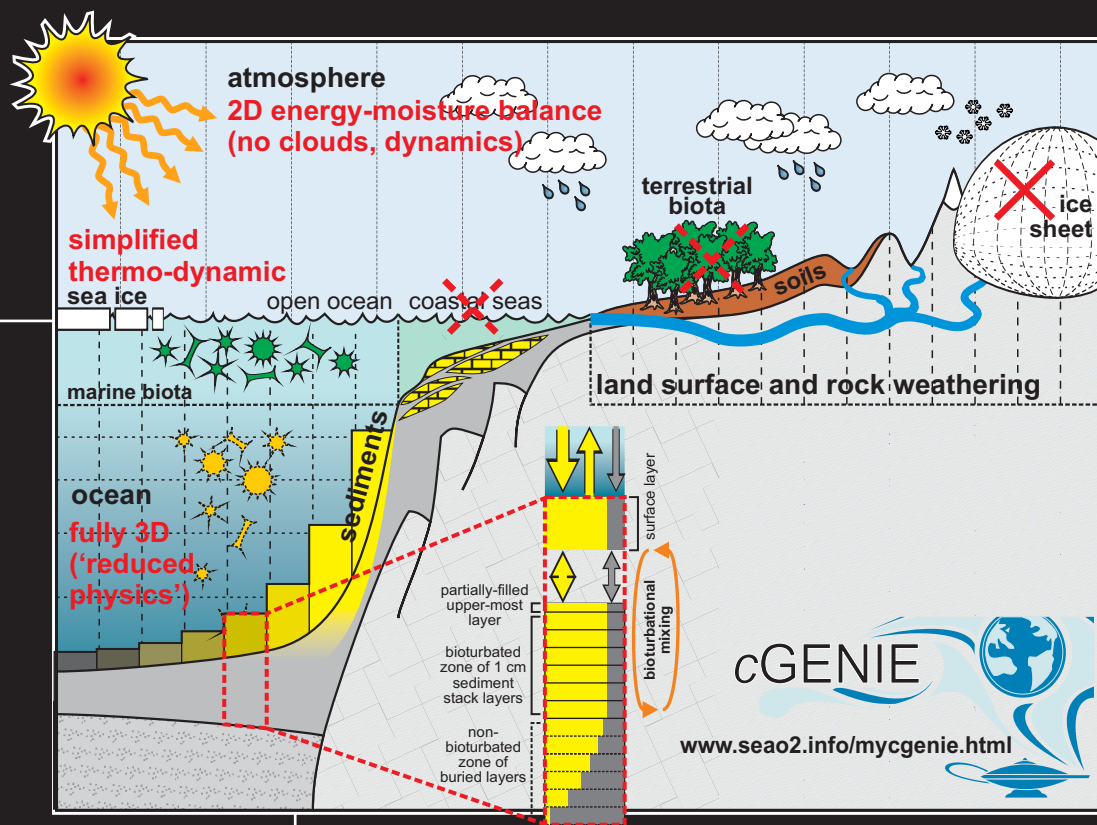
(1) Series of 1 Myr Earth system model experiments. CO₂ emissions from 1,000 to 20,000 PgC (GtC). Release interval: 1 yr.



Impulse response function analysis of the 'long tail' of CO_{2(excess)}



(1) Series of 1 Myr Earth system model experiments. CO₂ emissions from 1,000 to 20,000 PgC (GtC). Release interval: 1 yr.



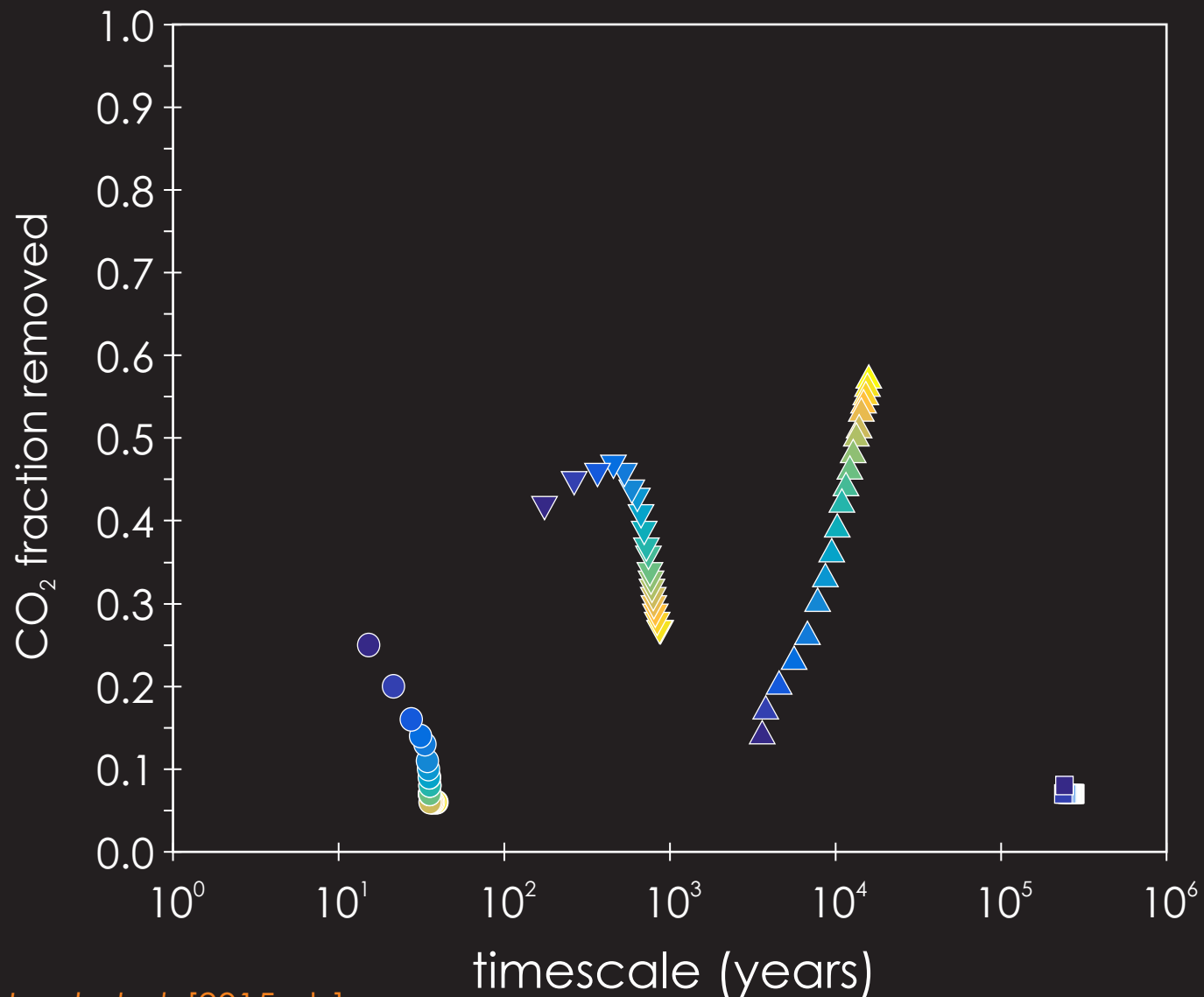
(2) Fit each CO₂ decay curve with a series (4 optimal) of exponentials. Extract the fraction of CO₂ and time-scale associated with each.

(The resulting empirical model can be used in place of a mechanistic model for projecting the long-term fate of carbon release.)

Lord et al. [2015a,b]

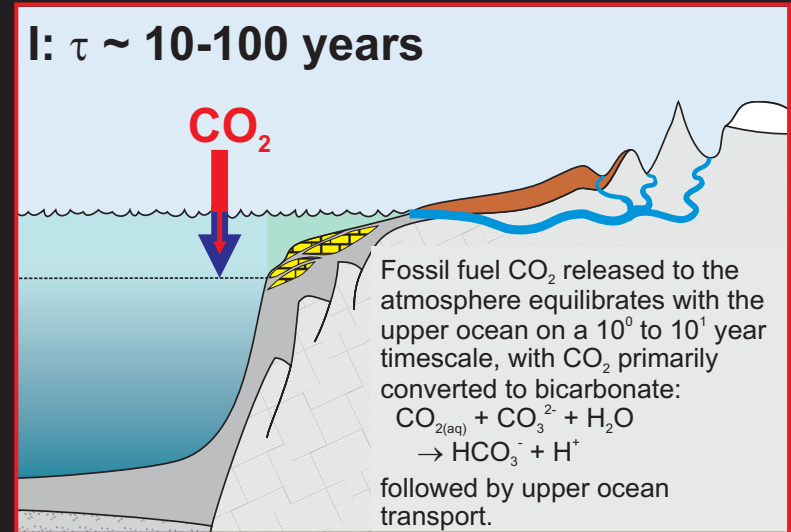
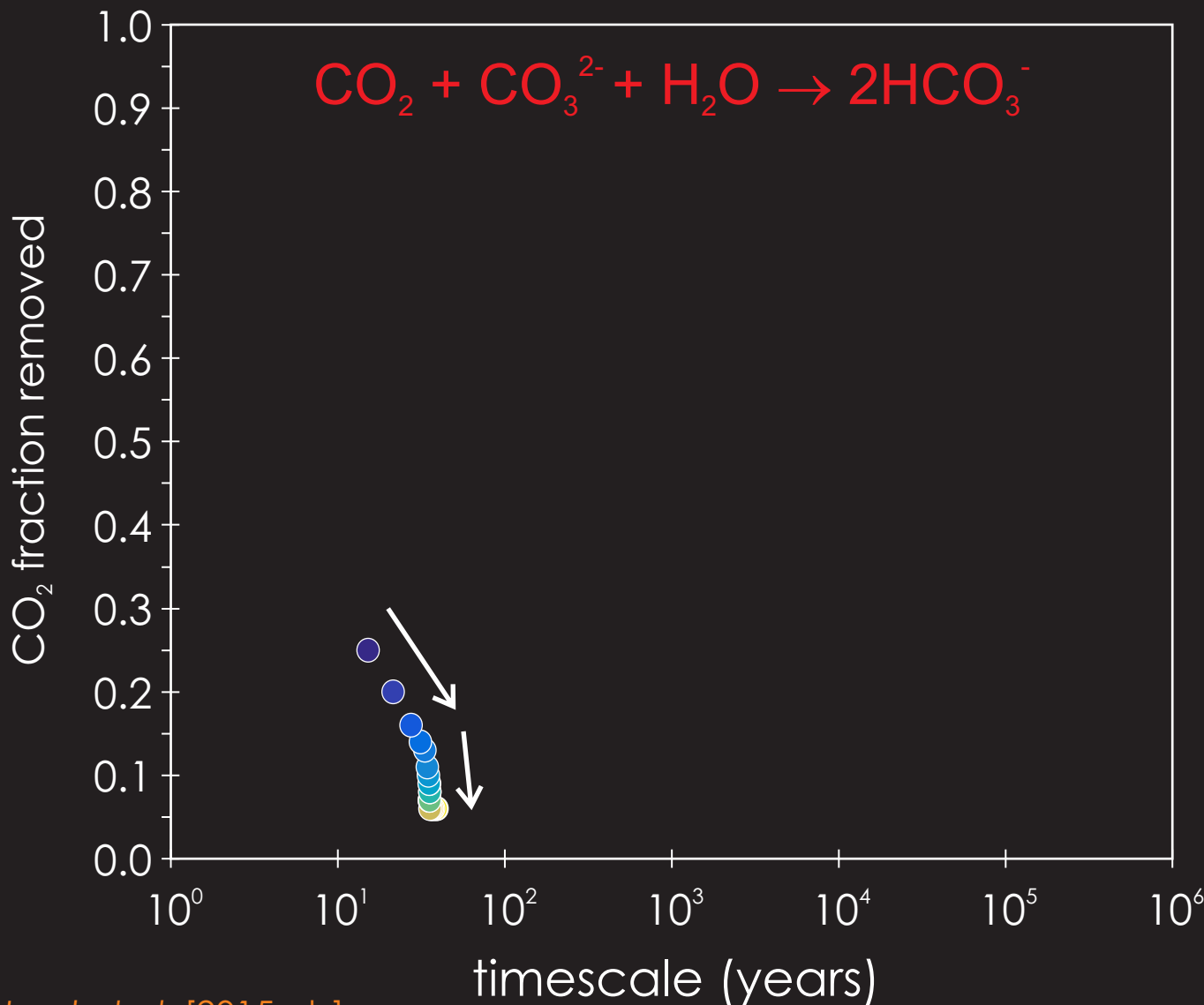


Response of fraction of CO_2 removed vs. the characteristic time-scale, as a function of total emissions, ranging from 1,000 PgC (dark blue) to 20,000 PgC (yellow).



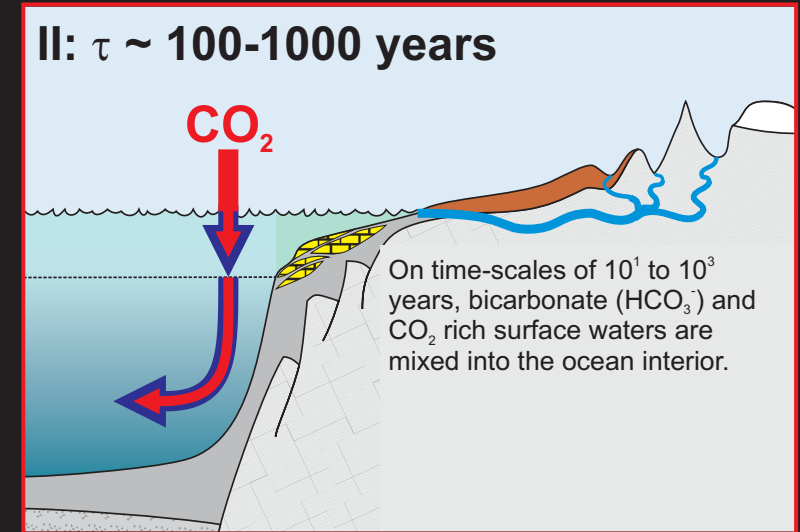
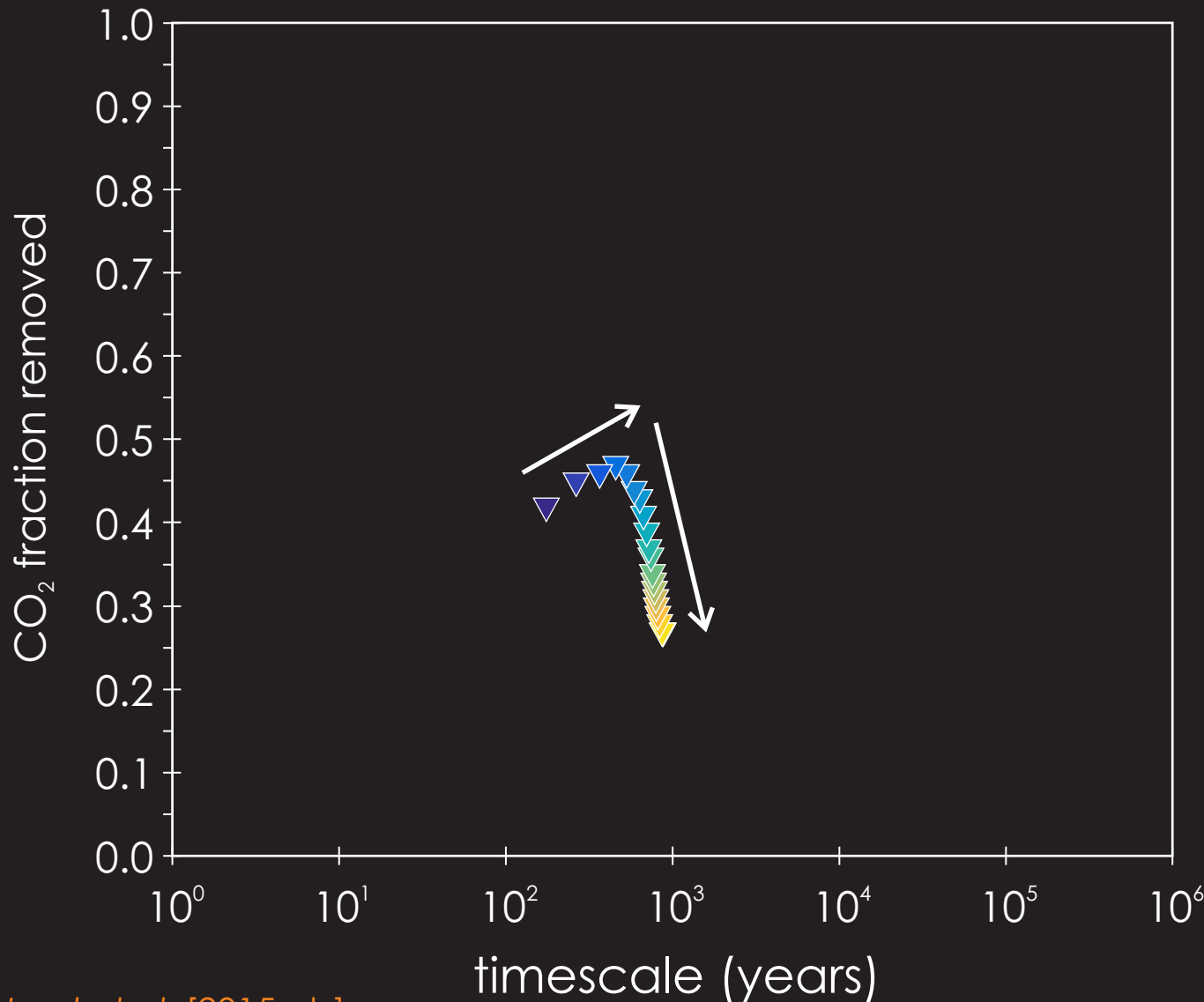


Depletion of mixed layer carbonate buffer;
ocean stratification and reduced surface
mixing. Warming and reduced CO_2 solubility.



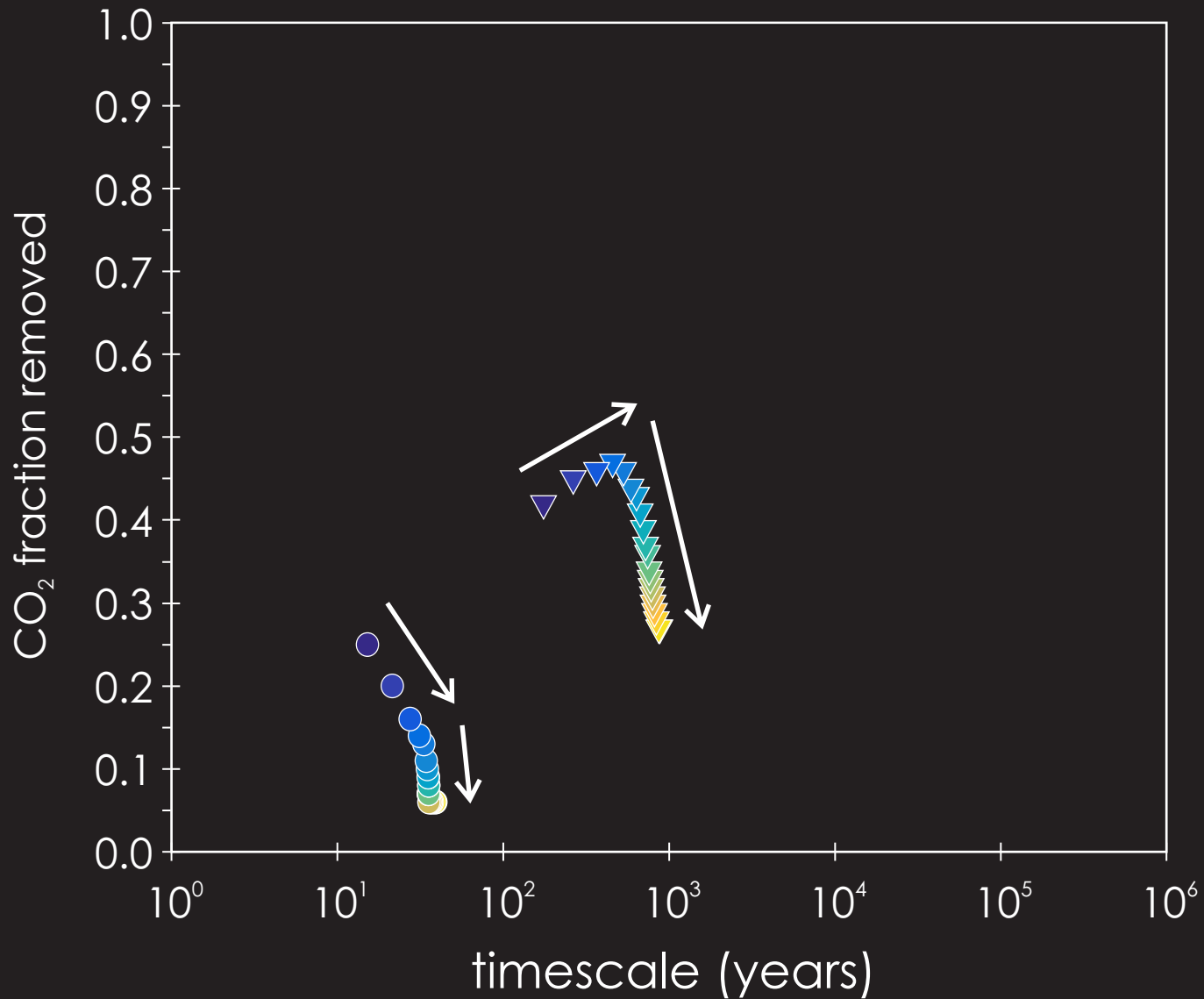


Ocean stratification and collapse of the AMOC
(in this particular model).
Threshold reached @ $\sim 4000 \text{ PgC}$?

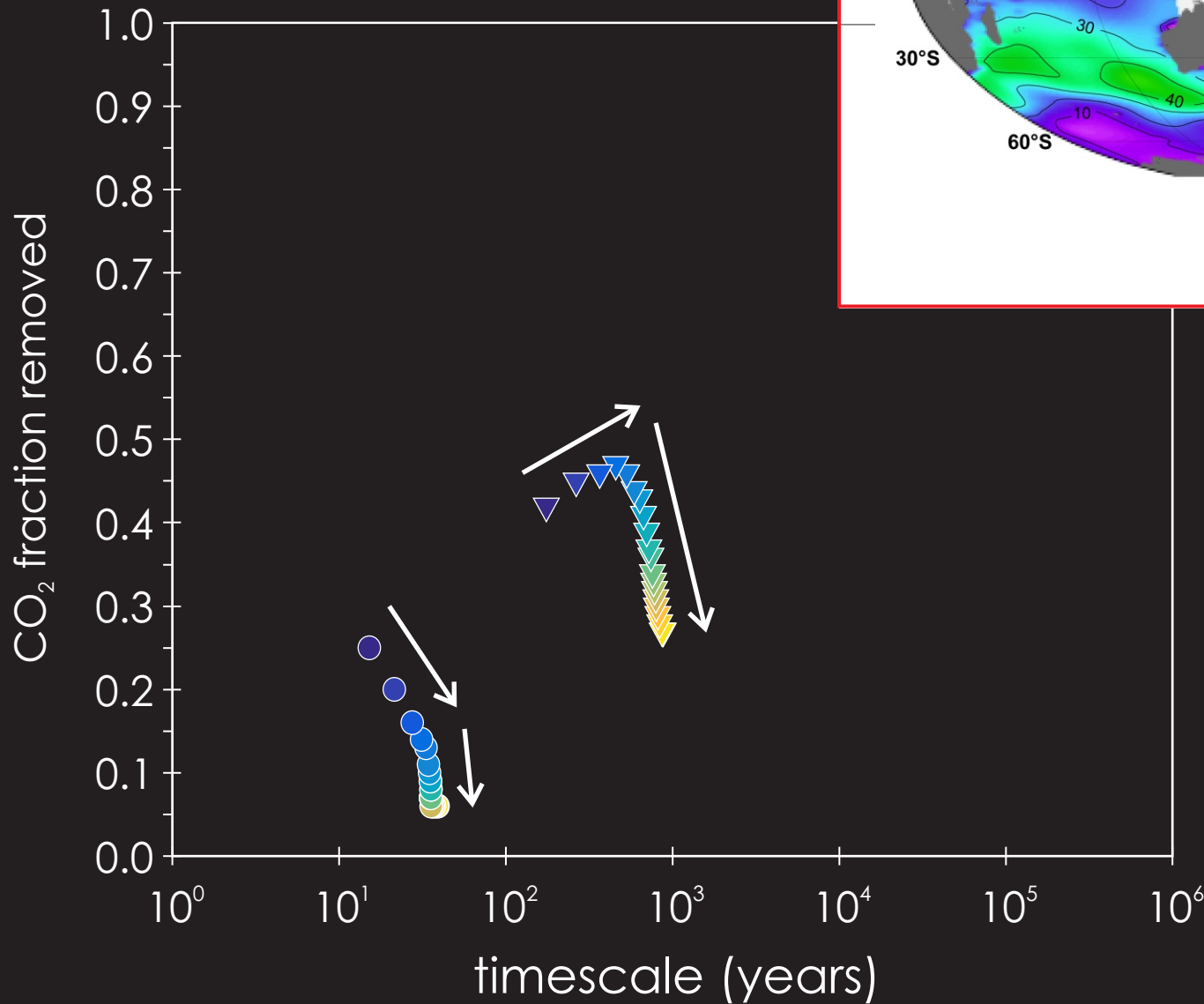
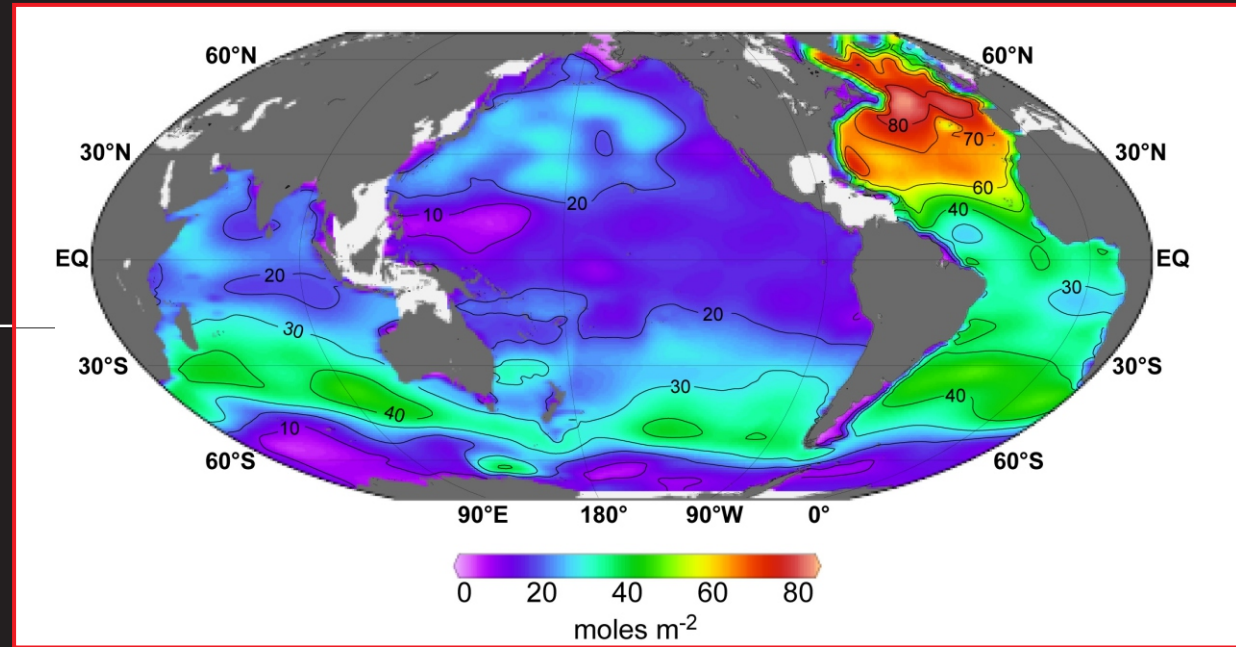




evidence?

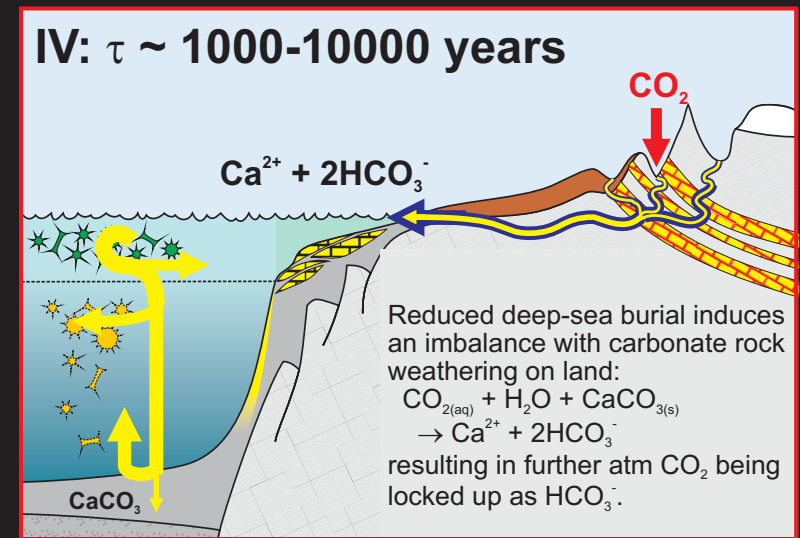
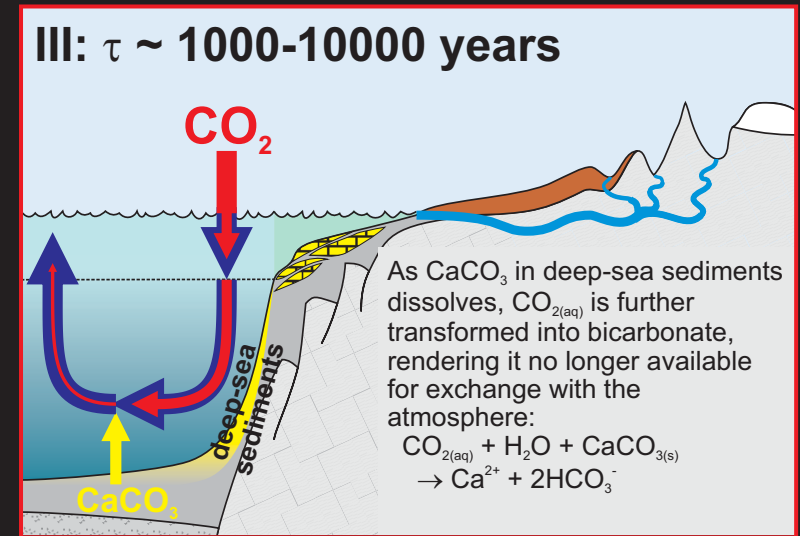
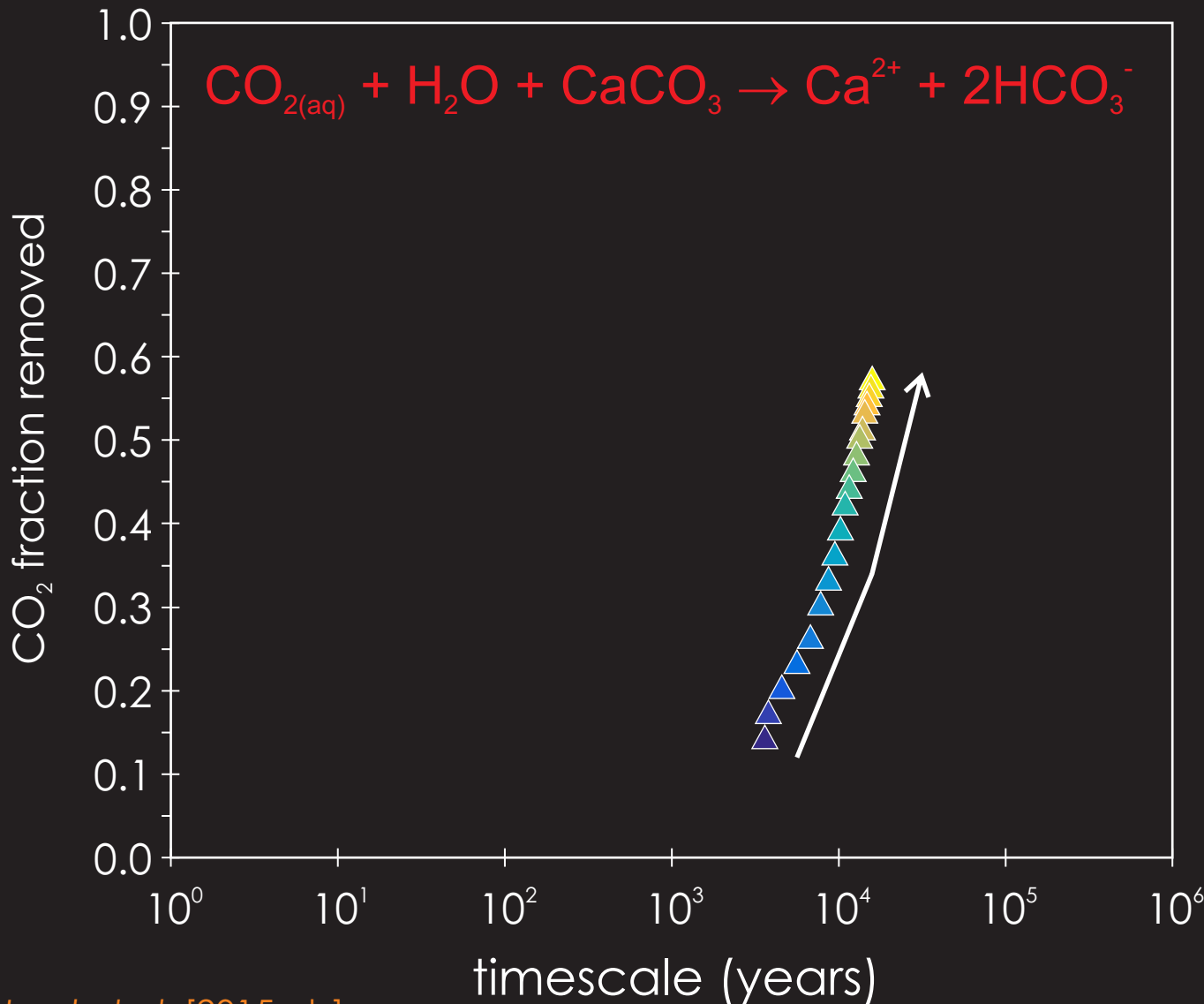


Impulse response function analysis of the 'long tail' of $\text{CO}_2(\text{excess})$



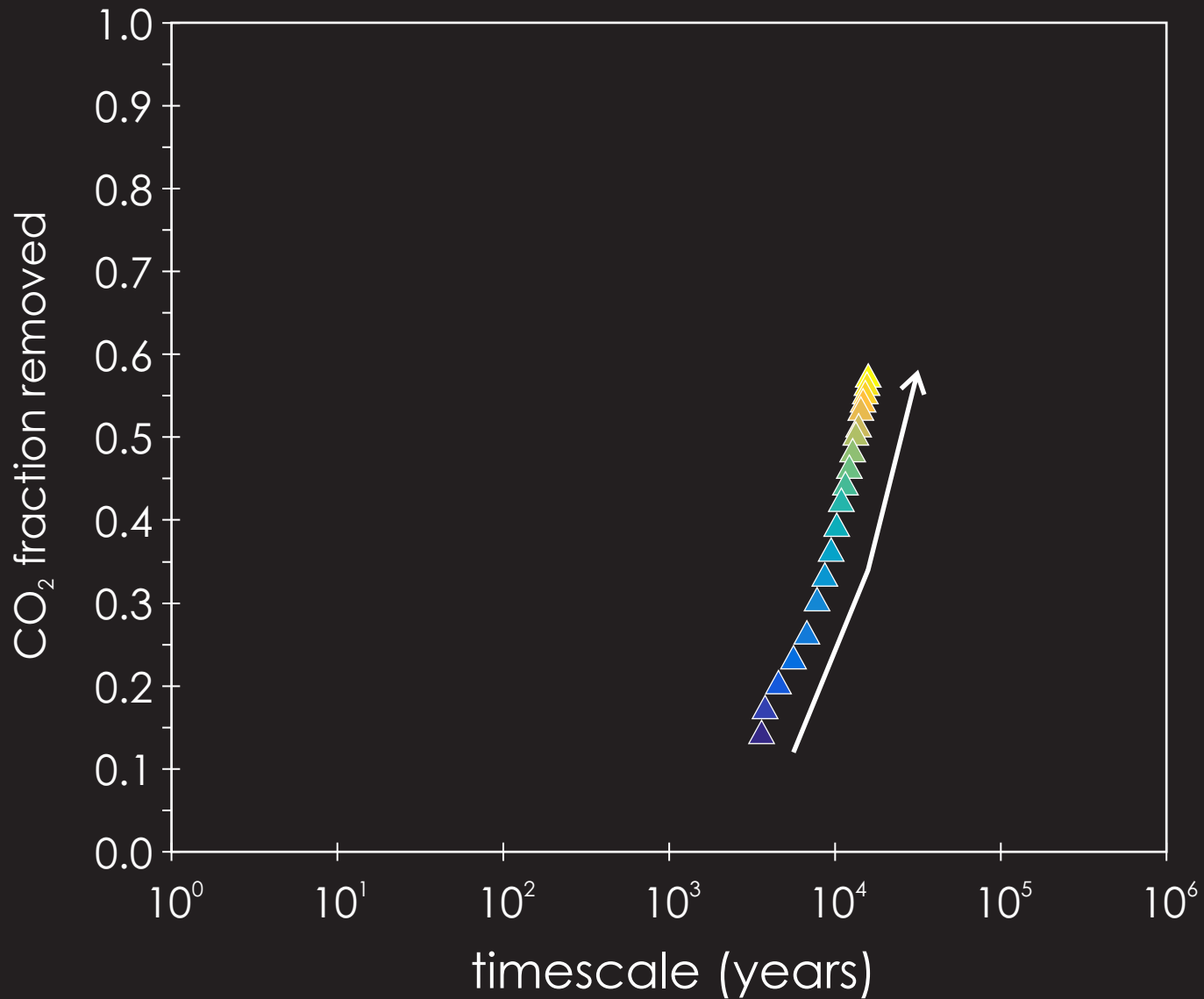


Geologic CO_2 removal via carbonate rocks and marine sediments – occurring on an increasing protracted time-scale.





evidence?

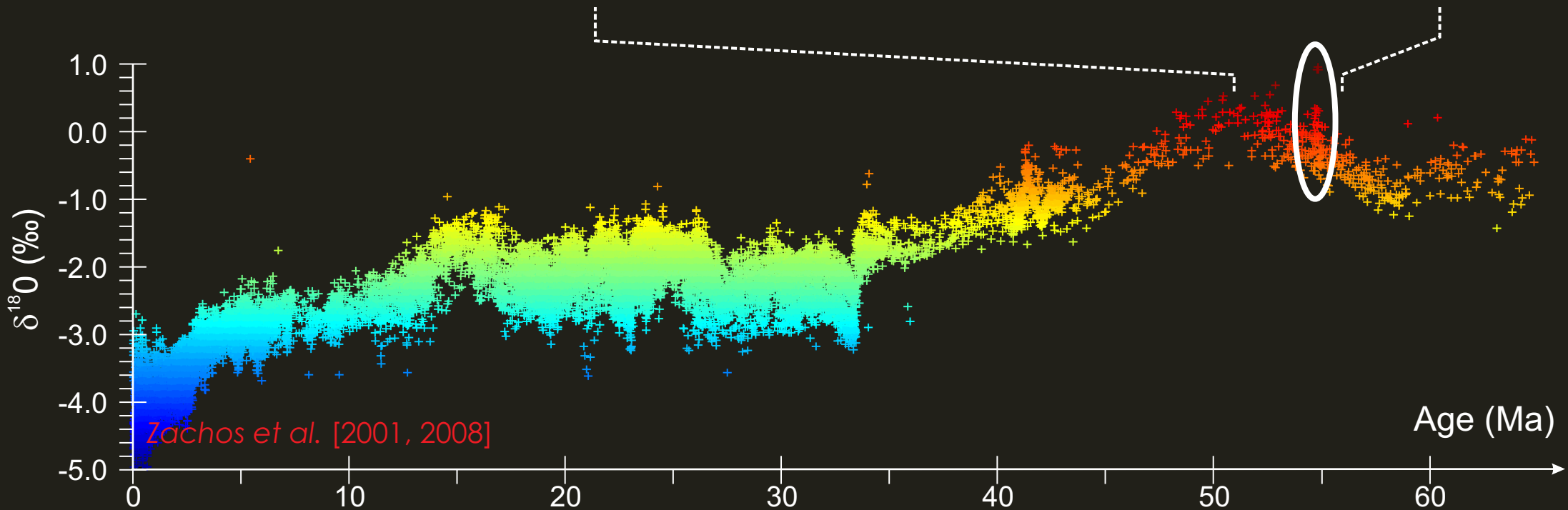
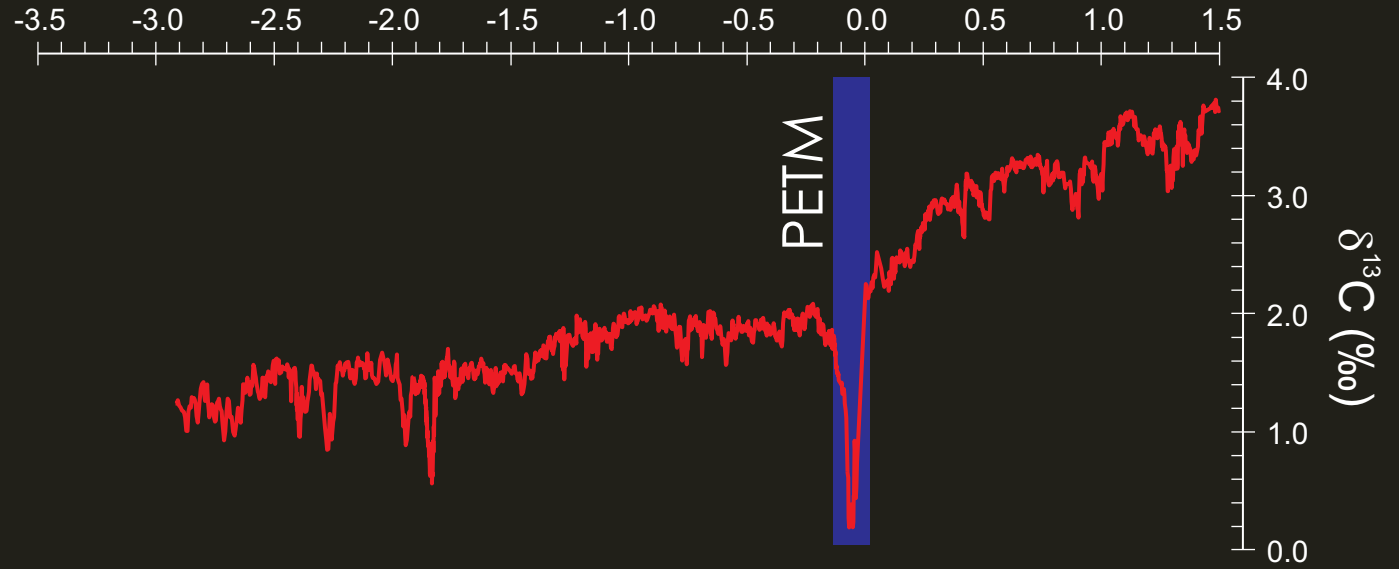


Impulse response function analysis of the 'long tail' of $\text{CO}_{2(\text{excess})}$

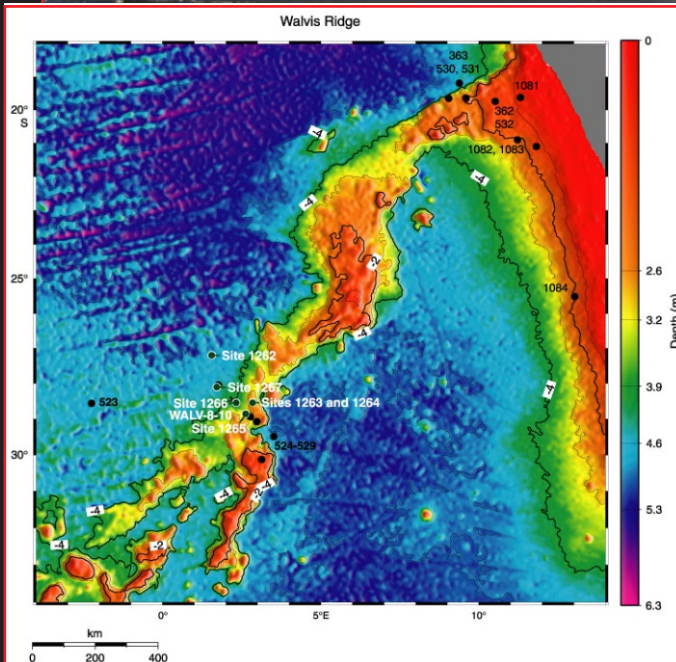
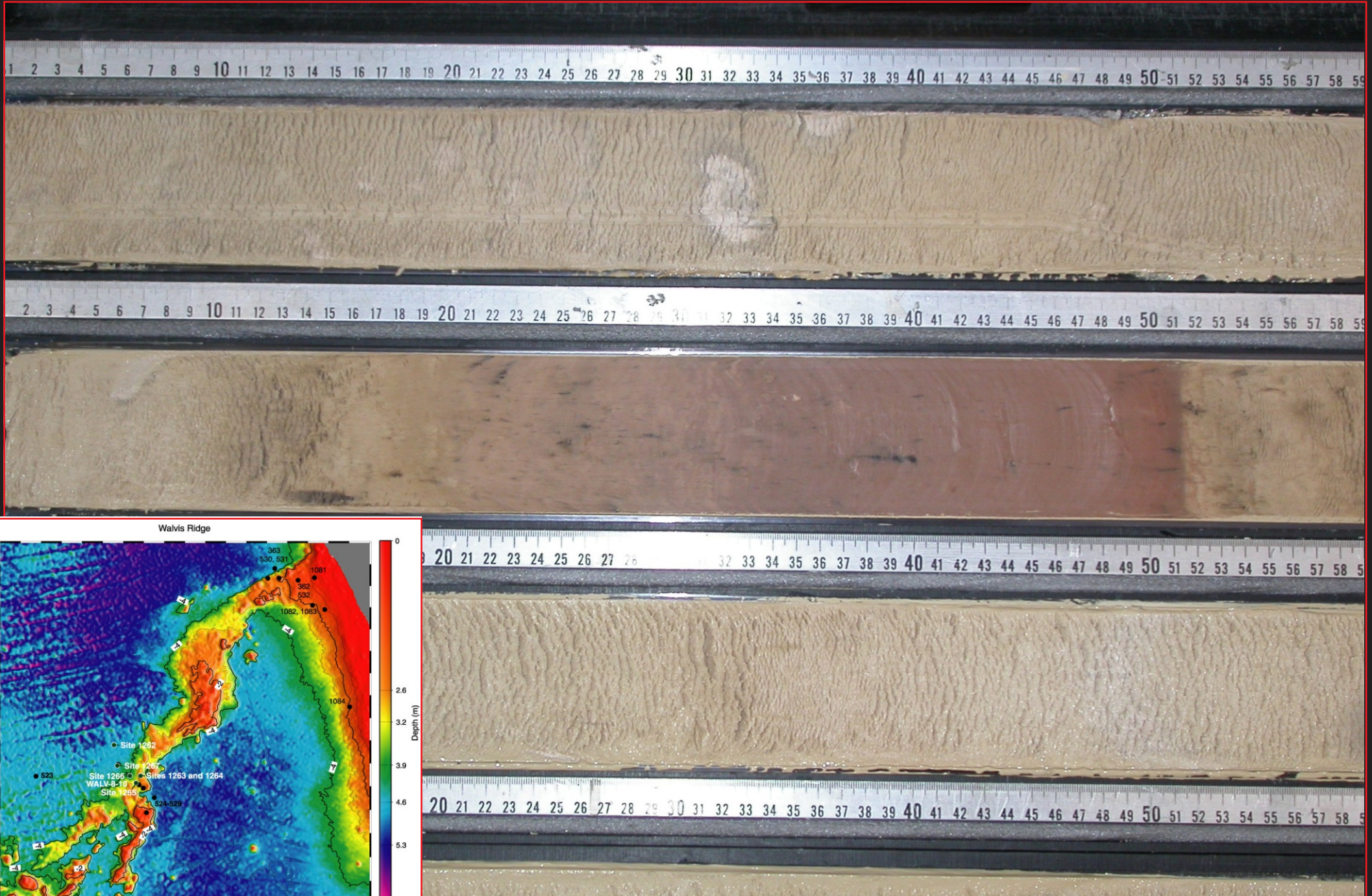


Zachos et al. [2010]
Lunt et al. [2011]

Age relative to the PETM (Ma)



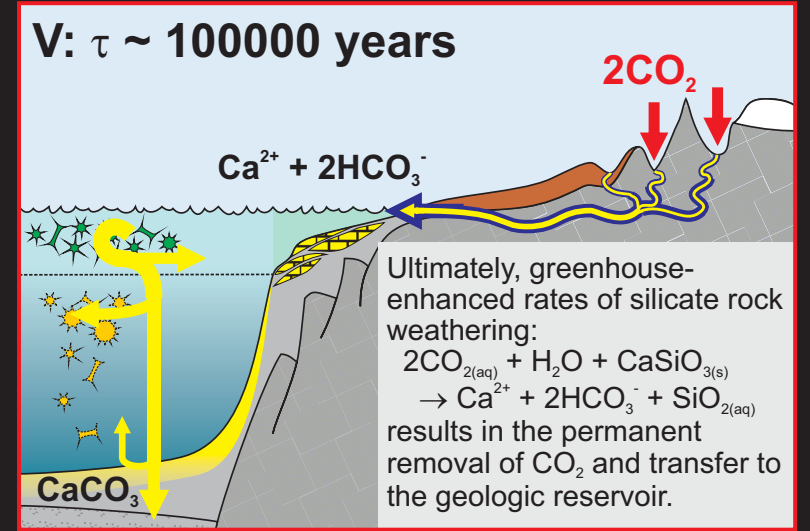
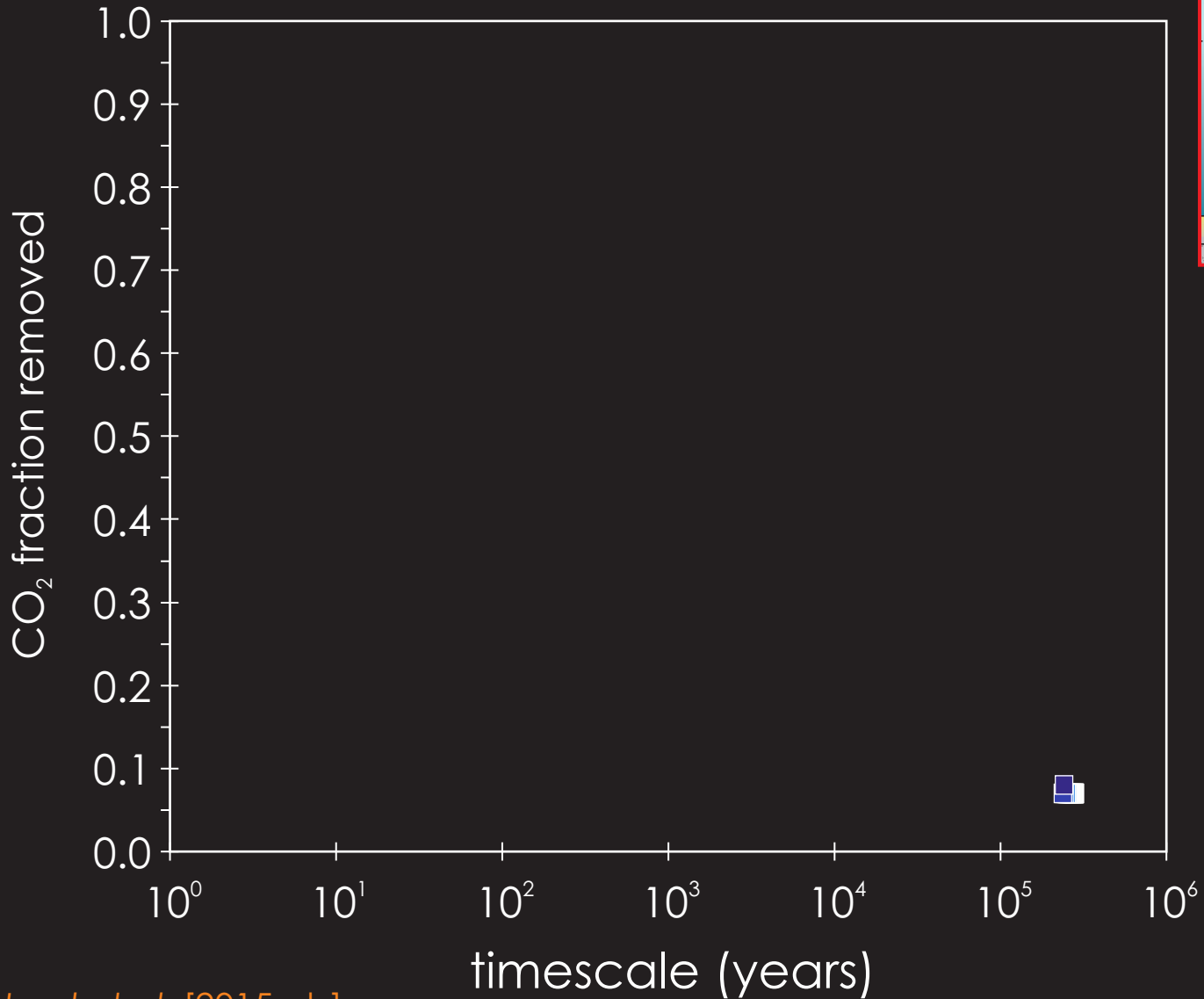
Impulse response function analysis of the 'long tail' of $\text{CO}_2(\text{excess})$



Sediments spanning the Palaeocene-Eocene boundary from ODP Leg 208 (Walvis Ridge)
Picture courtesy of Dani Schmidt (University of Bristol)

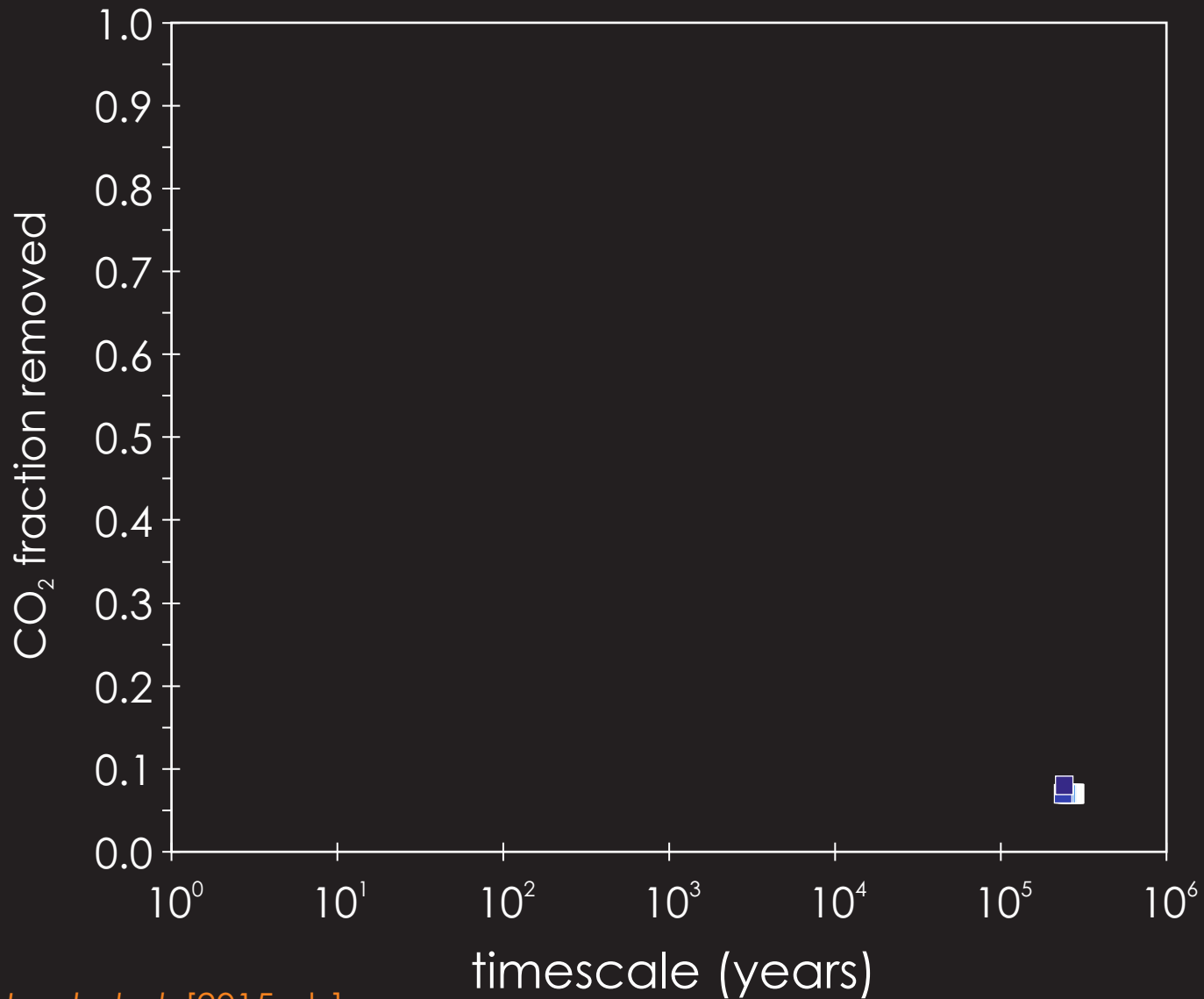


Silicate weathering (no time-scale response!).

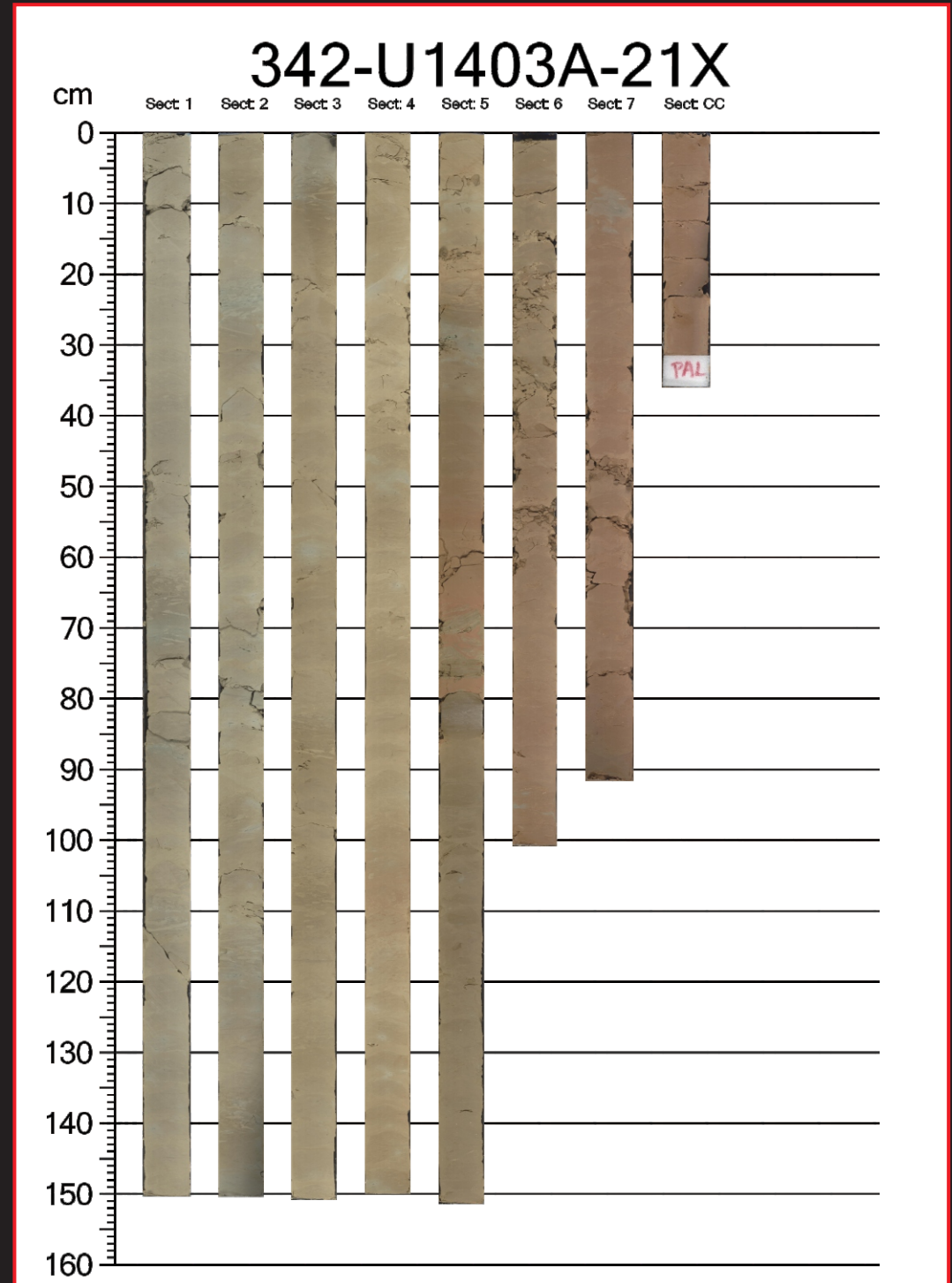
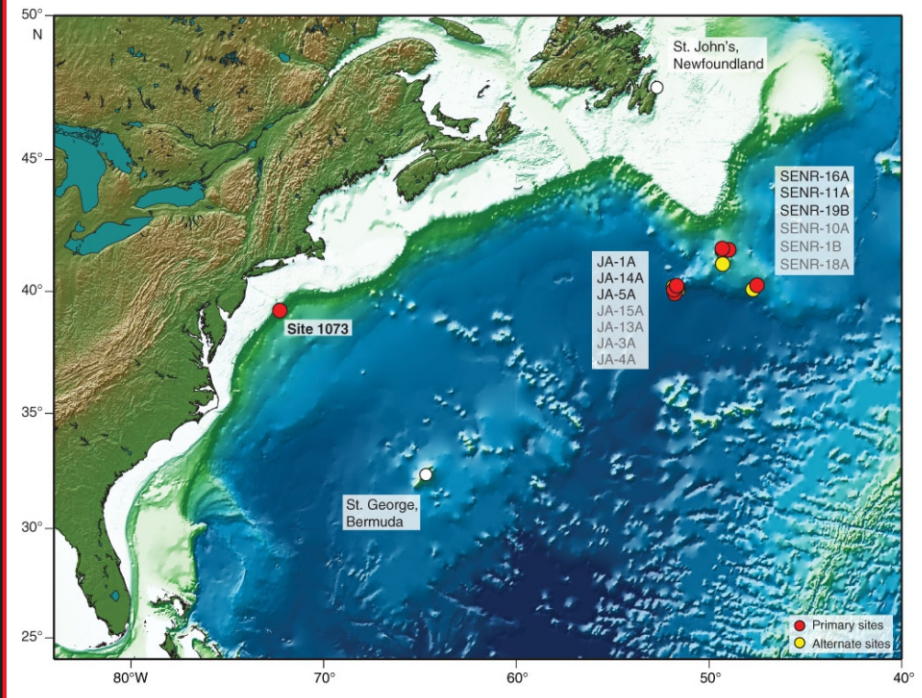




evidence?



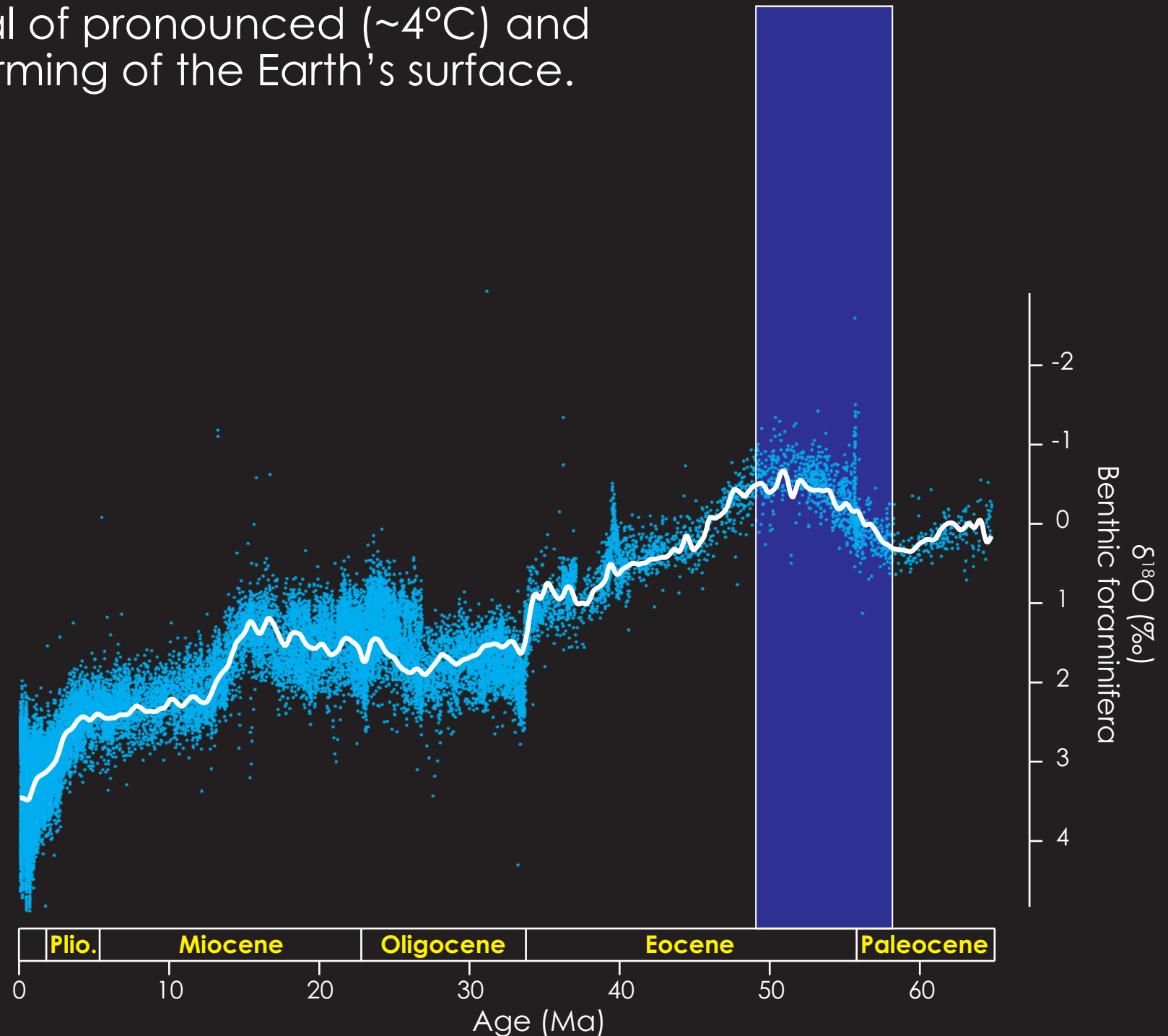
Impulse response function analysis of the 'long tail' of CO_{2(excess)}



Evidence for climate-CO₂ weathering feedback?



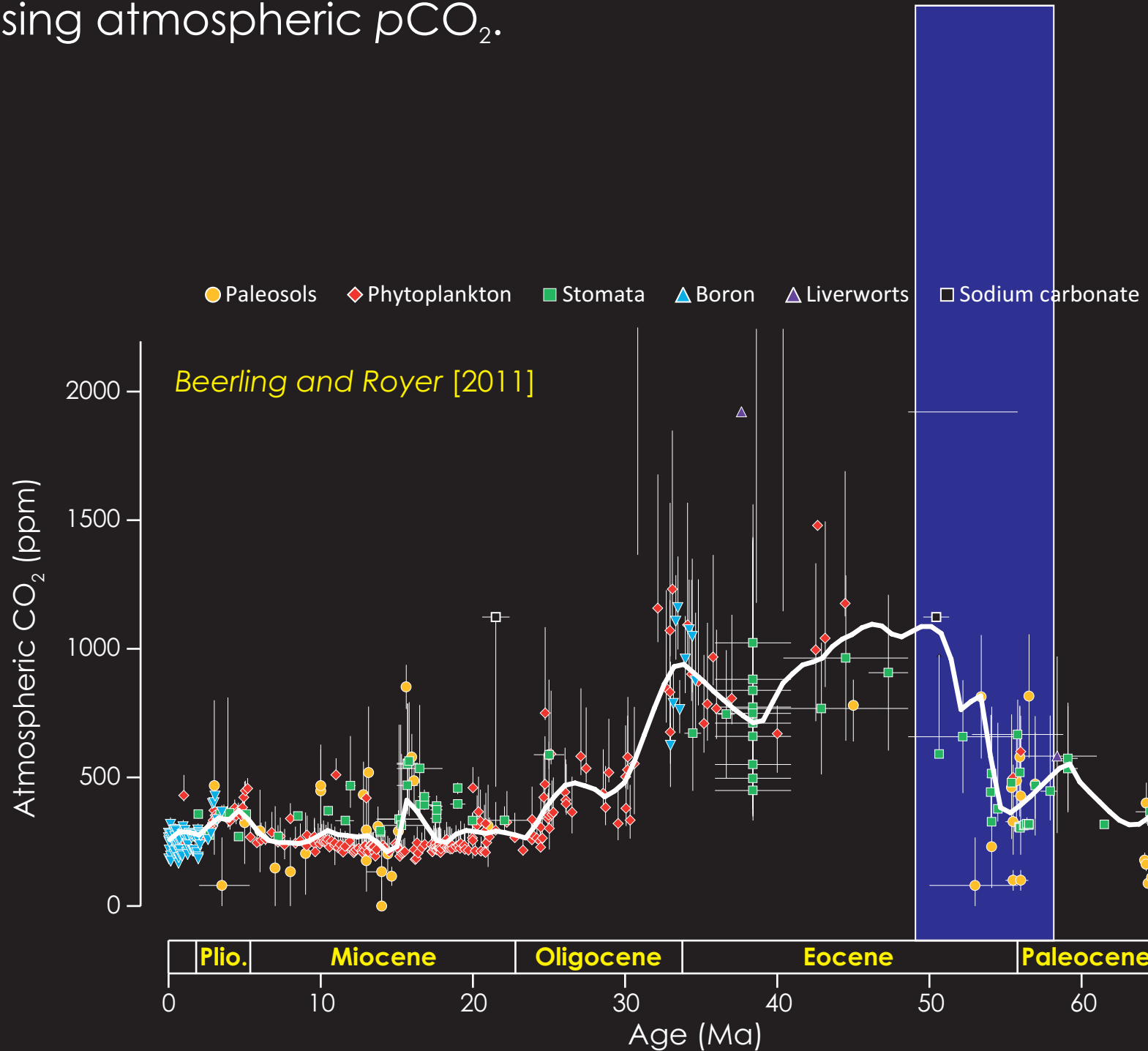
- ✓ ~9 Ma interval of pronounced (~4°C) and progressive warming of the Earth's surface.



Evidence for climate-CO₂ weathering feedback?



✓ Increasing atmospheric pCO₂.

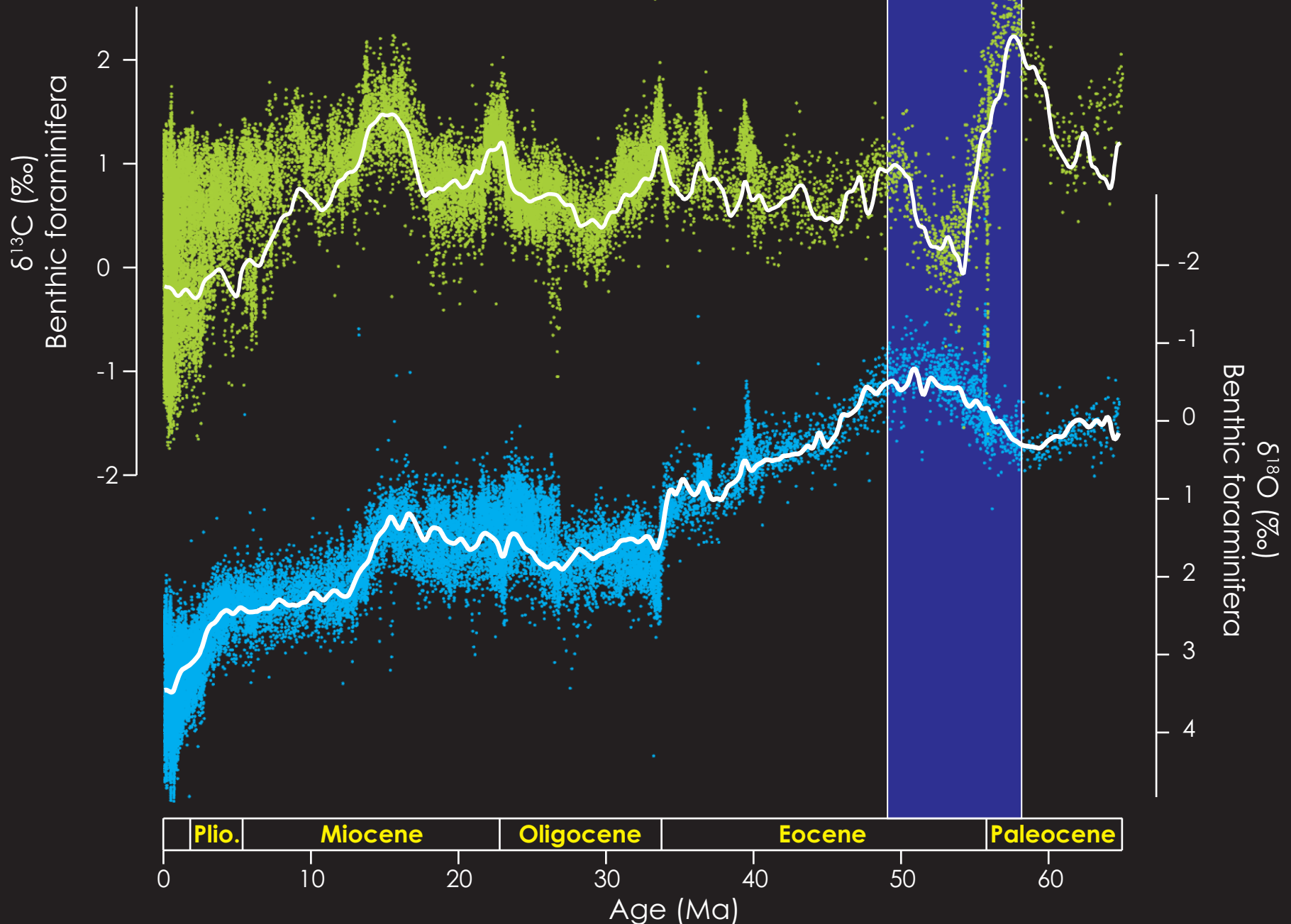


Evidence for climate-CO₂ weathering feedback?



- ✓ Mostly ... characterized by declining $\delta^{13}\text{C}$ values, consistent with net input of isotopically light carbon.

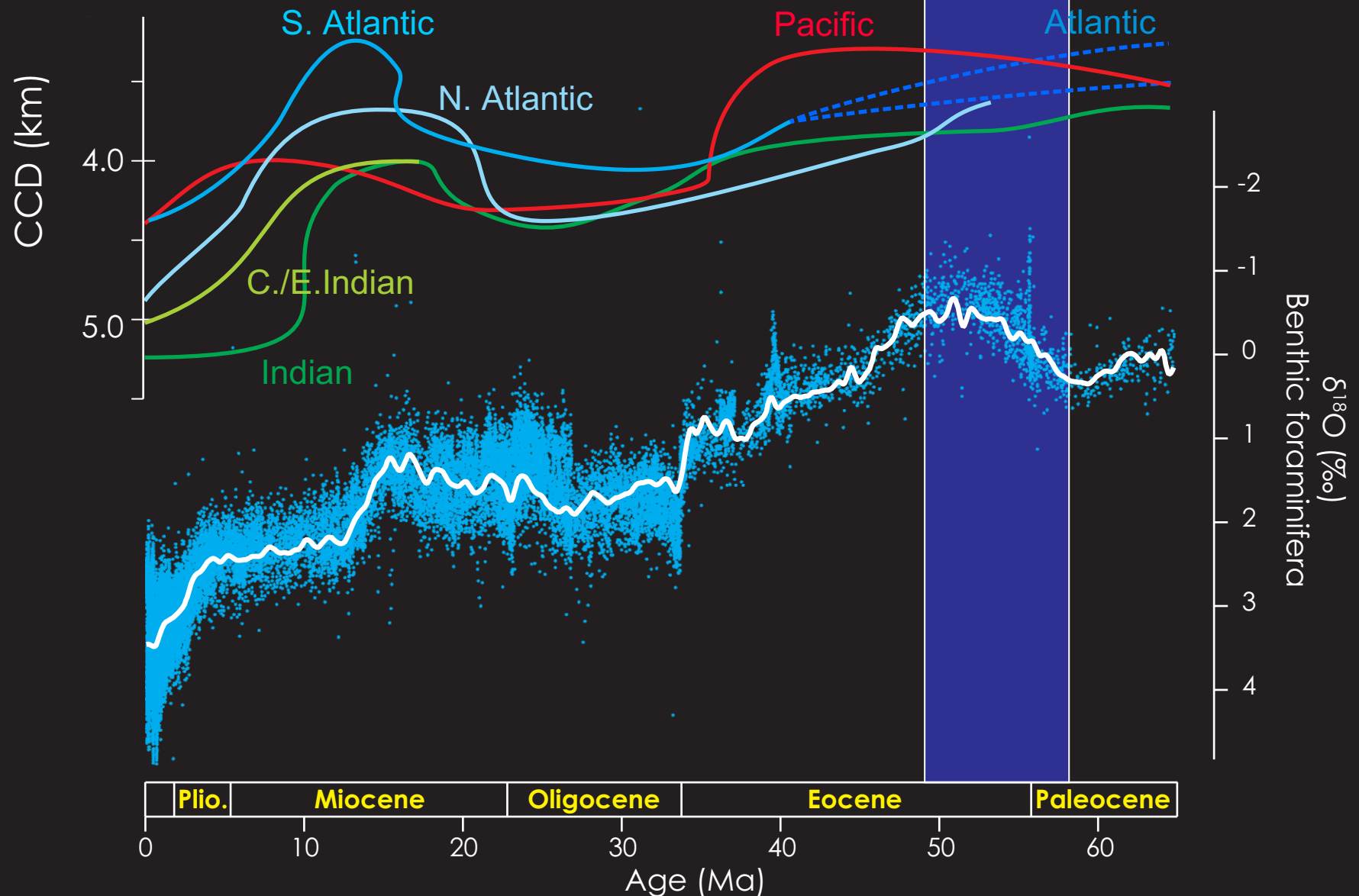
Cramer et al. [2009]



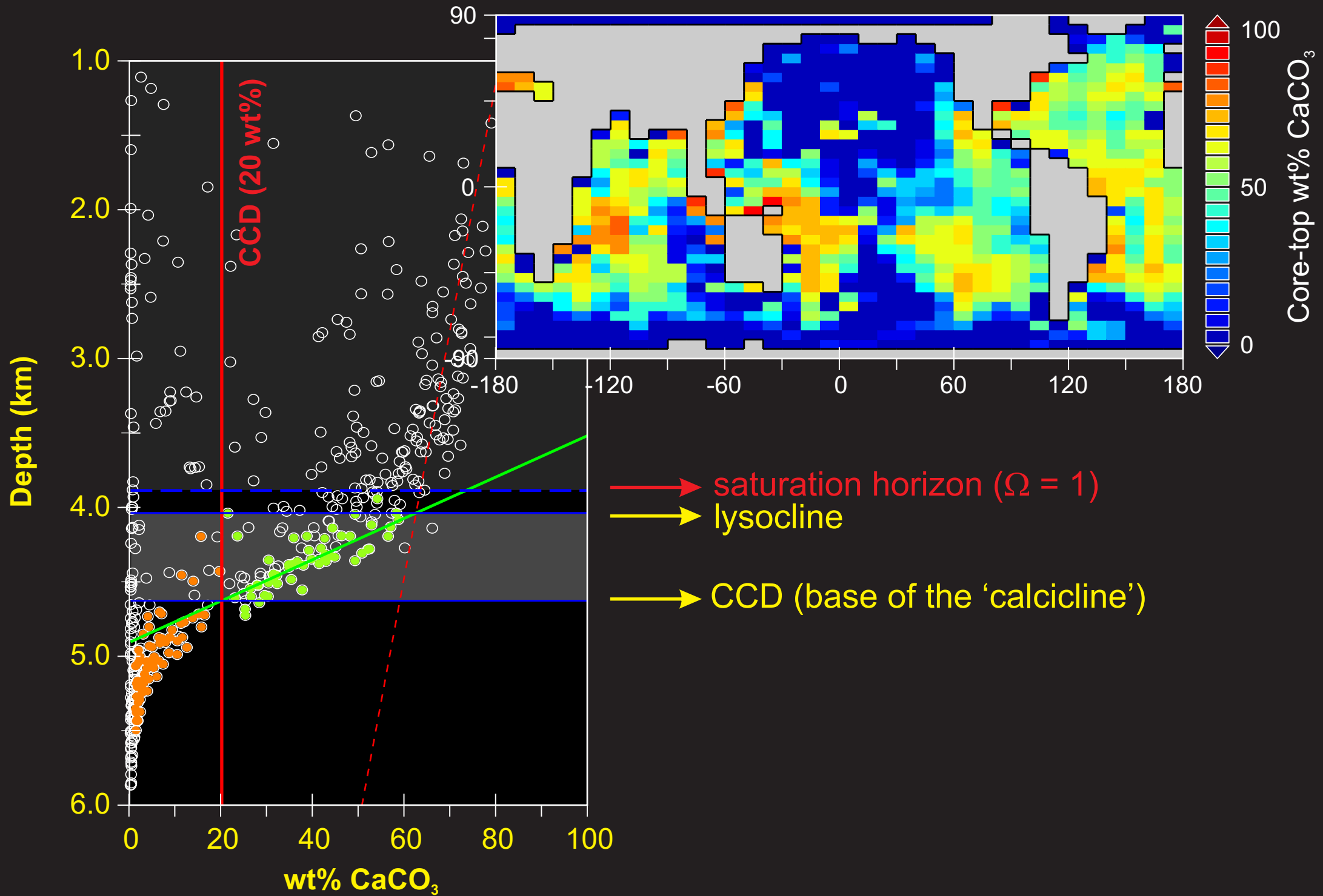
Evidence for climate-CO₂ weathering feedback?



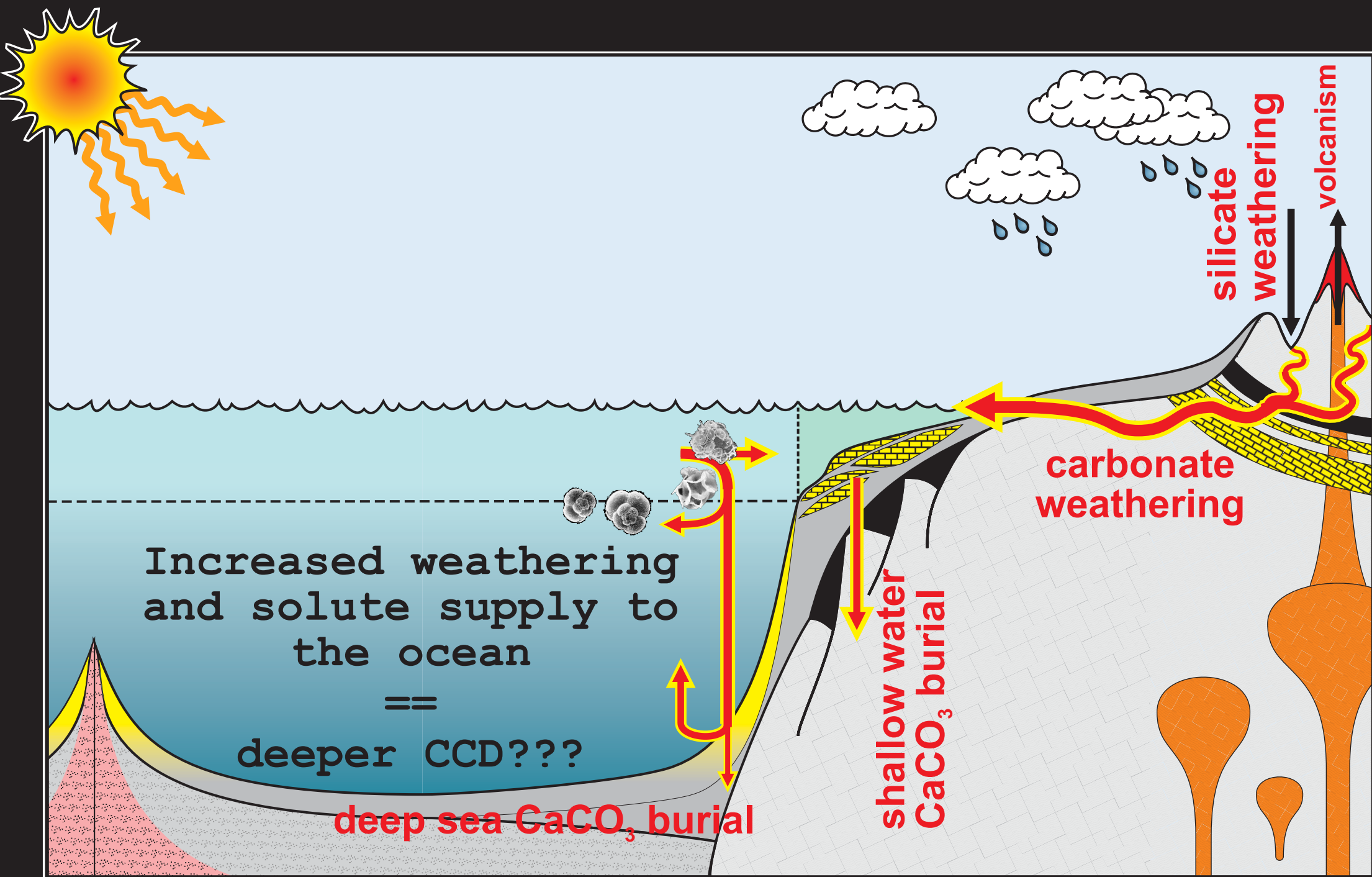
- ✗ Slightly deepening CCD ... but much less than box models predict (e.g. Komar *et al.* [2013]).
- ✗ Very sparse data coverage, not meaningfully updated since 1975.



Evidence for climate-CO₂ weathering feedback?



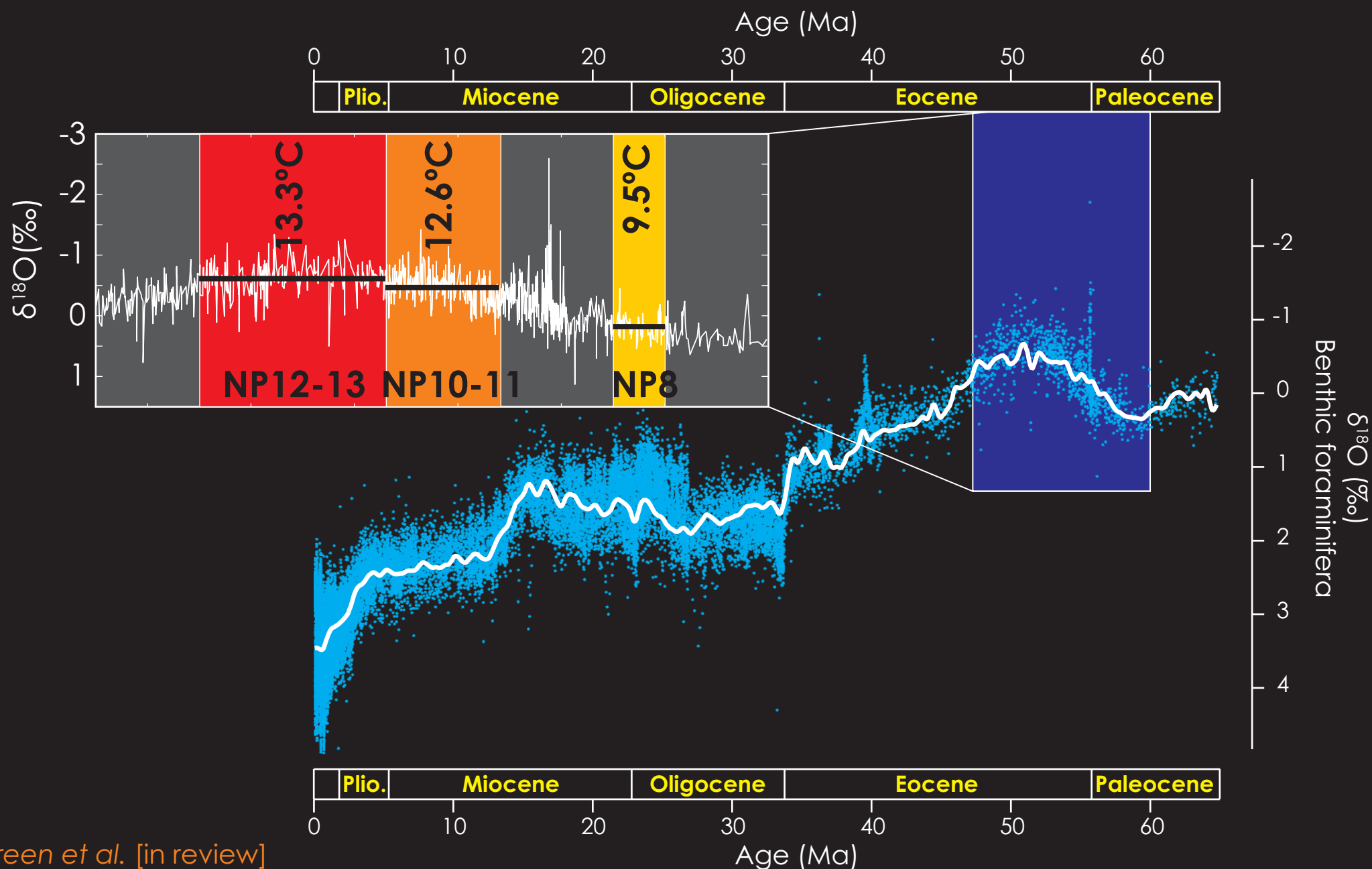
Evidence for climate-CO₂ weathering feedback?



Evidence for climate-CO₂ weathering feedback?



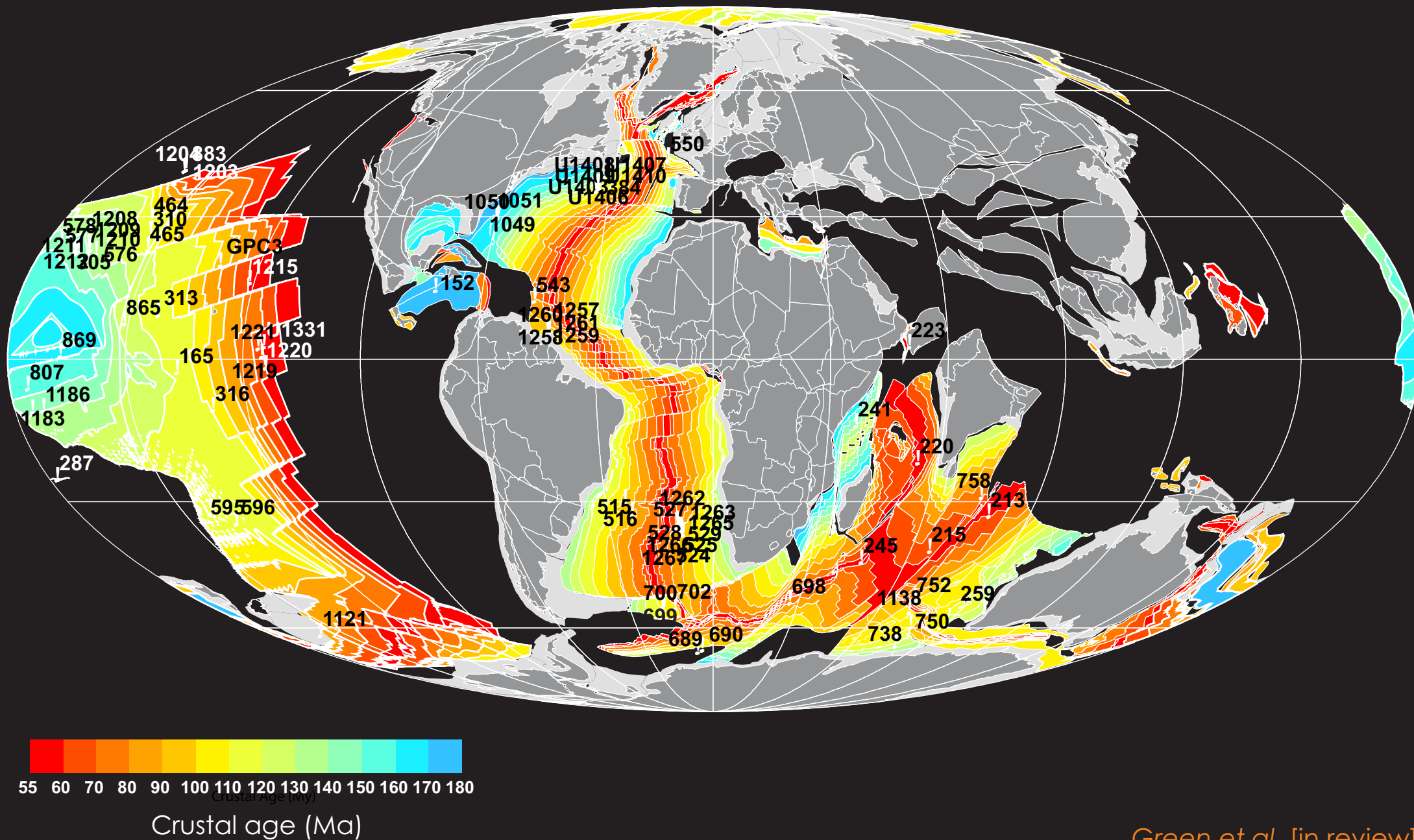
Three data slices spanning LPEE interval (and avoiding PETM).



Evidence for climate-CO₂ weathering feedback?



Site distribution (and existing crust older than 55 Ma).



Evidence for climate-CO₂ weathering feedback?



'CCD' plots.

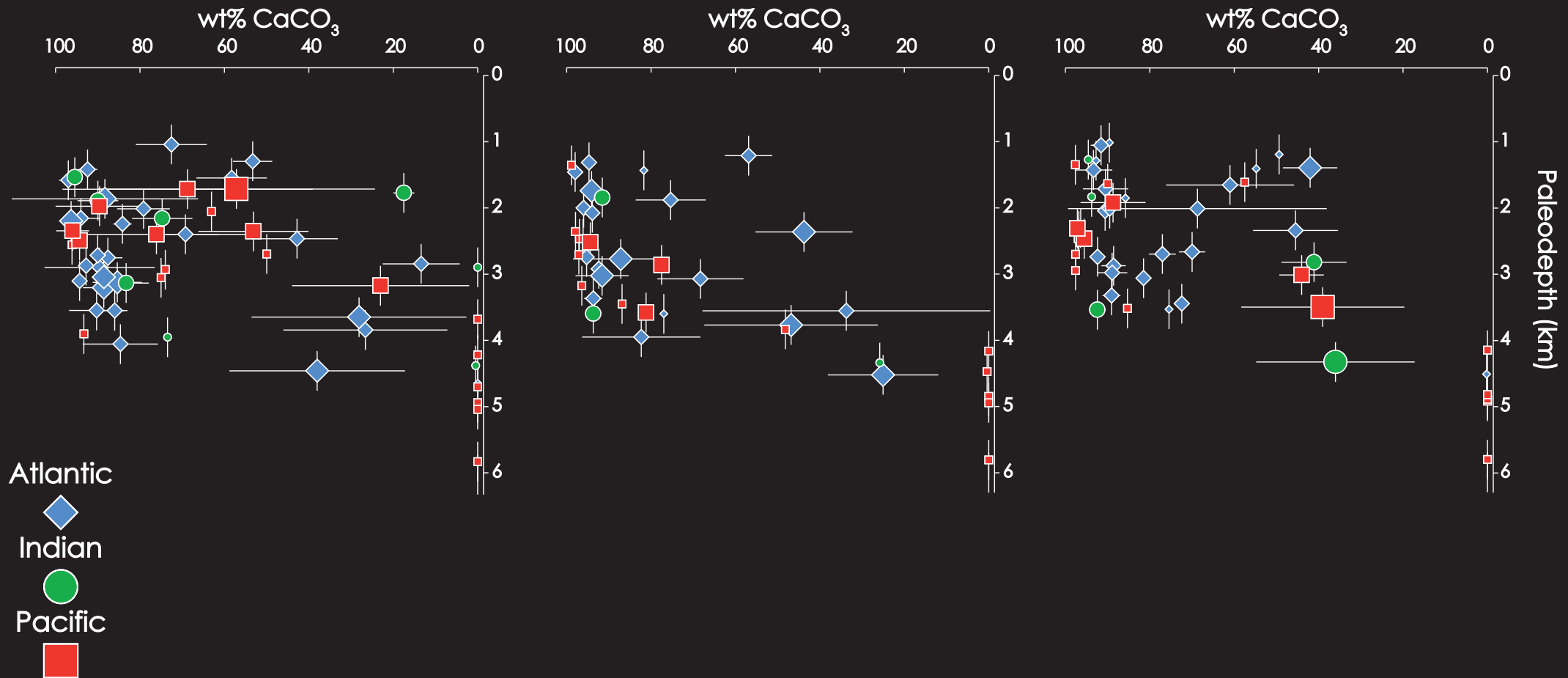
H₀: warming (=> increasing weathering?)



NP12-13 (~53-49 Ma)

NP10-11 (~55-53 Ma)

NP8 (~58-57 Ma)



Evidence for climate-CO₂ weathering feedback?



'CCD' plots.

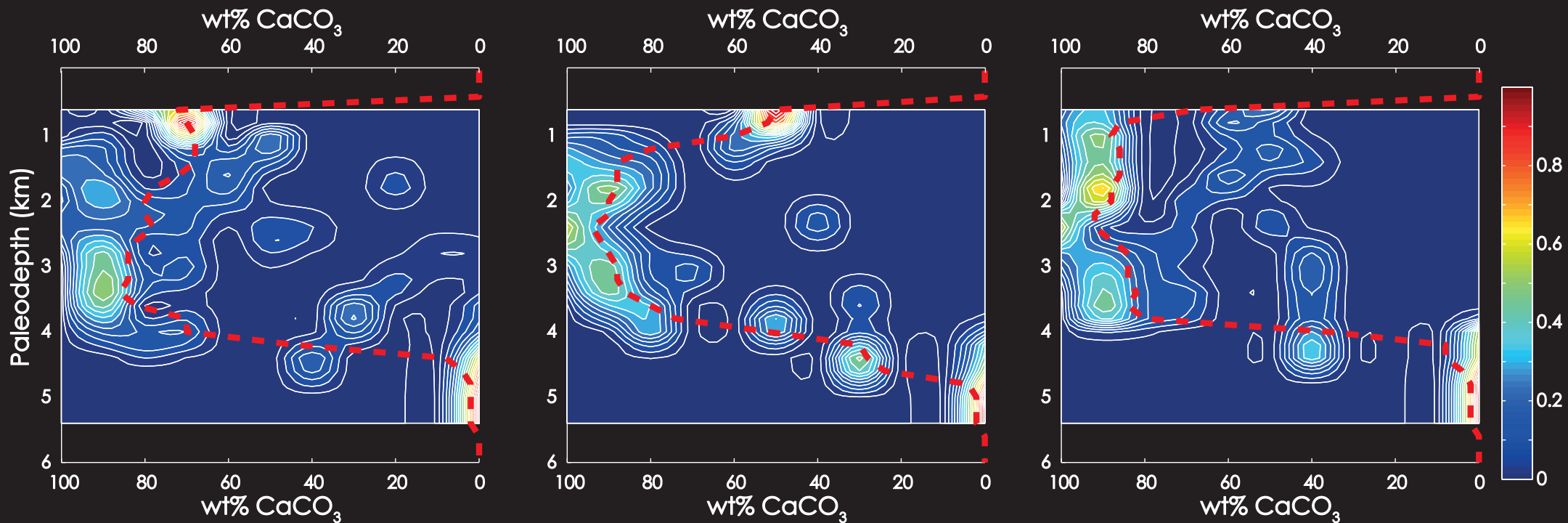
H₀: warming (=> increasing weathering?)



NP12-13 (~53-49 Ma)

NP10-11 (~55-53 Ma)

NP8 (~58-57 Ma)

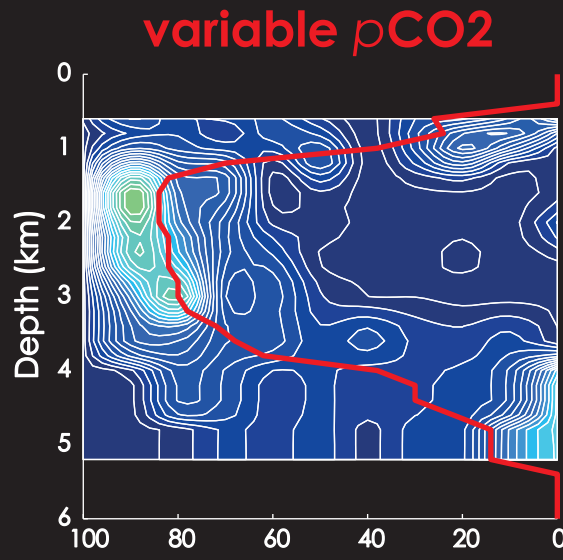


Contours are of relative data density within a sliding time-window (and wt% bin).
Red contour delineates 50% of the data.

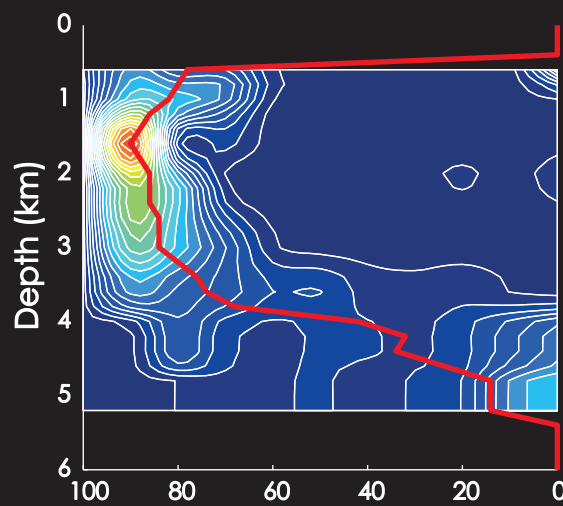
increased CO₂ out-gassing
=> higher atm pCO₂ and weathering @ steady state



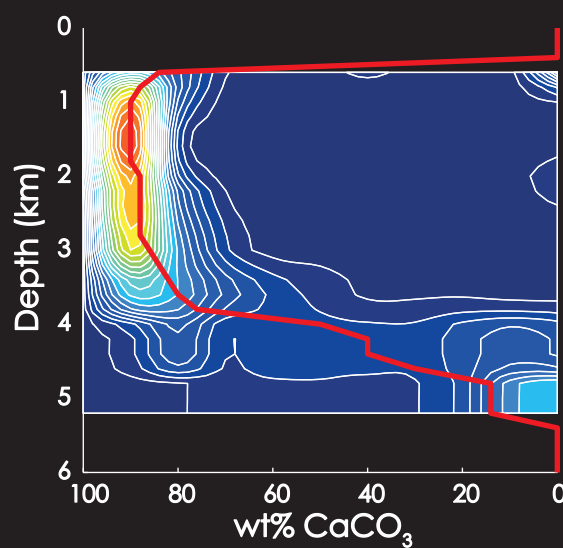
~3x
pre-industrial
pCO₂



~6x
pre-industrial
pCO₂



~12x
pre-industrial
pCO₂



anon model

```
! calculate carbonate alkalinity
loc_ALK_DIC = dum_ALK &
& - loc_H4BO4 - loc_OH - loc_HPO4 -
2.0*loc_PO4 - loc_H3SiO4 - loc_NH3 - loc_HS
&
& + loc_H + loc_HSO4 + loc_HF + loc_H3PO4

! estimate the partitioning between the
aqueous carbonate species

loc_zed = ( &
& (4.0*loc_ALK_DIC +
dum_DIC*dum_carbconst(icc_k) -
loc_ALK_DIC*dum_carbconst(icc_k))**2 + &
& 4.0*(dum_carbconst(icc_k) -
4.0)*loc_ALK_DIC**2 &
& )**0.5      loc_conc_HCO3 =
(dum_DIC*dum_carbconst(icc_k) -
loc_zed)/(dum_carbconst(icc_k) - 4.0)

loc_conc_CO3 = &
& ( &
& loc_ALK_DIC*dum_carbconst(icc_k) -
dum_DIC*dum_carbconst(icc_k) - &
& 4.0*loc_ALK_DIC + loc_zed &
& ) &
& /(2.0*(dum_carbconst(icc_k) - 4.0))

loc_conc_CO2 = dum_DIC - loc_ALK_DIC + &
& ( &
& loc_ALK_DIC*dum_carbconst(icc_k) -
dum_DIC*dum_carbconst(icc_k) - &
& 4.0*loc_ALK_DIC + loc_zed &
& ) &
& /(2.0*(dum_carbconst(icc_k) - 4.0))

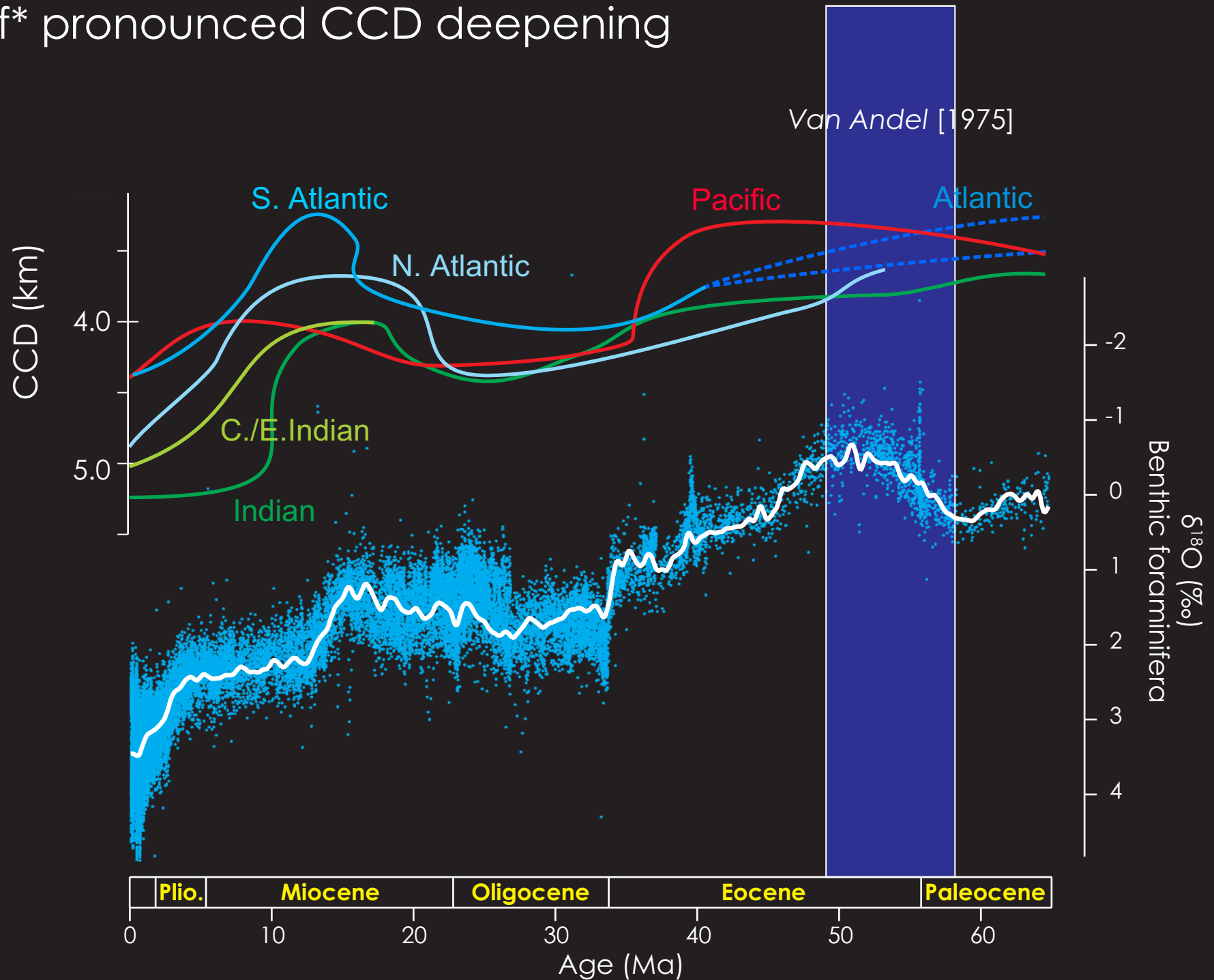
loc_H1 =
dum_carbconst(icc_k1)*loc_conc_CO2/loc_conc_
HCO3

loc_H2 =
dum_carbconst(icc_k2)*loc_conc_HCO3/loc_conc_
CO3
```

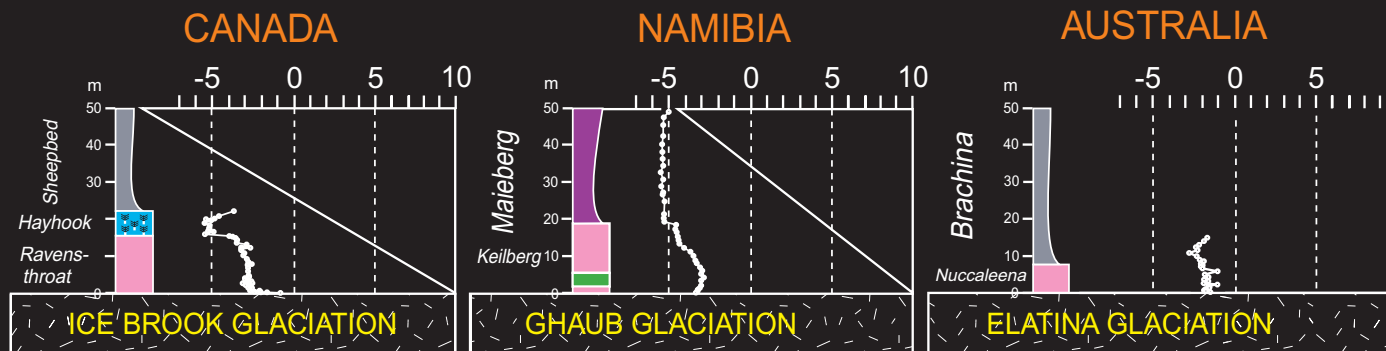
Evidence for climate-CO₂ weathering feedback?



✓ *lack of* pronounced CCD deepening

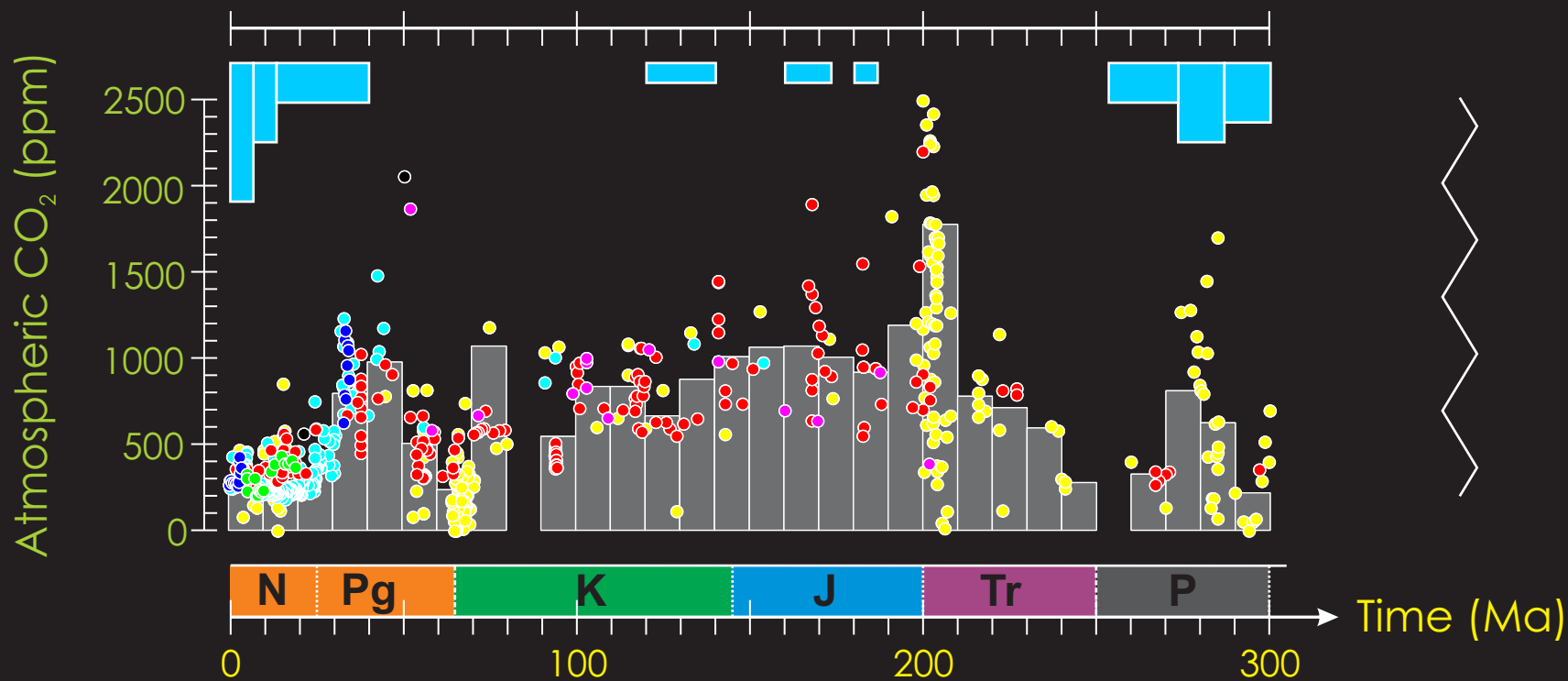


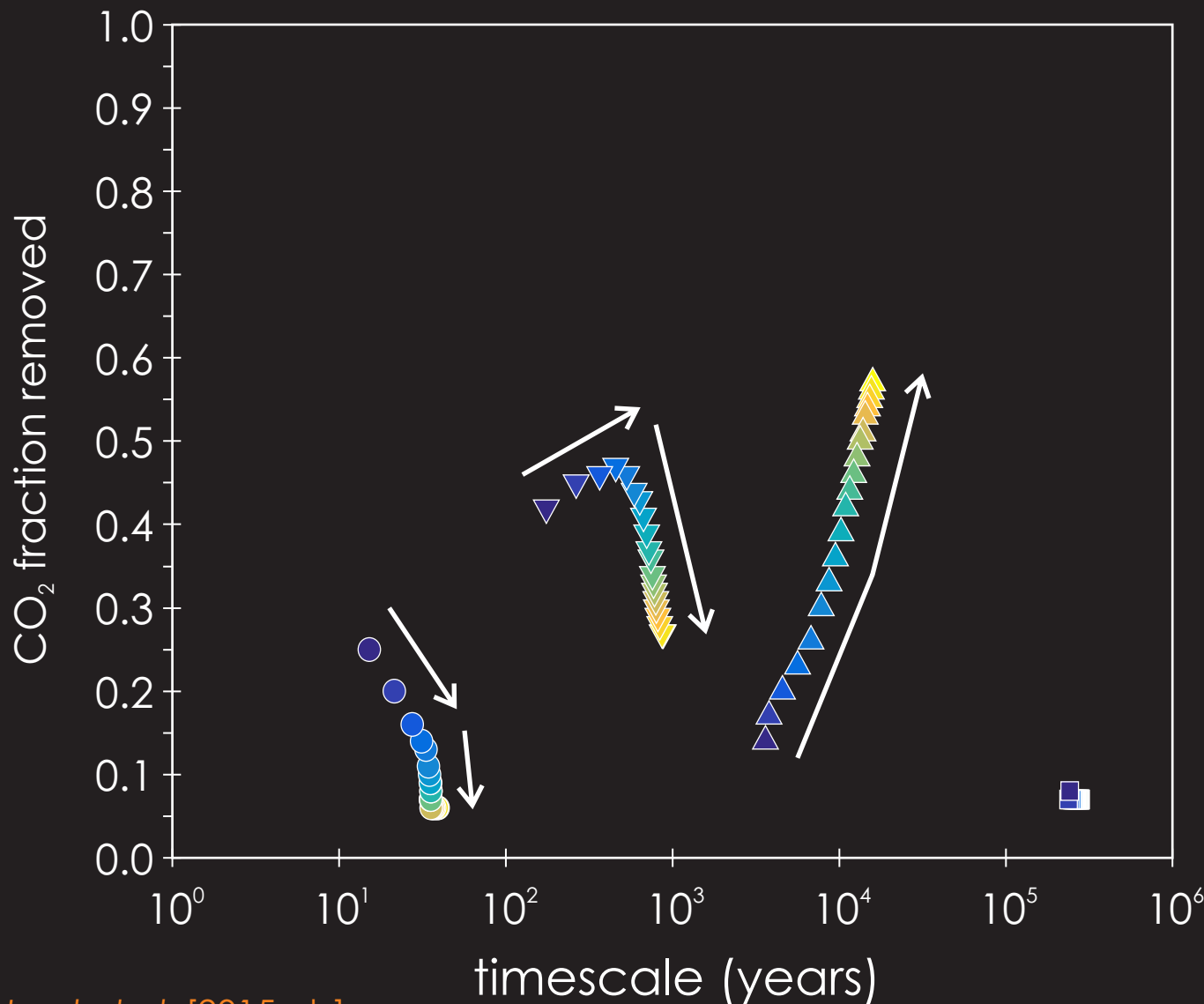
Evidence for climate-CO₂ weathering feedback?



- subaerial exposure surface
- cap limestone rhythmite
- cap limestone (cementstone)
- cap dolostone
- aeolian sandstone
- intertidal microbialaminite
- cross-bedded grainstone
- terrigenous sandstone
- columnar stromatolite
- wavy-laminated micrite
- flat-laminated micrite
- shale, siltstone

From: Hoffman and Schrag [2002]





With increasing total CO_2 emissions, the response time of all sinks (bar silicate weathering) lengthen, and the shorter time-scale two weaken at the expense of the ~10,000 year CaCO_3 burial process.

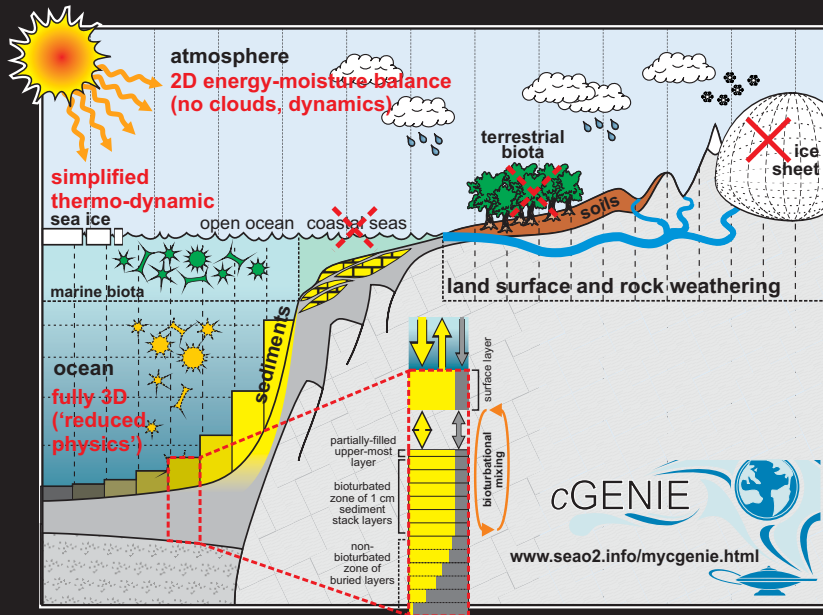
Elevated atmospheric $p\text{CO}_2$ hence becomes more persistent as the main short-term CO_2 feedbacks weaken.

The majority of carbon removal beyond ~10,000 PgC is removed only on time-scales exceeding 10,000 years.

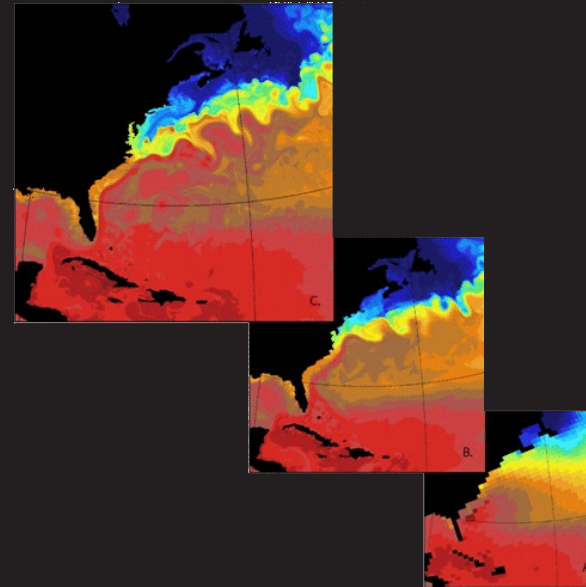
Melting Antarctica



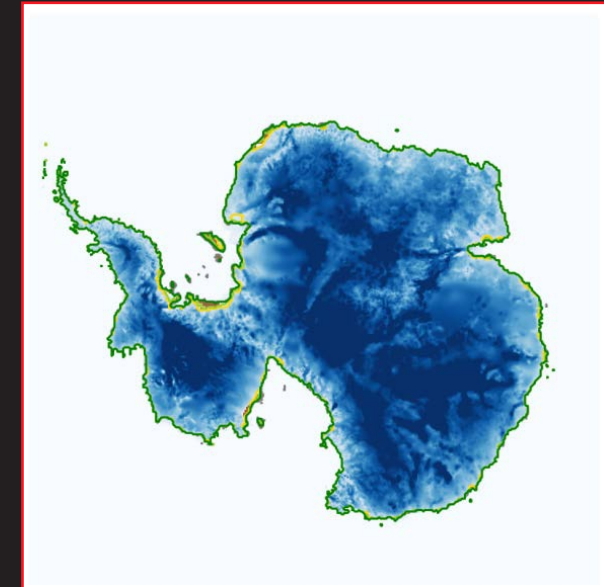
Melting Antarctica



Earth system model
(CO₂ and mean SST trajectories)

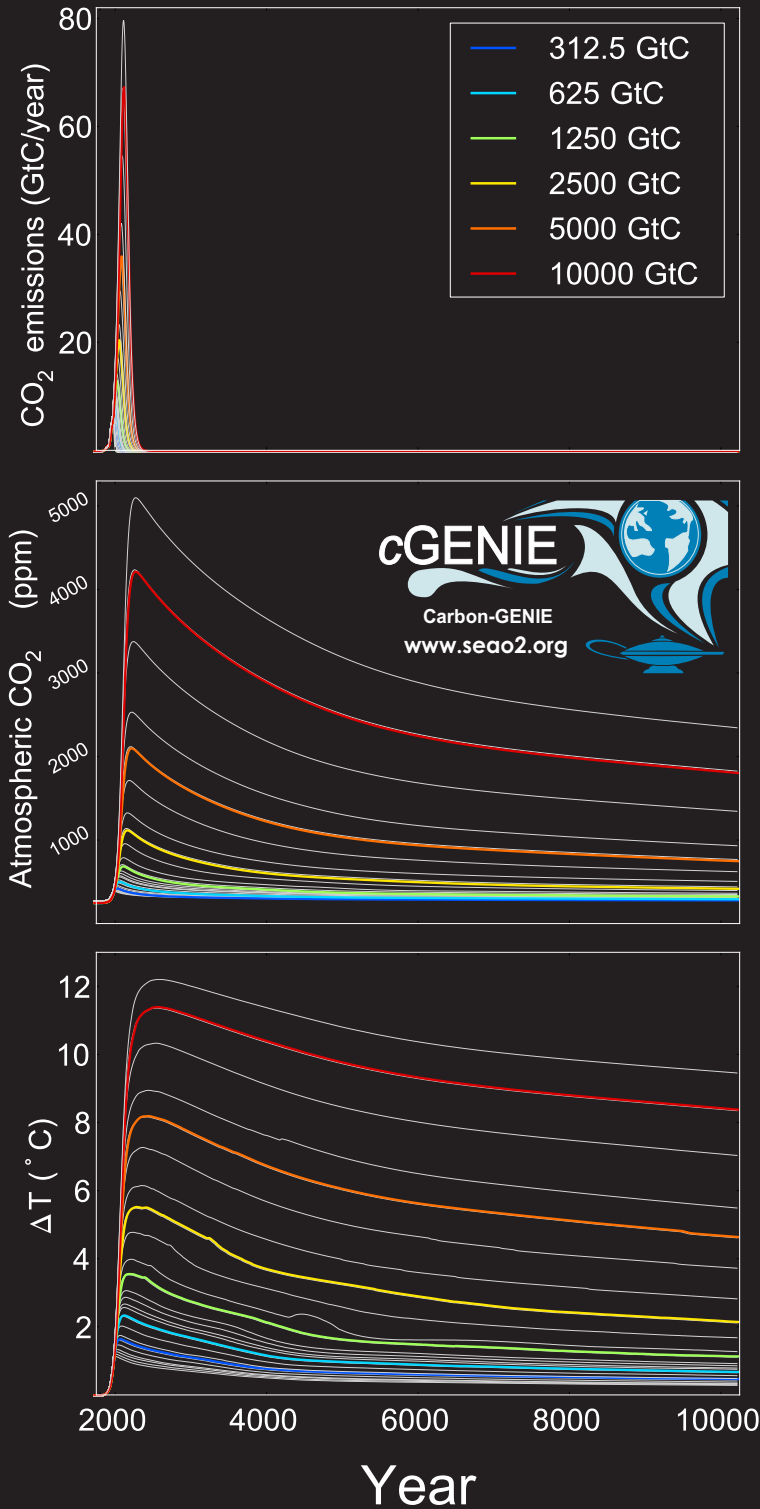


Downscaling
(SO SST and regional climate)

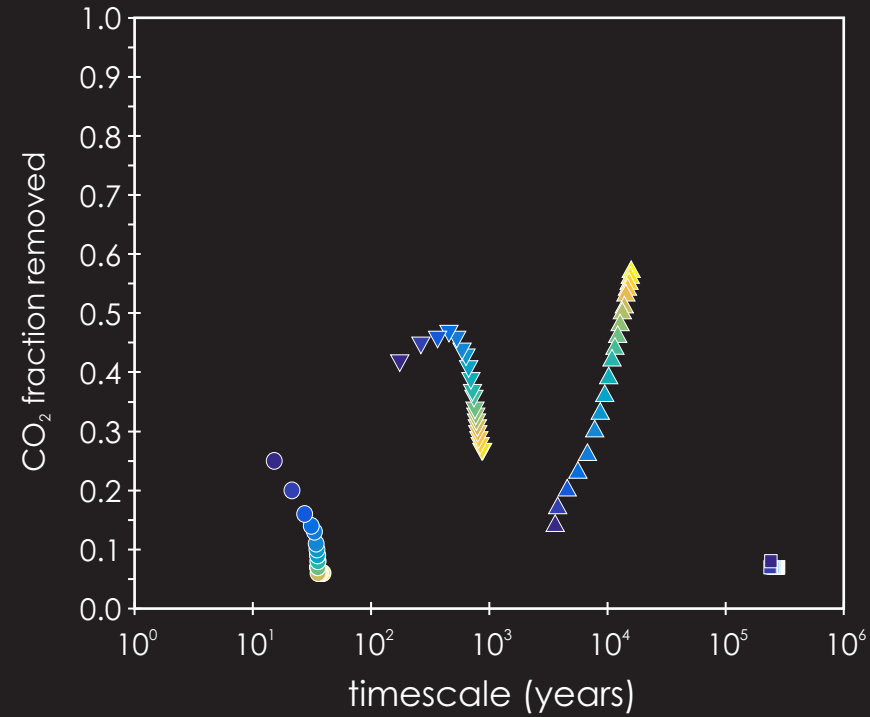
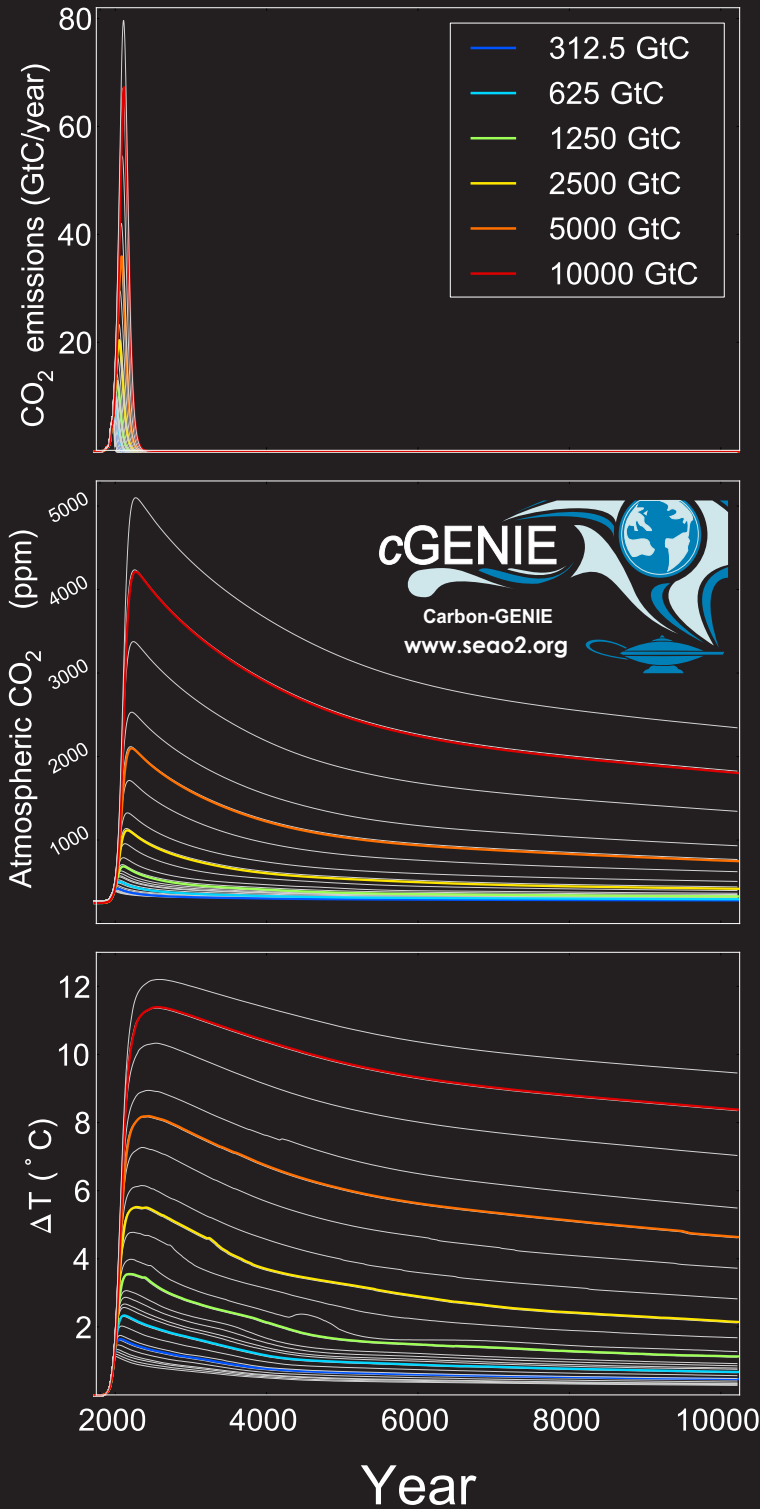


Ice sheet model

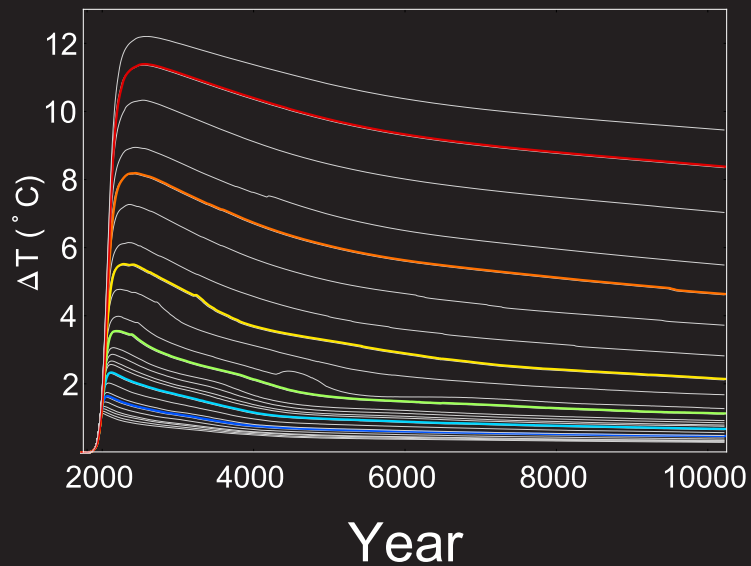
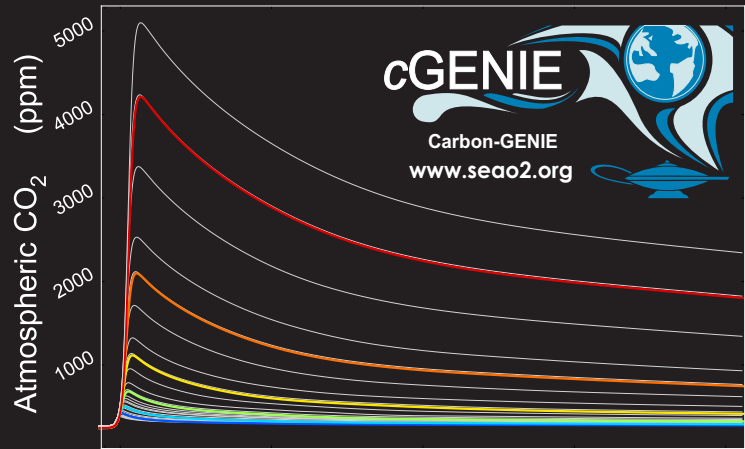
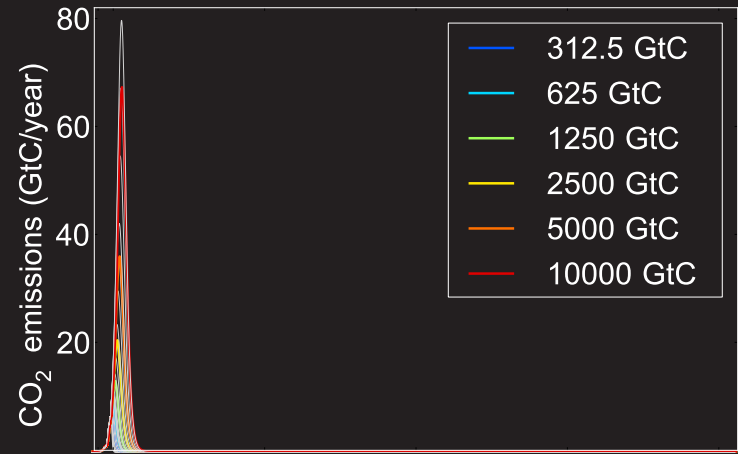
Melting Antarctica



Melting Antarctica

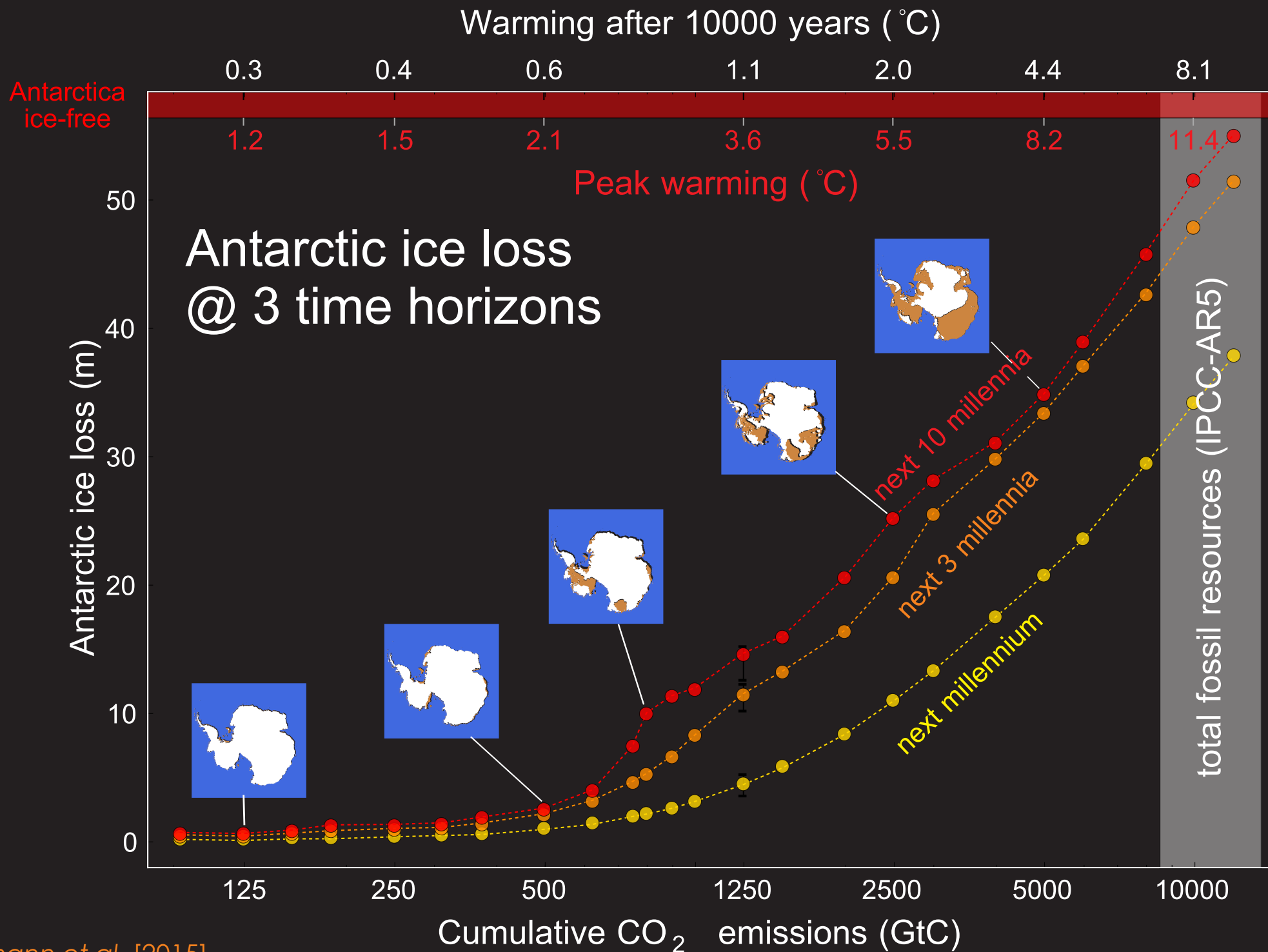


Melting Antarctica



↔ $\Delta F \propto \ln(C/C_0)$

Melting Antarctica



Enhanced weathering (CO₂ removal geoengineering)



Enhanced weathering (CO₂ removal geoengineering)



Enhanced weathering (CO₂ removal geoengineering)





granite ≈

SiO₂ = 72%

...

CaO = 1.8%

...

MgO = 0.7%

...

basalt ≈

SiO₂ = 50%

...

CaO = 10%

...

MgO = 10%

...

Enhanced weathering (CO₂ removal geoengineering)



~ plagioclase + pyroxene (+olivine)



Enhanced weathering (CO₂ removal geoengineering)



~ olivine + pyroxene

Harzburgite

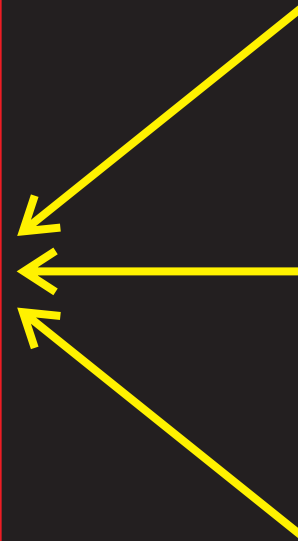


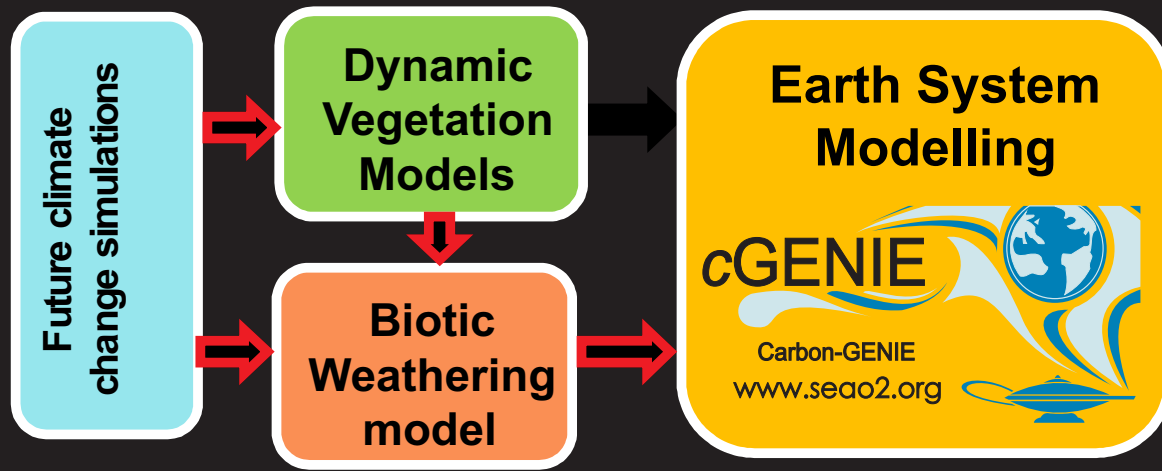
~ plagioclase + pyroxene (+olivine)



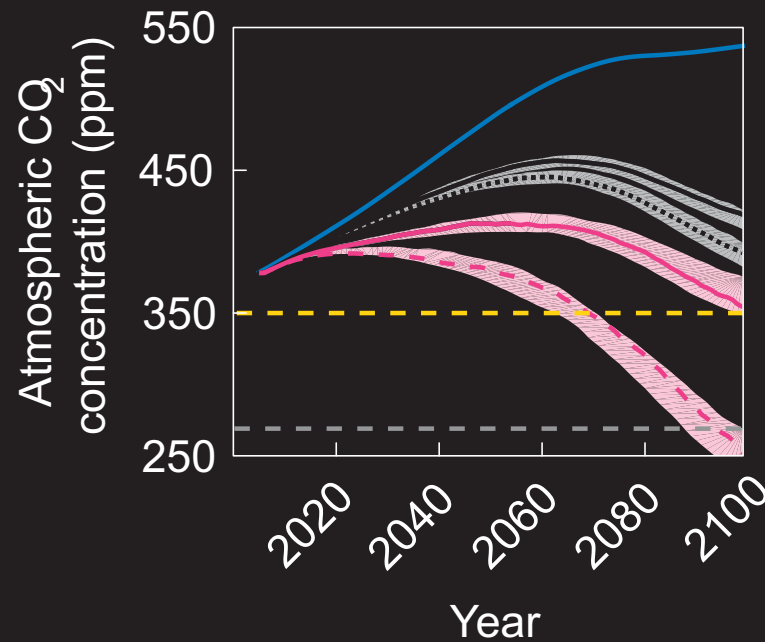
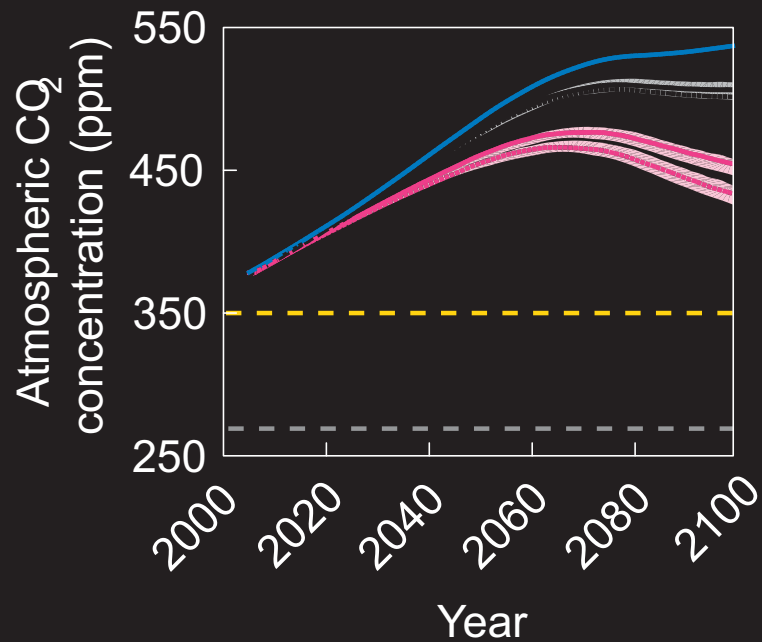
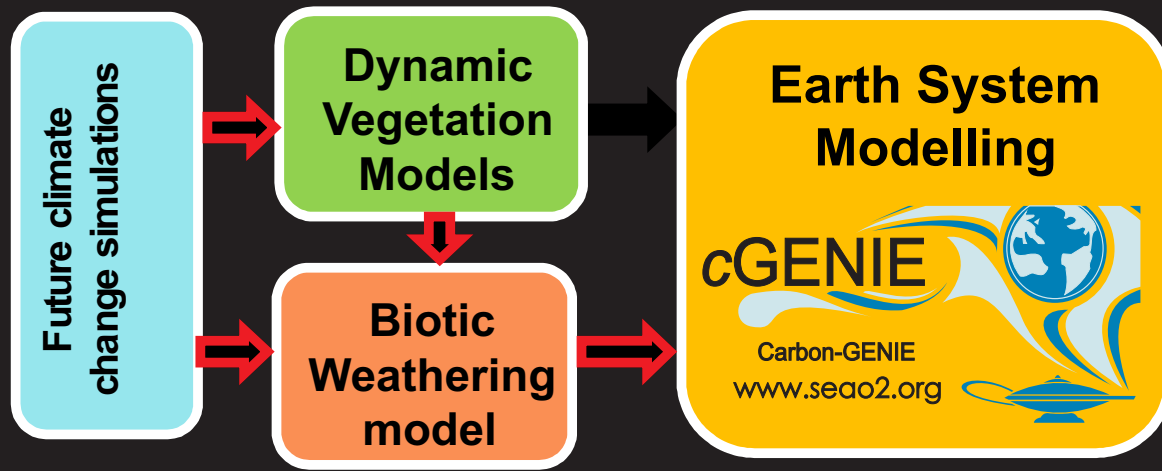
>90% olivine: $(\text{Mg}^{+2}, \text{Fe}^{+2})_2\text{SiO}_4$

Dunite



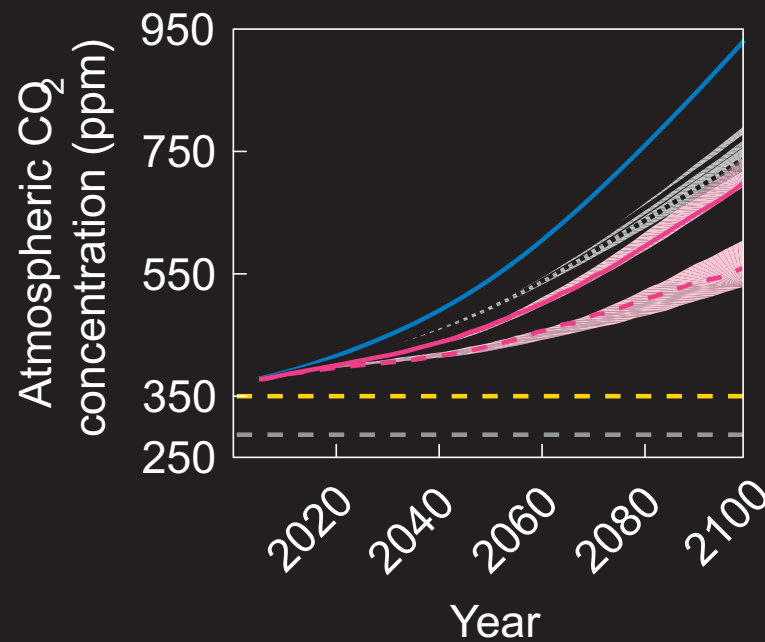
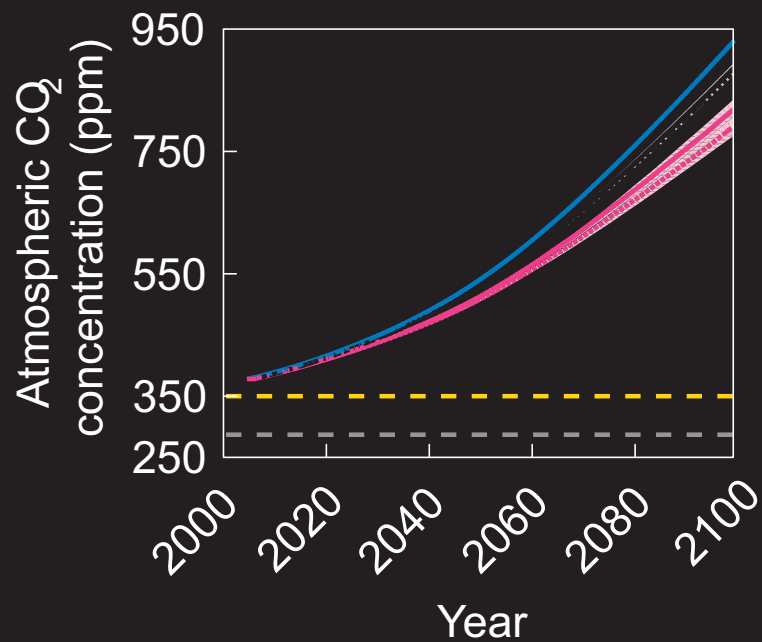
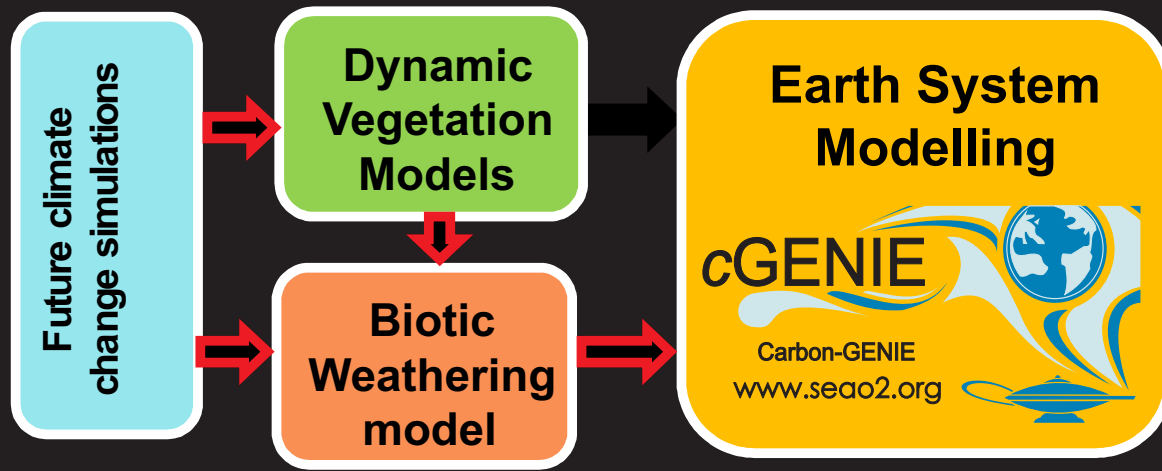


Enhanced weathering (CO₂ removal geoengineering)



- RCP scenario
- Basalt 10 cm
- Basalt 30 cm
- Harzburgite 10 cm
- Harzburgite 30 cm
- - - 350 ppm CO₂
- - - Pre-industrial CO₂

Enhanced weathering (CO₂ removal geoengineering)



- RCP scenario
- Basalt 10 cm
- Basalt 30 cm
- Harzburgite 10 cm
- Harzburgite 30 cm
- - - 350 ppm CO₂
- - - Pre-industrial CO₂



Current global oil
consumption =
90,136×10³ barrels per
day

$$\begin{aligned} 1.0 \text{ barrel} &= 159 \text{ l} \\ &= 159 \times 10^3 \text{ cm}^3 \end{aligned}$$

$$\begin{aligned} \Rightarrow \text{oil consumption} & \\ &= 5.23 \times 10^{15} \text{ cm}^3 \text{ year}^{-1} \\ &= \mathbf{5.23 \text{ km}^3 \text{ year}^{-1}} \end{aligned}$$

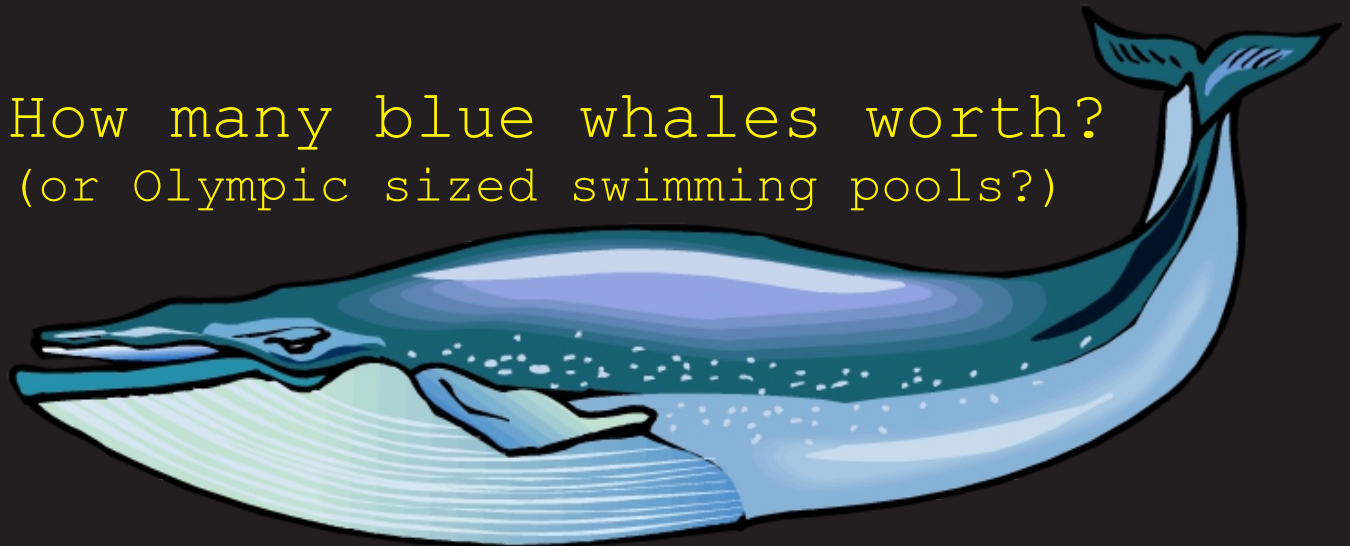


Current global oil
consumption =
90,136×10³ barrels per
day

1.0 barrel = 159 l
= 159×10³ cm³

⇒ oil consumption
= 5.23×10¹⁵ cm³ year⁻¹
= **5.23 km³ year⁻¹**

How many blue whales worth?
(or Olympic sized swimming pools?)





Current global oil
consumption =
 $90,136 \times 10^3$ barrels per
day

1.0 barrel = 159 l
= 159×10^3 cm³

⇒ oil consumption
= 5.23×10^{15} cm³ year⁻¹
= **5.23 km³ year⁻¹**

How many Yosemite Valleys?
(equivalent volume)





Current global oil
consumption =
90,136×10³ barrels per
day

$$\begin{aligned} 1.0 \text{ barrel} &= 159 \text{ l} \\ &= 159 \times 10^3 \text{ cm}^3 \end{aligned}$$

$$\begin{aligned} \Rightarrow \text{oil consumption} & \\ &= 5.23 \times 10^{15} \text{ cm}^3 \text{ year}^{-1} \\ &= \mathbf{5.23 \text{ km}^3 \text{ year}^{-1}} \end{aligned}$$

Yosemite Valley
(Wikipedia):

1,200m deep × 1,600m
across, 12.0 km long

⇒

$$\begin{aligned} \text{volume} &= 1.2 \times 1.6 \times 12.0 \\ &= \mathbf{23.0 \text{ km}^3} \end{aligned}$$

How many Yosemite Valleys?
(equivalent volume)





Thanks to:

Sarah Greene, Daniela Schmidt,

Ellen Thomas [Yale]; Sandy Kirtland Turner [UCR], et al.

Natalie Lord and Dan Lunt [Bristol], Mike Thorne [Mike Thorne and Associates Limited]

Lyla Taylor and David Beerling [Sheffield], et al.

Natural Environmental Research Council (NE/H023852/1 – ‘Evolution of Carbon Cycle Dynamics’)

AMEC and Quintessa (results part of Working Group 6 of the IAEA-sponsored MODARIA Programme)